NATIONAL INSTRUMENT 43-101 F-1 TECHNICAL REPORT ON DEL NORTE PROPERTY

LOCATED 34 KM EAST OF STEWART, BC

LATTITUDE: 56° 00' NORTH LONGITUDE: 129° 31' WEST N.T.S. MAPS 104A/4E and 104A/3W

> REPORT PREPARED FOR: DECADE RESOURCES LTD. 611 8th STREET STEWART, BC

> > BY

KRZYSZTOF MASTALERZ, Ph.D., P. Geo. krzys_mastalerz@yahoo.com

ALOJZY WALUS, M.Sc., P. Geo. alexwalus@hotmail.com

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TABLE OF CONTENTS

		Page
1.	SUMMARY	6
2.	INTRODUCTION	10
	2.1 Glossary of Technical Terms2.2 List of Abbreviations Used in the Report	11 12
3.	RELIANCE ON OTHER EXPERTS	15
4.	PROPERTY DESCRIPTION AND LOCATION	15
5.	ACCESSIBILITY, CLIMATE, LOCAL RESOURCES,	
	INFRASTRUCTURE AND PHYSIGRAPHY	16
	5.1 Accessibility5.2 Climate5.3 Local Resources & Infrastructure5.4 Physiography	16 16 16 16
6.	HISTORY	17
	6.1 Prior Ownership and Ownership Changes6.2 Summary of Previous Exploration Programs	17 19
7.	GEOLOGICAL SETTING AND MINERALIZATION	22
	 7.1 Regional Geology 7.2 Local Geology 7.3 Structure and Principal Zones 7.4 Mineralization 	22 27 32 39
8.	DEPOSIT TYPE	46
	8.1 Deposit Type8.2 Geological Model and Concepts Used	46 47
9.	EXPLORATION	49
10.	PREVIOUS DRILLING AND GEOPHYSICS	55
	10.1 Drilling 10.2 Geophysics	55 76

11.	SAMPLE PREPARATION, ANALYSES AND SECURITY	82		
12.	DATA VERIFICATION	84		
13.	MINERAL PROCESSING AND METALURGICAL TESTING	85		
14.	MINERAL RESOURCE ESTIMATE	85		
15.	MINERAL RESERVE ESTIMATES	85		
16.	MINING METHODS	85		
17.	RECOVERY METHODS	85		
18.	PROJECT INFRASTRUCTURE	85		
19.	MARKET STUDIES AND CONTRACTS	85		
20.	ENVIRONMENTAL STUDIES, PERMITTING AND SOCIAL OR			
	COMMUNITY IMPACT	85		
21.	CAPITAL AND OPERATING COST	85		
22.	ECONOMIC ANALYSIS	85		
23.	ADJACENT PROPERTIES			
24.	OTHER RELEVAN DATA AND INFORMATION	87		
25.	INTERPRETATION AND CONCLUSIONS	87		
26.	RECOMMENDATIONS	88		
27.	REFERENCES	95		
28.	DATE AND SIGNATURE PAGE	100		
29.	CERTIFICATE OF QUALIFIED PERSON	101		
LIST	OF TABLES			
Table	1 Selected drill results from LG Vein and Kosciuszko Zone			
Table 2	2. Claims information			
Table 3	3. Decade 2019 exploration – Sample locations and descriptions			
Table 4. Decade 2019 exploration - Laboratory results (selected eleme				

by Bishop and Gal, 1991: converted to metric system)

(according to: Cremonese, 2003).

Selected significant 1990 drill-hole intercepts (based on original data presented

Significant results of the 2002 diamond drilling program on the Kosciuszko Zone

Table 5.

Table 6.

- Table 7. Significant results of the 2003 diamond drilling program on the LG vein (according to: Cremonese, 2004).
- Table 8. Significant results of the 2004 diamond drilling program on the LG vein (according to: Mastalerz and Cremonese, 2004).
- Table 9. Significant results of the 2005 diamond drill program on the LG Vein area (Cremonese, 2020; unpublished file).
- Table 10. Significant results of the 2006 drill program on the Del Norte property (according to Teuton News Release of December 18th, 2006.
- Table 11. Significant results of the 2014 drill program (according to Cremonese, 2015)
- Table 12. Significant results of the 2016 diamond drill program on the Del Norte property (according to Cremonese, 2017).
- Table 13. Significant intersections of the drill holes DN-2019-1 and DN-2019-2
- Table 14. Exploration work recommended for the Del Norte property and its estimated costs.

LIST OF FIGURES

- Figure 1. Del Norte Property Location Map.
- Figure 2. Del Norte Property Claim Map.
- Figure 3. Del Norte Property Mineral Showings.
- Figure 4. Regional Geology Stikine Terrane..
- Figure 5. Regional Geology Terranes and Project Location
- Figure 6. Del Norte Property Local Geology Map.
- Figure 7. Del Norte Property Simplified Cross Section.
- Figure 8. LG Vein 2004 Drill Hole Plan and Selected Drill Hole Cross-Sections
- Figure 9. "LG Vein Extension" Drill Hole Cross-Section Pad Ext 1 (2004)
- Figure 10. Humdinger Zone Drill Hole Cross-Section Pad 2 (2016)
- Figure 11. LG Vein Kosciuszko Zone (with 2002-2006 drill-dole projections and preliminary contours of 2005 airborne EM anomaly
- Figure 12. Hardpan Creek Area Mineral Showings/Occurrences
- Figure 13. 2019 Decade Exploration Location of Sampled Areas
- Figure 14. 2019 Decade Exploration Sample Locations Map 1
- Figure 15. 2019 Decade Exploration Sample Locations Map 2 (SP Vein Area)
- Figure 16. Hardpan Creek Area Drill Hole Locations (1990)
- Figure 17. Hardpan Creek Area Drill Hole Cross-Sections DDH90-1, 2 and DDH90-3, 4 (1990).
- Figure 18. Kosciuszko Zone Drill Hole Cross-Section DN02-1 and DN02-2 (2002)
- Figure 19. LG Vein 2003 Drill Hole Locations and Results

- Figure 20. LG Vein 2004 Drill Hole Locations
- Figure 21. LG Vein 2004 Drill Hole Cross-Sections (Pad C and K)
- Figure 22. Hardpan Creek Area 2014 Drill Hole Locations and D14-01 D14-02 Cross-Section
- Figure 23. Hardpan Creek Area 2016 Drill Hole Locations and DN16-03, DN16-04, DN16-05 Cross-Section
- Figure 24. LG Vein Area 2016 Drill Hole Locations and DN16-10 through DN16-13 Cross-Section
- Figure 25. LG Vein Area 2019 Drill Hole Locations and Cross-Section DN2019-01
- Figure 26. Del Norte 2005 Airborne Survey Total Magnetic Field (TMI), Del Norte Valley
- Figure 27. Del Norte 2005 Airborne Survey AeroTEM Off-Time Profiles, LG Vein Kosciuszko Zone Area
- Figure 28. Del Norte 2019 Airborne Survey ZTEM Inversion results by Geotech, LG Vein Area
- Figure 29. Core Recovery vs. Gold/Silver Grade 2004 Drill Program, LG Vein.
- Figure 30. Interpreted Northern Extension of LG Vein (and recommended drill hole locations)

LIST OF PHOTOGRAPHS

- Photo 1. Microscope photomicrograph of the Kosciuszko Zone mineralization in reflected light (a core sample).
- Photo 2. SP vein area sampled by the authors in autumn 2019 (view from the "Crest Zone"; looking northwestward).

APPENDICES

Appendix 1. Laboratory certificates: analyses of 2019 rock samples collected by the authors on the SP Zone area, Del Norte property.

1. SUMMARY

1.1 PROPERTY DESCRIPTION AND LOCATION

The Del Norte property is located 34 km east of Stewart in the Skeena Mining Division. It is centered on Latitude 56° 00' North and Longitude 129° 31' West on NTS map sheets 104A/4E and 104A/3W. The property consists of 13 mineral tenures (claims) covering 5830 hectares which are 100 percent owned by Teuton Resources.

1.2 ROYALITIES AND AGREEMENTS

Decade Resources has the right to earn up to a 55 % interest under the following terms:

- Payment of \$400,000 over 4 years with an initial payment of \$20,000.
- Issuing 800,000 shares of Decade on signing.
- Issuing \$180,000 of Decade stock over a 4-year period.
- Expenditures of \$4,000,000 over 5 years.

Decade has the right to earn an additional 20% by placing the property into production.

The property is subject to the following royalties (the "Pre-Existing Royalties"): one percent net smelter royalty in favour of Sabina Gold & Silver Corp. pursuant to purchase and sale agreement between Teuton Resources and Sabina dated June 16, 2014 and a royalty agreement between the Teuton and Sabina dated June 16, 2014. Teuton Resources has the right to repurchase one-half of the Sabina royalty for \$1,000,000.

One percent net smelter royalty in favour of Excellon Resources Inc. pursuant to a royalty agreement between Teuton and Excellon dated June 16, 2014. Teuton Resources has the right to repurchase one-half of the Excellon royalty for \$ 500,000.

Teuton Resources permit to conduct exploration work on the property expired in 2019. Several months ago, Teuton applied for another permit. The authors are not aware of any environmental liabilities or any other significant factors and risks that may affect access, title, or the right or ability to perform work on the property.

1.3 MINERALIZATION

The Del Norte property contains several mineralized zones. Brief description of the three most prominent zones are provided below.

LG vein, MINFILE No. 104A-161

The LG vein is a gold-silver bearing quartz-calcite vein which has been traced by drilling for 1.2 km and vertical extent of 240 metres. The true width ranges from 0.5 to 1.25 metres. The vein is spatially restricted to a narrow package of felsic volcaniclastic rocks located on the contact between intermediate volcanic/volcaniclastic rocks of the Hazelton Group and argillites/mudstones of the Salmon River Formation. Sulphides include mostly pyrite, galena and sphalerite with trace to minor chalcopyrite, tetrahedrite, silver-sulphosalts and native gold.

Kosciuszko (Del Norte) Zone, MINFILE No. 104A 176

The zone is not a part of LG vein, it is located some 60-70 metres NE from the presumed location of the LG vein. The zone is striking NNW-SSE and dipping 60 degrees to the northeast, what is roughly conformable with the bedding of the surrounding argillites/ siltstones. The zone, which is 50 metres long and from 3 to 10 metres wide is comprised of quartz-carbonate- sulphide cemented breccias, replacement zones and veins. Breccia fragments, composed mostly of argillite are very angular ranging in size from less than 1 cm to 10 cm across. Occasionally there are vuggy cavities. Gangue minerals include quartz, carbonates and minor amounts of zoisite/clinozoisite. Sulphides include pyrite, sphalerite, galena and tetrahedrite along with trace amount of realgar, arsenopyrite and electrum.

Hardpan, MINFILE No. 103P 318

The Hardpan Creek area contains eight mineral showings which were first evaluated in 1990. The mineralization is primarily fissure vein and replacement type. Geologic setting suggests the potential for vein, gold-copper porphyry and possible volcanogenic massive sulphide mineralization. Two types of mineralization are prominent, base metal and/or precious metal rich fissure type veins (+/- stringer stockwork) and gold-copper rich replacement type horizons. Mineralized zones are accompanied by strong alteration. Fissure type mineralization is by far the most commonly found and can be further subdivided into zinc-lead rich or copper-gold rich veins or stringer-stockwork zones. The orientation of the veins and stringers coincides with the main fracture and fault systems, trending towards the northwest, northeast and rarely east.

1.4 DRILLING

The above described mineralized zones witnessed several diamond drilling programs conducted from 1990 to 2005. Table 1 below show selected drill results from the LG Vein and Kosciuszko Zone.

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Drill Hole #	From (m)	To (m)	Interval (m)	Au (g/t)	Ag (g/t)	Target
DN02-01	11.90	43.50	31.60	3.66	213.6	Kosciuszko Zone
DN02-02	19.80	52.70	32.90	4.61	144.3	Kosciuszko Zone
DN02-03	1.30	24.70	23.40	7.46	277.3	Kosciuszko Zone
DN03-01	3.72	6.10	2.38	10.80	565.0	LG Vein
DN03-02	5.61	8.66	3.05	9.53	545.1	LG Vein
DN03-03	12.65	14.51	1.86	13.92	242.1	LG Vein
DN03-05	67.79	68.88	1.10	14.26	2619.7	LG Vein
DN03-06	121.49	122.5	1.01	15.46	1177.0	LG Vein
DN03-07	144.99	146.49	1.49	11.55	1346.0	LG Vein

DN04-01	96.62	97.35	0.73	9.26	958.3	LG Vein
DN04-02	148.96	152.55	3.60	8.37	297.3	LG Vein
DN04-04	103.91	107.05	3.14	12.27	640.8	LG Vein
DN04-07	63.98	65.04	1.07	8.98	315.8	LG Vein
DN04-08	79.89	81.29	1.40	6.89	640.1	LG Vein
DN04-16	66.14	66.63	0.49	16.8	1842.1	LG Vein
DN04-18	83.00	84.09	1.1	14.74	558.5	LG Vein
DN04-20	118.35	121.13	2.77	19.58	1434.8	LG Vein
DN04-26	106.60	106.8	0.20	1923.4	24.0	LG Vein
DN04-32	101.44	103.78	2.35	8.23	992.9	LG Vein
DN04-33	142.43	143.8	1.37	36.03	944.9	LG Vein
	2005 Drilling					
DN05-02	184.1	199.64	15.54	6.47	346.0	LG Vein
including	190.78	191.64	0.86	71.5	3269.9	LG Vein
DN05-05	180.44	184.71	4.27	4.48	476.5	LG Vein
DN05-06	34.44	37.03	2.59	5.42	734.7	LG Vein
DN05-07	168.3	170.09	1.79	6.42	129.1	LG Vein

1.5 INTERPRETATION AND CONCLUSIONS

Historic exploration programs completed on the Del Norte property led to discovery of several significant mineral occurrences. The most prominent mineralized zone discovered on the property to date is the LG vein which has been traced by drilling for 1.2 km and vertical extent of 240 metres. The vein is open in both directions as well as at depth. The true width of the vein ranges from 0.5 to 1.25 m, however in some places mineralization is not restricted to the vein but occurs also in the surrounding rocks forming breccia and stockwork zones, as well as satellite veins. A good example of such place is the area sampled by the authors in 2019, where several grab samples collected from semi-stockwork and satellite veins close to main LG vein returned up to 19.6 g/t gold and 3920 g/t silver. Another example is the Kosciuszko breccia zone located 60-70 metres from the presumed location of the LG vein.

Because of the snow and ice coverage the Kosciuszko zone as well as the stratigraphic contact which host the LG vein located directly to the north and south of the zone were not adequately drill tested. This portion of the contact zone which is approximately 700 metres long was tested only by six holes drilled from two set-ups. The LG vein likely continue further south across a large snow patch as indicated by the 2019 sample A19-260 collected 350 metres south from the most southerly located drilling pad which was used to test the LG vein. The sample, taken from suboutcrop of a vein returned 462 ppb gold and 25.8 g/t silver along with highly anomalous lead, zinc and arsenic values. The NMG vein and Bullion zone which are located within the same stratigraphic contact zone which hosts the LG vein may represent a more distant, southern continuation of the LG vein system. These two mineral occurrences were never drill tested.

The presence of chalcedonic quartz, vuggy cavities, realgar, silver sulphosalts and electrum within LG vein and Kosciuszko zone unequivocally point to a low-sulphidation epithermal environment.

Considering its epithermal origin, the LG vein shows remarkable horizontal and vertical continuity.

Silver sulfosalts and native gold of the LG vein often occur within soft, black, graphitic matrix which fills interstices between breccia fragments. In small-diameter (BQ size) drill holes, large part of this soft matrix with accompanying gold and silver mineralization, was very likely lost in the drilling process. Significantly, lower core recovery was frequently recorded in numerous historic drill intersections which cut across the strongly fractured vein wall rock and/or across the vein itself.

According to the heli-borne geophysical survey conducted on the property in 2005 by Aeroquest, the area north of the 2004 Esh drillpad which include the "LG extension" showing was tectonically shifted 500-600 metres east. Surprisingly, this important fact was never drill- tested or even acknowledged by previous property operators. Tectonic displacement would certainly explain why several 2004 drillholes which tested the "LG extension" zone failed to intersect the LG vein. It also indicates that the "LG extension" zone is not a part of the LG vein. The broad area of potential true LG northern extension becomes one of the primary targets for the next exploration program.

Another important area which warrants further investigation is the area of so-called Hardpan Creek zone which encompasses 8 mineral occurrences. Some of these occurrences display characteristics of porphyry gold-copper mineralization. However, geology of the entire area is complex and requires a detailed geological mapping as well as compilation of all available data to create a comprehensive geological model which will help to plan further exploration.

The presence of several large (up to 2.0 metres across) argillite boulders with syngenetic zinc mineralization found in 1994 on Del Norte Glacier rises the possibility of finding VMS mineralization similar to one which comprises the nearby BA zone which is a major Kuroko style VMS occurrence tested by 178 drillholes.

Large areas of the Del Norte property are weakly explored due to extensive ice and snow coverage. However, the rapidly receding ice enables better access to these areas which may host more mineralized zones.

1.6 RECOMMENDATIONS

Two areas related to the LG vein appear the most important targets for further investigations. The first area which likely host the northern extension of the LG vein appears to be detached and displaced eastward by postulated offset faults. The second area comprise the southern portion of the vein and its potential continuation to the south which include the NMG and Bullion occurrences. Both these areas will require a detailed geological mapping with emphasis on the structural and lithological observations which should shed additional light on the structural and lithostratigraphic context of the vein.

It must also be stressed that the results of the authors' reconnaissance rock sampling and geological observations on the SP vein area (the LG vein southern part) indicate that the LG vein is at least locally accompanied by a much wider zone of a stockwork-type mineralization. This widened zone when approached systematically may bring about a potential economic

improvement in evaluation of the vein. A problem of artificially diminished gold and silver grades due to inadequate drill practices can be addressed by drilling larger-diameter core and applying denser and more viscous drilling mud when intersecting the vein and its mineralized wall rocks.

The Hardpan Creek area is the third area of recommended exploration. In spite of the extensive exploration the area still requires a comprehensive model which would include all the existing information and provide basis for the further exploration. The company should also consider hiring a geophysical consultant who should address the relationship between the physical properties as measured on the core material and the results of the historic geophysical surveys.

It is also recommended to try locate the source of large (up to 2.0 metres across) argillite boulders with syngenetic zinc mineralization found in 1994 on Del Norte Glacier. The presence of boulders with this type of mineralization rises the possibility of finding VMS mineralization similar to one which comprises the nearby BA zone which is a major Kuroko style VMS occurrence.

The next exploration program should include drilling of 3,200 metres of core, geological mapping, rock and soil sampling and geophysical consultation. The total cost of the program is estimated at \$957,000.

2. INTRODUCTION

- (a) This report was prepared on the request of Decade Resources Ltd., a mineral exploration company, currently listed on the TSX Venture Exchange (TSX-V) under the symbol "DEC". The company is engaged in exploration for precious and base metals in British Columbia. The company was originally registered (incorporation number BC0750660) in British Columbia as Golden Oil Resources Ltd. on March 3, 2006. On December 01, 2006 Golden Oil Resources Ltd. changed its name to Decade Resources Ltd.
- (b) Decade Resources Ltd. has hired the authors to review the project, summarize the exploration results and prepare a written report to NI 43-101 standards. The authors are independent of Decade Resources and have no beneficial interest in the Del Norte project. Fees from this Technical Report are not dependent upon conclusions of this report. The authors are qualified persons as defined by Canadian Securities Administrator ("CSA") National Instrument ("NI") 43-101. This report has been prepared in accordance with the requirements of NI 43-101 of the CSA as set out in Form 43-101F1.
 - This report was prepared in order to enable Decade Resources Ltd. raising funds necessary to fulfill the earning requirements on the claims that form the Del Norte property.
- (c) The main source of information has been the geological field work conducted on the property by the authors in the period from 1993 to 1994 and from 2002 to

2005. The report also relies on assessment reports prepared by Teuton and Goodgold Resources as well as Teuton and Lateegra Resources press releases. It also relies on B.C. Dept. of Mines Annual Reports as well as bulletins published by various government geologists.

(d) Mr. Mastalerz conducted field work on the property during the summers of 2004 and 2005 directly supervising exploration programs. Mr. Walus conducted field work on the property in 1993 and 1994. In 2002 and 2003 Mr. Walus conducted field work and supervised the exploration programs. On September 29, 2019 the authors visited the property again fulfilling the due diligence work on behalf of Decade Resources.

2.1 Glossary of Technical Terms

Unless otherwise indicated, the following terms used in this report have the meanings ascribed to them below.

Atomic Absorption (AA) - Atomic absorption spectroscopy and atomic emission spectroscopy is an analytical procedure for the quantitative determination of chemical elements using the absorption of optical radiation by free atoms in the gaseous state. Atomic absorption spectroscopy is based on absorption of light by free metallic ions.

Adit - an entrance to an underground mine which is horizontal or nearly horizontal,

Anastomosing - Irregularly branching and reconnecting veins.

Breccia – Rock made up of angular or sub-angular fragments >2mm embedded in a fine-grained matrix.

Cataclasite - a type of fault rock that has been wholly or partly formed by the progressive fracturing and comminution of existing rocks.

Dacite – Volcanic rock rich in quartz and plagioclase

Ductile deformation – This type of deformation occurs when a substance is stressed to a point where it begins to behave in a plastic manner.

Dip – An angle of inclination between a geological feature/rock and horizontal plane.

Fault – A fracture in a mass of rocks accompanied with relative movement between its two blocks. Faults are the result of the rock's mechanical response when submitted to sufficient stress as to induce permanent deformation.

Fault gouge – Unconsolidated, often soft rock formed along fault plane.

Facies - a body of rock with specified characteristics, which can be any observable attribute of rocks (such as their overall appearance, composition, or condition of formation).

Felsic – igneous rock rich in elements which for feldspar and guartz

Foliation in **geology** refers to repetitive layering in rocks. Each layer can be as thin as a sheet of paper, or over a meter in thickness.

Footwall - Part of a fault which occurs below the fault plain

Hanging wall - Part of a fault which occurs above the fault plain

Hyaloclastite - a volcanoclastic accumulation or breccia consisting of glass fragments formed by quench fragmentation of lava.

Igneous – A primary crystalline rock formed by the solidification of magma.

Intrusion – A body of igneous rock formed by the consolidation of magma intruded into other rocks, in contrast to lavas, which are extruded upon the surface.

Induced polarization (**IP**) - geophysical imaging technique used to identify the electrical chargeability of subsurface materials, such as ore.

Intermediate volcanics - refers to the chemical composition of a volcanic rock that has 52-63 wt % SiO2 being an intermediate between felsic and mafic compositions. Typical intermediate rocks include andesite, dacite and trachyandesite.

Matrix - the fine-grained materials that surround larger grains in a rock.

Mylonitization - Deformation of a rock by extreme microbrecciation, due to mechanical forces applied in a definite direction, without noteworthy chemical reconstitution of granulated minerals.

Outcrop – The part of a rock formation that is exposed at the Earth's surface

Pluton – A general term applied to a body of intrusive igneous rock, irrespective of its shape, size or composition.

Propylitic alteration – Hydrothermal alteration which convert existing minerals to chlorite, epidote, carbonates, quartz and pyrite.

Sedimentary – Pertaining to rocks formed by the accumulation of sediments, formed by the erosion of other rocks.

Sericitization – alteration process in which minerals are converted to fine grained mica called sericite.

Siliciclastic sediments - silica-based sediments

Silicification – alteration process in which minerals are replaced by fine grained silica.

Shear zone - Deep level equivalents of faults. It forms as a response to inhomogeneous deformation partitioning strain into planar or curviplanar high-strain zones.

Slickensides - a polished and striated rock surface that results from friction along a fault or bedding plane.

Strike – A direction of line formed by the intersection of strata surfaces with the horizontal plane, always perpendicular to the dip direction.

Thrust fault - type of reverse fault that has a dip of 45 degrees or less.

VLF – geophysical electromagnetic method that relies on transmitted radio waves.

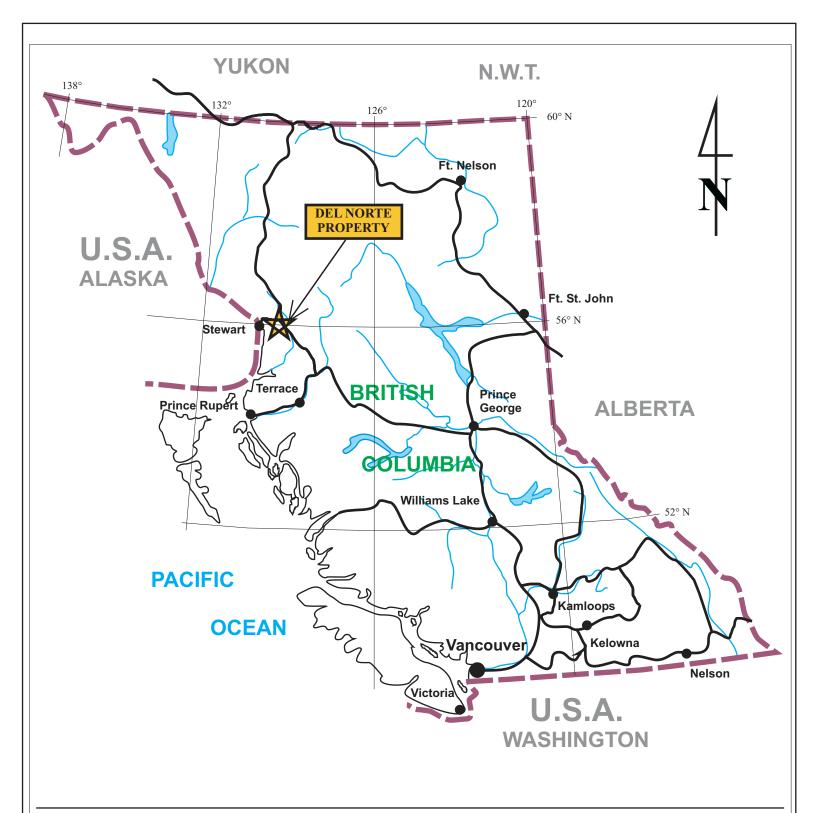
Volcanoclastic - undisturbed deposits of volcanic materials.

ZTEM (Z-Tipper Axis Electromagnetic) - an airborne electromagnetic **survey** system which detects anomalies in the earth's natural magnetic field. These disruptions are caused by zones of rock that conduct or resist electrical current more than the surrounding rock, like ore deposits.

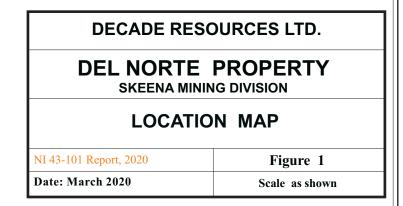
2.2 Abbreviations Used in the Report

AR Assessment Report g/t Gram per tonne

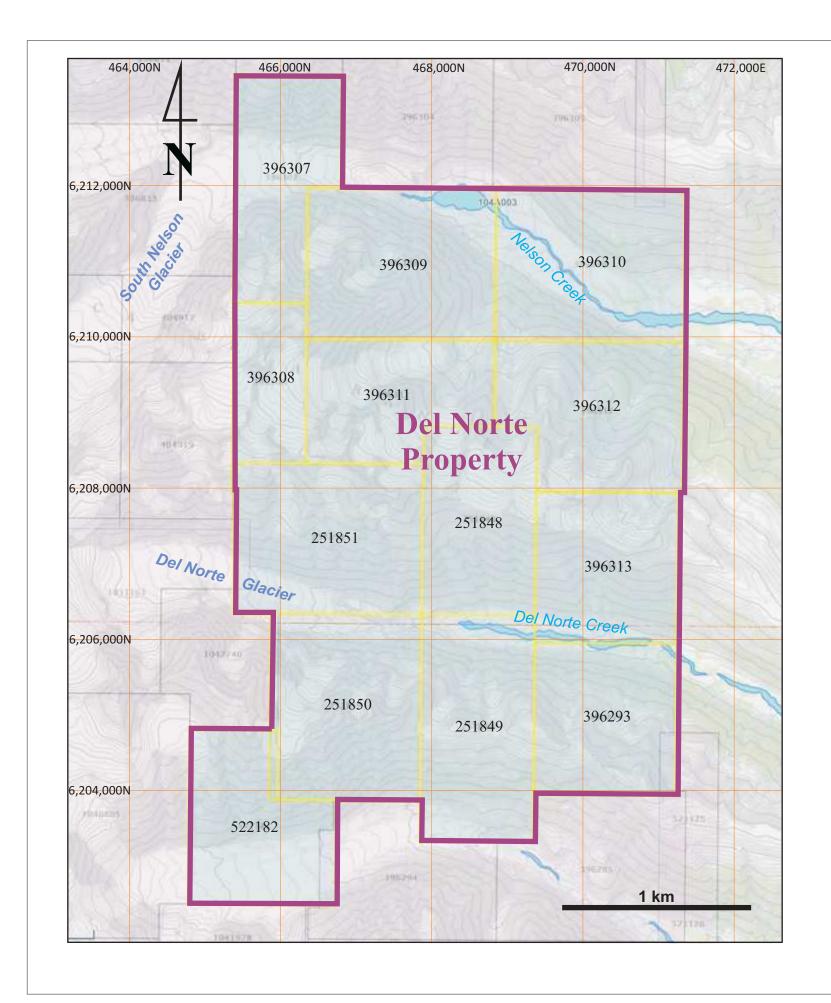
m meterkm kilometerppm parts per millionppb parts per billion

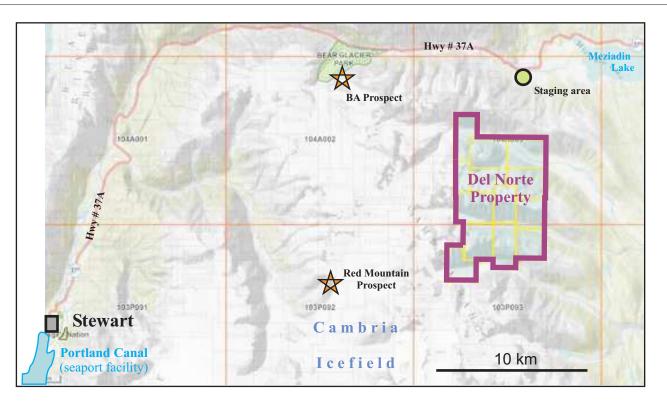


To accompany report by K. Mastalerz & A. Walus

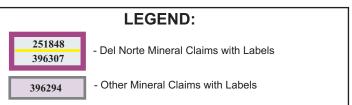


100 km

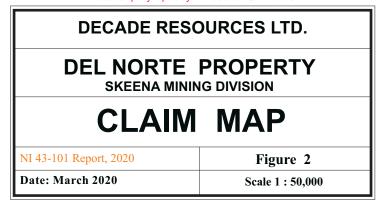




Topography from MTO Map: https://www.mtonline.gov.bc.ca/mtov/map/mto/cwm.jsp?site=mem_mto_min-view-title



To accompany report by K. Mastalerz & A. Walus



3. RELIANCE ON OTHER EXPERTS

The authors of this report have the firsthand knowledge about the exploration programs conducted in 1993 and 1994 and from 2002 to 2005 since they either participated or directly supervised these programs. For information about the exploration conducted in other periods the authors relied on reports prepared by geologists who worked in this area, as well as various government publications. Full list of these reports is provided in References.

4. PROPERTY DESCRIPTION AND LOCATION

The Del Norte property is located 34 km east of Stewart in the Skeena Mining Division (see Figs. 1 and 2). It is centered on Latitude 56° 00' North and Longitude 129° 31' West on NTS map sheets 104A/4E and 104A/3W.

The property consists of 13 mineral tenures (claims) covering 5830 hectares. They are 100 percent owned by Teuton Resources. Expiration dates of the claims and other relevant information are shown in table 1 below. Claims location is shown on Fig. 2. According to Minerals Titles Online there are no Crown Granted claims in the area covered by the Del Norte claims.

Table 2. Claims information

Title #	Claim Name	Owner	Issue Date	Good to Date	Area (ha)
251848	CROESUS 1	126630 (100%)	1987/MAY/04	2020/SEP/30	375.00
251849	CROESUS 2	126630 (100%)	1987/MAY/04	2020/SEP/30	450.00
251850	CROESUS 3	126630 (100%)	1987/MAY/04	2020/SEP/30	500.00
251851	CROESUS 4	126630 (100%)	1987/MAY/04	2020/SEP/30	500.00
396293	MIDAS 1	126630 (100%)	2002/SEP/10	2020/SEP/30	400.00
396307	LORD NELSON 6	126630 (100%)	2002/SEP/10	2020/SEP/30	450.00
396308	LORD NELSON 7	126630 (100%)	2002/SEP/10	2020/SEP/30	375.00
396309	HORATIO 1	126630 (100%)	2002/SEP/09	2020/SEP/30	500.00
396310	HORATIO 2	126630 (100%)	2002/SEP/09	2020/SEP/30	500.00
396311	HORATIO 3	126630 (100%)	2002/SEP/09	2020/SEP/30	500.00
396312	HORATIO 4	126630 (100%)	2002/SEP/09	2020/SEP/30	500.00
396313	HORATIO 5	126630 (100%)	2002/SEP/10	2020/SEP/30	400.00
522182	CORDIERITE	126630 (100%)	2005/NOV/10	2020/SEP/30	380.18

Decade Resources has the right to earn up to a 55 % interest under the following terms:

- Payment of \$400,000 over 4 years with an initial payment of \$20,000.
- Issuing 800,000 shares of Decade on signing.
- Issuing \$180,000 of Decade stock over a 4-year period.
- Expenditures of \$4,000,000 over 5 years.

Decade has the right to earn an additional 20% by placing the property into production.

The property is subject to the following royalties (the "Pre-Existing Royalties"): one percent net smelter royalty in favour of Sabina Gold & Silver Corp. pursuant to purchase and sale agreement between Teuton Resources and Sabina dated June 16, 2014 and a royalty agreement between the Teuton and Sabina dated June 16, 2014. Teuton Resources has the right to repurchase one-half of the Sabina royalty for \$1,000,000.

One percent net smelter royalty in favour of Excellon Resources Inc. pursuant to a royalty agreement between Teuton and Excellon dated June 16, 2014. Teuton Resources has the right to repurchase one-half of the Excellon royalty for \$ 500,000.

Teuton Resources permit to conduct exploration work on the property expired in 2019. Several months ago, Teuton applied for another permit. The authors are not aware of any environmental liabilities or any other significant factors and risks that may affect access, title, or the right or ability to perform work on the property.

5. ACCESSIBILITY, CLIMATE, LOCAL RESOURCES, INFRASTRUCTURE AND PHYSIGRAPHY

5.1 ACCESSIBILITY

Access to the property is by helicopter based in Stewart. The shortest trip (7-8 km) to the main showings (LG vein/Kosciuszko zone) is from a staging area located on Highway 37A near the bridge on the Surprise Creek (Fig. 2).

5.2 CLIMATE

The weather is typical of the North Coast of British Columbia with wet summers and heavy snowfall in the winters. Large snow-drifts cover parts of the property until mid-June, with minor areas of permanent snow found at the highest elevations and in sheltered areas. Because of the mountainous terrene and large snowfall, the surface exploration in the Stewart area is restricted to summer and early fall with the maximum rock exposure occurring in late August to October. However, once development starts, year-round core drilling and underground work can proceed and was done on many properties in the general area.

5.3 LOCAL RESOURCES & INFRASTRUCTURE

The closest settlement is the municipality of Stewart (population 500) approximately 30 km WWS of the property (Fig. 2). The town has two grocery stores, gas station, hotels and two helicopter bases and an air strip. The Stewart area also features a mill located by the Premier Mine as well as the year-round seaport. Excellent paved roads connect Stewart with Smithers and Terrace, which are major supply centers in this part of British Columbia.

5.4 PHYSIOGRAPHY

Several claims which comprise the property are located on the eastern edge of the extensive Cambria Icefield, along the headwaters of Willoughby, Del Norte and Nelson Creeks. Elevations vary from approximately 600 metres on the creek bed at the eastern edge of the property to more than 2000 metres near ridge tops. Slopes range from moderate to precipitous. Vegetation in the area changes from mountain hemlock and balsam at low-lying elevations to shrubs, mountain grasses and heather at higher elevations.

6. HISTORY

6.1 PRIOR OWNERSHIP AND OWNERSHIP CHANGES

1913-1922

The property was originally staked as the "Bullion" claim, sometime prior to 1913. Between this first staking and 1922, the property was re-staked as the Delnorte Group by Green and Ficklin of Hyder, Alaska (MANDY, 1939).

Late 1920s

Several placer mining leases were recorded on Nelson Creek in the late 1920s (Property File - Placer Mining Lease, 1928).

<u>1939</u>

That year, Owen McFadden of Stewart, backed by a syndicate, explored the ground by a series of fifteen open-cuts and some small "popholes". At this time the property was known as the "'Meziadin Group".

1987

Teuton Resources Corp. acquired the Croesus claims.

1989

Teuton Resources Corp. staked the Bond 1-7 claims over the area. Goodgold Resources Ltd. entered into an option agreement with Teuton.

<u>1991</u>

Goodgold terminated option agreement with Teuton.

2003

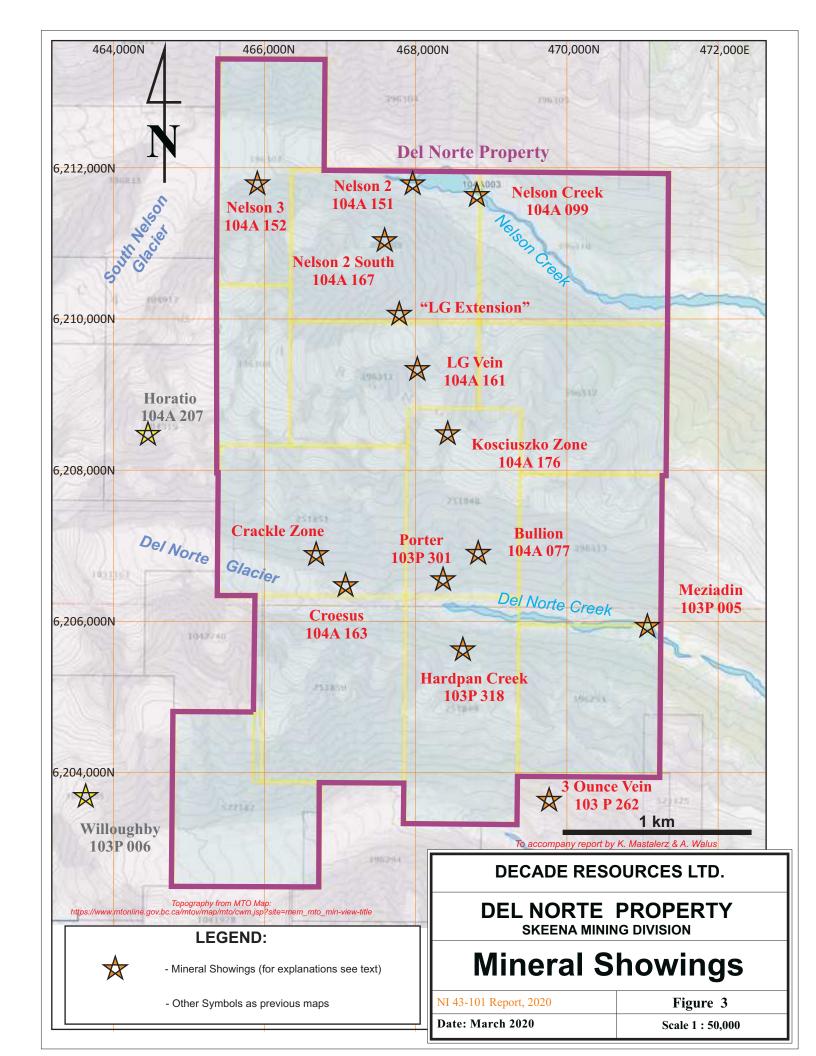
Teuton optioned the property to Lateegra Resources Corp.

2005

Lateegra Resources Corp. terminated option agreement with Teuton. The same year Teuton optioned property to Sabina Gold and Silver Corporation.

<u>2007</u>

Sabina terminated option agreement with Teuton.



6.2 SUMMARY OF PREVIOUS EXPLORATION PROGRAMS

1913-1922

A small adit was driven on the north side of Del Norte creek to test a zone of quartz veining paralleling the contact between Bowser sediments and Hazelton volcaniclastics (Mandy, 1939).

Late 1920s

Moderately fine, flaky, but still rough gold has been panned from the Nelson Creek near the foot of the glacier and at other locations along the creek (Geological Survey of Canada Memoir 32).

1930s

Cominco run a gold sluicing operation on Nelson Creek. In 1939, Owen McFadden of Stewart, backed by a syndicate, explored the ground by a series of fifteen open-cuts and some small "popholes". At this time the property was known as the "'Meziadin Group". In the same year, the property was visited by Dr. Mandy of the B.C. Department of Mines; Mandy examined and sampled several of the showings. Samples results indicated erratic low-grade gold mineralization associated with copper and occasional zinc values. According to historic records, most of this sampling was done on the north side of Del Norte Creek (MANDY, 1939).

1960s

The area presently covered by Del Norte property was explored again by companies searching for porphyry copper deposits (MINFILE).

1987

Teuton Resources Corp. staked the Croesus claims and carrie d out a program of rock and silt sampling. Grab sample collected in Hardpan Creek area (Fig. 3) returned 19.3 g/t gold (Cremonese, 1988).

1988

Teuton followed up on the discovery of a gold-bearing sulphide lens with a limited program of geological mapping, prospecting, and rock and soil sampling in the Bullion and Hardpan Creek areas (Fig. 3). Two zones, one featuring lead-zinc mineralization, the other copper-gold, were discovered in the Hardpan Creek drainage (Cremonese, 1989).

1989

Goodgold Resources Ltd. flew a heli-borne VLF-EM and magnetometer survey over the southern part of the Bond claims. Structural analysis of the survey data suggests the presence of numerous NW-SE oriented structural features (faults?). These inferred structures are considered to be of prime exploration importance because of possible presence of structurally controlled mineralized bodies (Mallo & Dvorak, 1989).

1990

That year Goodgold mounted a major \$500,000 plus program focusing mostly on the Hardpan Creek portion of the property. The program consisted of detailed work on the Croesus 2 claim

(Hardpan Creek area) and reconnaissance exploration on the remainder of the property. Exploration in the Hardpan Creek area (Fig. 3) consisted of establishing a grid (12km of picketed lines), mapping and prospecting (66 grab samples), blasting and trenching (600m, 523 chip samples), soil geochemical survey (347 samples), geophysical surveys (VLF-EM, magnetic, UTEM) and diamond drilling (12 holes, 1119 m of BQTW core, 536 core samples). Reconnaissance work included mapping and prospecting (71 grab samples), stream sediment sampling (98 samples), trenching (300m, 64 chip samples) and contour soil sampling (64 samples). Highlights of the program include the discovery of the gold-copper "O" zone, the gold-silver "Humdinger" zone, the lead-zinc "Grizzly" zone and gold-silver LG vein (Bishop and Gal,1991). Chip sample DNLGR-298 collected from LG vein assayed 30.3 grams per tonne gold, 1,42 grams per tonne silver, 2.12 per cent lead, 0.45 per cent copper and 0.38 per cent zinc across 1 metre (Bishop and Gal, 1991). No follow up work was done on the LG vein as Goodgold dropped the property the same year.

1991

Goodgold carried out another \$100,000 of work before relinquishing its option. During this phase, which concentrated on the north side of Del Norte Creek, geochemical sampling, prospecting and mapping was carried out. A quartz-sulphide vein, the NMG vein, was discovered about 150 metres north of the Bullion showing. The vein outcrops intermittently over a distance of 225 metres and lies within the same sedimentary-volcanic horizon which hosts the LG vein discovered in 1990 (Fig. 3). The NMG vein yielded significant gold and silver values. The highest assay came from a one metre chip sample across the vein at its southern exposure. The sample assayed 10.6 grams per tonne gold and 571.45 grams per tonne silver (sample DM-MR-82646, Assessment Report 22103). A zone of quartz calcite stringers, some highly auriferous, was also discovered north of the toe of Del Norte Glacier. Soil sampling over this area, named the "Crackle" zone (Fig. 3), disclosed widespread anomalous copper values.

1993

That year Teuton carried out a modest program of reconnaissance rock sampling at four sites within the Del Norte property. Sampling in the Crackle zone area resulted in the discovery of several new areas of Au-Ag-As bearing quartz-sulfide veins, some with high gold values. The best gold assay of 66.38 g/t (sample DN93-123) came from a 5 cm wide vein carrying abundant arsenopyrite. These veins are now known to occur over an area roughly 700 metre square encompassing both sides of Del Norte Glacier (Cremonese, 1994).

1994

Teuton Resources continued its reconnaissance rock sampling program collecting a total of 139 samples. Several highly anomalous gold-silver grab samples associated with tetrahedrite and galena mineralization were collected along the south flowing creek located approximately 300 m west from the center of Crackle zone (Fig. 3). The samples returned up to 16.83 g/t gold and 4279 g/t silver. Several boulders up to 2.0 metres across of hydrozincite stained black argillite/ash tuff with "sedex" type zinc mineralization were found in medial moraine of the Del Norte glacier (Cremonese, 1995). Two samples taken from two separate boulders containing 7-10% sphalerite and 2-3% pyrite returned 6.79% and 7.96% zinc.

2002

Teuton crew investigating the area north of the Bullion zone discovered Kosciuszko zone (Minfile No. 104A 176; Fig. 3). Chip sample across the northern end of the zone returned a weighted average of 6.14 grams per tonne gold and 630.86 grams per tonne silver over 10.0 metre interval (Teuton Resources Corp. News Release, September 26, 2002). Three holes drilled from a single pad which targeted the zone intersected true widths of mineralization ranging from 8.5 to 10 metres and carrying gold values ranging from 3.22 to 6.93 g/t and silver values ranging from 162.34 to 251.60 g/t. The best drilling result came from hole DN02-3 which returned a weighted average of 7.64 g/t gold and 277 g/t silver over 23.4 m.

One metre chip sample across LG vein yielded 13.89 g/t gold, 545.89 g/t silver, 1.82% zinc and 0.63% lead.

2003

Lateegra Resources Corp. optioned the property from Teuton and drilled nine holes testing the LG vein (Fig. 3) with seven of these holes returning significant gold-silver values. The success of this program led to a large 2004 program.

2004

Lateegra Resources continued to drill the LG gold-silver vein. Thirty-six holes totaling 4,519 metres were completed in a series of two to four-hole fans targeting the LG vein. The intersection in hole 2004-01 is representative of the vein tenor: 9.25 grams per tonne gold and 958 grams per tonne silver over 0.7 meters. Drilling in 2004 in-filled the area tested in 2003, delineating the LG vein over a horizontal distance of 750 meters. However, a series of large step-out holes along strike to the north did not locate the structure.

2005

Sabina Gold and Silver Corporation acquired the Del Norte property from Teuton Resources Corporation. Ten holes totaling 1,400 metres were completed along the southern 500 metres of the 1.2-kilometre-long LG vein. Later that year a helicopter-borne AeroTEM II electromagnetic and magnetic survey was carried out on behalf of Teuton Resources Corp. over the Del Norte-Midas property with a total of 1,299.5 line-kilometres flown.

2006

Sabina embarked on extensive prospecting program covering numerous areas of the property as a follow-up of the 2005 geophysical airborne survey.

2014

Teuton Resources drilled 5 short holes totaling 592 metres from one pad located in the area of the Hardpan zone. Hole DN14-02 returned 0.26 g/t gold and 0.09% copper over 140.5 m. Hole DN14-05 yielded 0.44 g/t gold and 0.07% copper over 107.9 m (Cremonese, 2014).

<u>2016</u>

Teuton Resources drilled 13 holes totaling 1,822.4 m from 4 pads of which 12 reached the planned depth. Teuton tested 4 separate areas with one pad assigned to each area. Three pads

tested targets in the Hardpan and Bullion areas, one pad tested LG vein. The most significant results were obtained from drilling on pad # 2 (located 150 meres from the 2014 pad) where holes DN16-03, 04 and 05 encountered a significant porphyry copper-gold mineralization associated with strong K-feldspar alteration. The best results came from hole DN16-05 where a 74.2 m interval averaged 0.27 g/t gold and 0.15% copper (Cremonese, 2017).

2017

Teuton conducted rock and soil sampling program over four different areas of the property collecting 82 rock and 99 soil samples. Crackle zone soil geochemical anomaly was extended by 300 metres. Talus fines sampled in the vicinity of Kosciuszko zone detected a copper-zinc anomaly.

2018

Geotech Ltd. of Aurora, Ontario acting on behalf of Teuton Resources carried out a ZTEM and magnetic airborne geophysical survey. The survey has selected 11 targets for precious metals mineralization and 3 targets for porphyry/stockwork mineralization (Geotech, 2018).

2019

Teuton Resources drilled four holes on the Del Norte property. The first two holes targeted a sharp EM anomaly (identified by a 2005 airborne survey) which lies parallel to and just east of the LG vein. The remaining 2 holes were drilled into recently discovered mineralized zone located in the western part of the property. No significant mineralization was encountered in the holes.

Extent and the results of the earliest exploration programs conducted on the Del Norte area are not precisely known. However, the following sub-chapters summarize the most important exploration efforts and achievements.

7. GEOLOGICAL SETTING AND MINERALIZATION

7.1 REGIONAL GEOLOGY

The entire Canadian Cordillera is interpreted to comprise a series of oceanic and island-arc crustal fragments (terranes) which were successively accreted to the proto-North American continent. These fragments occur within three large-scale geomorphological units which are labelled Insular, Coastal and Intramontane belts. These accreted terranes adjoin the deformed units of the continental crust to the east – the Omineca crystalline belt and the Foreland fold-and-thrust belt (Fig. 4).

The Del North property lies entirely within the Stikine Terrane (Stikinia) of the northern Intermontane Belt of the western Canadian Cordillera. It is located along the easternmost rim of the Coast Ranges and the Bowser Basin to the west (Fig. 5). Stikinia comprises Late Paleozoic and Mesozoic volcanogenic and sedimentary successions with their coeval intrusive complexes. It is overlain by a thick succession of the Jurassic Bowser Lake and, further east, the

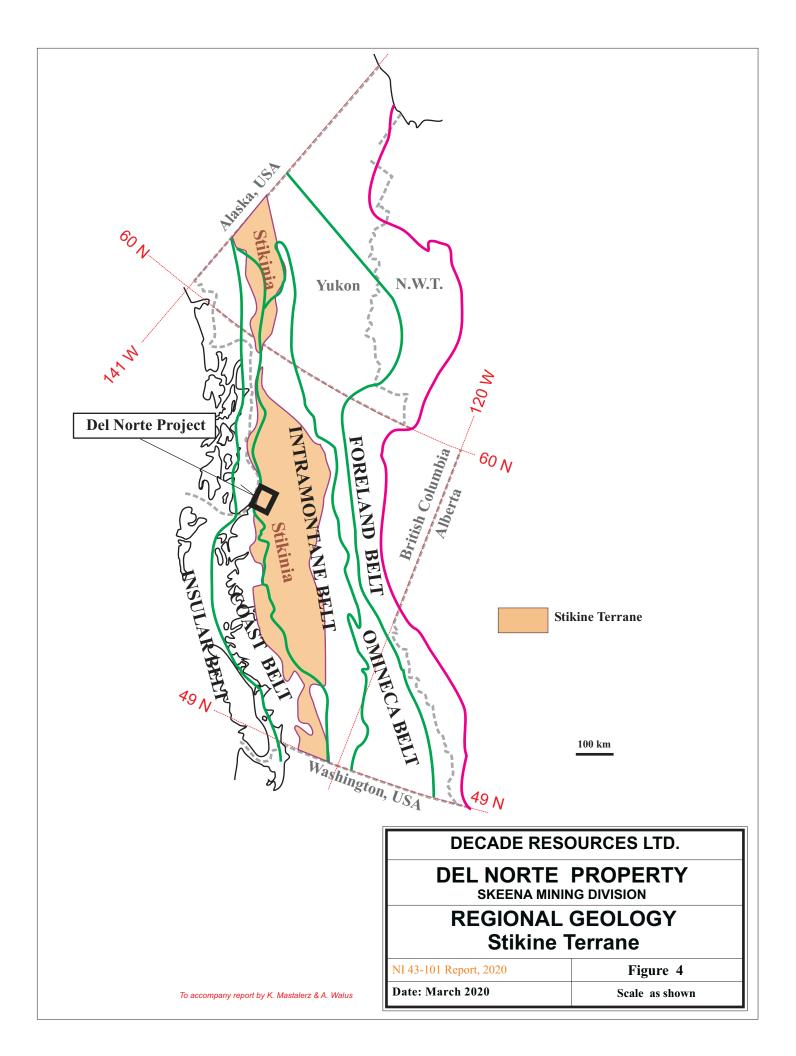
Cretaceous Sustut Groups which comprise predominantly marine siliciclastic deposits. Much younger continental succession of sedimentary rocks and minor accompanying volcanics (Sloko Group) occur rather fragmentarily over the area.

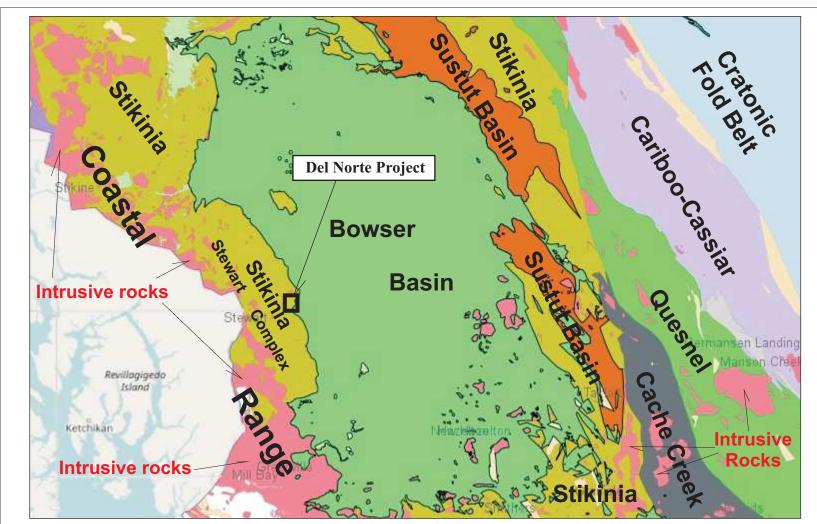
The oldest strata of the Stikinia – Paleozoic Stikine assemblage – are exposed only fragmentarily, mostly along the western flank of the Stikine Terrane (Monger, 1977). The assemblage includes three distinct limestone complexes which are interbedded with volcanic rocks and minor siliciclastics. These rocks have been invaded by moderately diversified plutonic intrusions. It is interpreted that this predominantly marine assemblage represents remnants of an oceanic arc system of the eastern paleo-Pacific.

The oldest package of the Stikine assemblage includes the Devonian limestone, bimodal volcanic and some associated plutonic rocks (Gunning, 1993). The middle package consists of the Carboniferous limestone and mafic volcanics, and the youngest package includes the thick but discontinuous Lower to Upper Permian limestone with minor argillites. The latter limestone is frequently fossiliferous and, locally, very well stratigraphically dated (age-specific conodonts and foraminifers which occur locally within this unit). The Permian limestone of the Stikine assemblage has been lately lithostratigraphically incorporated into the Ambition Formation (e.g. Gunning et al., 1994; Nelson et al., 2008; Barresi et al., 2015). The Ambition Fm. limestone is locally overlain, apparently conformably, by the pelagic cherty to argillaceous sediments (cf. Brown et al. 1996). The age of this latter unit is not precisely defined so far.

The Permo-Triassic contact is relatively poorly exposed, however it appears that it is commonly associated with a significant stratigraphic hiatus, change in character (or deformation style) of the rock formations and/or locally it was described as an angular unconformity. It is commonly accepted that the Upper Triassic package of the Stuhini Group overlays the older succession unconformably (Brown et al. 1996).

The Stuhini Group is the second important assemblage of the Stikine terrane and was originally defined (along with its lower boundary) in the Tulsequah map area as a succession of volcanogenic and sedimentary rocks that lie above the Middle Triassic unconformity (Kerr, 1948, Souther, 1971). Their equivalent strata are known as the Takla Group further south and east. The Stuhini Group is facially diversified, characterized by relatively rapid facies changes, and commonly dominated by volcanogenic rocks (e.g. Brown et al., 1996). The Triassic Stuhini Group also incorporates, locally very thick packages of sedimentary rocks which frequently include characteristic, poorly sorted polymictic conglomerates (with abundant fragments of limestone and the older volcanic rocks) and sedimentary breccias, as well as immature finer grained sediments such as tuffaceous wackes and siltstones or arkosic greywacke (Brown et al., 1996). Locally (e.g. Terrace area) the dark green polymictic conglomerates and breccias are found overlying directly the surface of a distinct erosional unconformity (Barresi, 2008, Mastalerz 2018). Some of the finer-grained sedimentary rocks and limestone layers contain Norian (middle Late Triassic) age-diagnostic ammonites and some other, stratigraphically less important fossils.





To accompany report by K. Mastalerz & A. Walus

DECADE RESOURCES LTD.

DEL NORTE PROPERTY SKEENA MINING DIVISION

REGIONAL GEOLOGY Terranes and Project Location

NI 43-101 Report, 2020	Figure 5
Date: March 2020	Scale as shown

100 km

Volcanic facies of the Stuhini Group include various end-members of predominantly mafic and intermediate extrusive and effusive rock varieties (Brown et al., 1996; Barresi et al., 2015). However, felsic volcanic units are also distinguished locally within these facies.

The Lower to Middle Jurassic Hazelton Group is the youngest significant and the most economically important, assemblage of the Stikinia. It comprises diversified volcanogenic and minor sedimentary rocks. Similarly to the Stuhini Group the succession of the Hazelton Group is characterized by abrupt lateral facies changes, which causes difficulties in a consistent lithostratigraphic subdivision of this package over the entire region. As a result, variable formation (and minor lithostratigraphic units) names have been applied for potentially equivalent units in specific distinct areas. The constituent lithologies include felsic, intermediate and minor basic volcanic rocks (both effusive and extrusive varieties). Thick packages of poorly sorted volcaniclastic rocks, frequently with significant admixtures of epiclastic component fragments, are very common over the entire Stikinia area. Massive and locally flow banded, commonly fragmental porphyritic dacite, as well as andesite flows, crystal-rich tuff, thick sequences of lapilli tuff, locally ignimbrites are probably the most important volumetrically end-members of this Group. Quite common are sequences of thick bedded hyaloclastic tuffs. Rhyolite flow units, locally flow-banded) and basalt flows are considerably less common and less widespread.

Sedimentary components of the Hazelton Group include thinner sequences of lime wacke (locally containing Toarcian ammonites), black to dark-grey tuffaceous mudstone and argillites, and argillaceous tuffs.

Rock formations of the Hazelton Group are separated from the underlying Stuhini Group by a significant stratigraphic hiatus and its lower contact is interpreted as a prominent both, erosional and angular unconformity (cf. Brown et al. 1996).

In the Stuart and nearby areas a typical lithostratigraphic subdivision of the Hazelton Group includes three widespread, distinct formations. The oldest succession, Unuk River Formation, comprises predominantly andesitic volcaniclastic rocks (typically ash tuffs) and minor sedimens. Typical lithologic components of the following, Betty Creek Formation are thick sequences of massive, locally gradded bedded, coarse-grained volcanoclastics, with common admixture of sub-rounded epiclastic fragments, and flow units of intermediate composition. The rock formation is also characterized by a common maroon, brown, red and less frequently, green colors of the rock units. The youngest, Salmon River Formation includes predominantly darkgrey to black tuffaceous mudstone and argillaceous tuffs. Locally, a relatively thinner, predominantly felsic unit of the Mt. Dilworth Formation occurs along the contact between the Betty Creek and Salmon River formations.

The rock formations of the Hazelton Group are overlain (by overlap, most likely also by local onlap?) by a very thick succession of the Middle-to-Late Jurassic Bowser Lake Group which constitutes a sedimentary infill of the Bowser Lake Basin (Fig. 5). These sediments comprise predominantly dark-grey to black coloured turbiditic sandstone, siltstone and minor conglomerates. Considerably younger (Cretaceous) sedimentary rocks are filling-up the successor basin compartments along the eastern margin of the Bowser Basin – Sustut Basin

(Fig. 5) and are related to a considerable eastward shift of local depo-centers which resulted from the late-to-post Jurassic tectonic folding and uplift of the Stikine terrane area.

7.2 LOCAL GEOLOGY

The large lobe of the Stikinia terrane exposed along the western boundary of the Bowser Basin near Stuart has been named the Stewart Complex (Grove, 1986) – Fig. 5). It stretches from the Iskut River valley in north-west to Kitsault (Alice Arm) area at south-east for almost 150 kilometres and is 30- 40 km wide. This relatively well-defined segment of Stikinia is known from numerous mineral occurrences and several ore deposits, many of which are incorporated into the Iskut, Sulphurets, Stewart and Kitsault mining camps. Several mapping programs at various scales have been completed in this area since 1911 (for a discussion see: Alldrick, 1989, Greig et al., 1994b, Lewis et al., 2001). However, in spite of these mapping projects and numerous exploration programs several aspects of stratigraphy, lithostratigraphy and structure of this area are not fully explained.

There are a few factors which strongly contributed to difficulties in deciphering these features. The region is partly covered with ice fields and the glaciated areas are not available for standard surface investigations. The rock formations of the complex are thick successions of predominantly volcanogenic rocks which typically display rapid facies and thickness changes. Additional complications resulted from numerous penetrations of the complex by syn- and post-accumulation intrusive rocks and associated metamorphism, which both strongly overprinted the primary features. Finally, the Stikinia succession of the Stuart complex is very strongly tectonically deformed including diversified folding and associated locally strong faulting.

In spite of a significant progress in exploration and mapping since the mid-1980', the nomenclature for the Mesozoic strata of the Stuhini and Hazelton groups in northwestern British Columbia is still not completely consistent and remains slowly evolving. However, there already exists a commonly accepted and applied lithostratigraphic subdivision of the Hazelton Group strata over a significant part of the Stewart complex (Grove, 1986; Alldrick, 1989; Lewis et al. 2001).

The volcanogenic and sedimentary rocks of the Hazelton Group overlay the Stuhini rocks with an erosional unconformity (cf. Lewis et al. 2001). The Stuhini rock formations exposed only locally in the area of the Stewart Complex are moderately diversified and comprise pyroxene porphyry flows and tuffs, turbidites, conglomerates and limestones (Alldrick, 1989). The lowermost accumulations of the Hazelton Group are usually lumped together into a Unuk River Formation (Grove, 1986). This formation was preliminarly labelled as the "Andesite Sequence" (Alldrick, 1985) and, if complete, it includes three informal andesite members (composed predominantly of massive ash tuffs and volcaniclastic sediments) which are separated by two members of predominant turbidites, and topped with so called "Premier Porphyry" (feldsparhornblende-phyric flows and tuffs). However, lithologies of the Unuk River Formation may differ quite significantly from this classic sub-division in some local areas of the Stewart complex.

However, the lowermost part of the Hazelton Group comprises locally the granitoid-clast conglomerate and fossiliferous limy sandstone/siltstone and is commonly described as the Jack Formation in the northwestern part of the complex (Iskut River-Eskay Creek-John's Peak area; see Lewis et al., 2001). This unit has been proposed as a distinct lithostratigraphic formation in the Iskut-Unuk River area (Henderson et al., 1992).

The following thick succession of predominant volcaniclastic rocks (previously called the "Epiclastic Sequence"; Aldrick 1985) is usually collectively named as the <u>Betty Creek Formation</u> (Grove, 1986). This lithostratigraphic unit is characterized by thick, relatively monotonous sequences of characteristic hematite-bearing epiclastic and other volcaniclastic rocks, from very coarse- and medium-grained, commonly thick-bedded and homogenous, but also showing variable graded-bedding typical of debris-flow to high-density turbidite deposition. Their common hematite content results in a distinct brownish-red rock and/or maroon rock colour. However, some layers and sequences display dirty-green to grey, and/or yellowish colouring. Andesite and/or dacite tuff and flow units are less common and volumetrically less significant components of the Betty Creek Formation. The latter end-members occur more frequently in the lower part of the Formation (Alldrick, 1989).

The Mt. Dilworth Formation is the next, younger formal litostratigraphic unit of the Hazelton Group distinguished within the Stewart complex (Alldrick. 1987). Because of its predominant felsic volcanic components, this unit was preliminarly called the "Felsic Volcanic Sequence" (Alldric, 1985). This Formation comprises quite highly diversified volcanic rocks, predominantly of felsic composition. They include fine tuff, ash flow tuff, pumice tuff, lapilli tuff (sometimes welded or ignimbritic), tuff breccia and a limited amount of flow banded dacite and rhyolite flows. Locally the topmost part of the formation includes the dark-grey pyritic felsic tuff and/or black argillaceous lapilli tuff (Alldrick, 1985). Thickness of this formation is highly variable, locally it pinches out and is absent in some areas.

The Salmon River Formation (Shofield and Hanson, 1922; formerly also labelled as the "Sedimentary Sequence" by Alldrick 1988) is considered the final, youngest end-member of the Hazelton Group in the Stewart complex. It consists predominantly of black to dark-grey tuffaceous mudstone, argillite and minor poorly sorted, matrix-supported argillaceous coarse tuff. Locally, an impure, wacke limestone, sometimes fossiliferous (Alldrick, 1985, 1989) occurs at the base of this formation. It is worth noting that in some areas of the Stewart complex the lithostratigraphic components of the Salmon River (as well as Mt. Dilworth) Formations vary significantly. For example, in the Eskay Creek-Sulphurets area a characteristic mafic flow unit has been identified in the uppermost portion of the Salmon River Formation (Lewis et al., 2001).

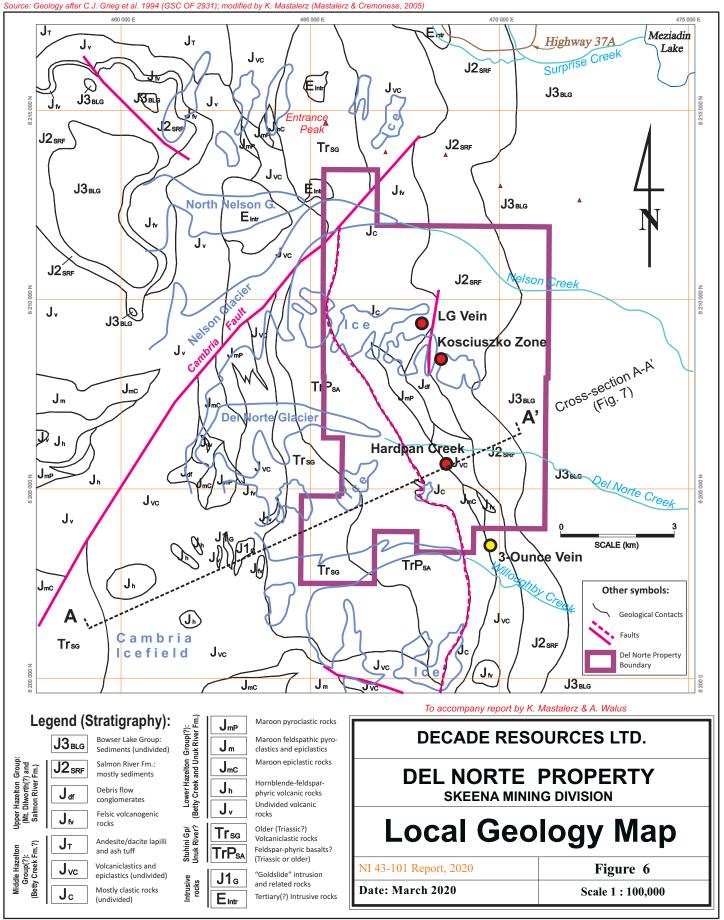
Thickness of all the formations of the Hazelton Group is highly variable within the Stewart complex, what is considered rather a very common attribute of numerous volcanogenic successions. The thickness of laterally equivalent units strongly depends on their location within an original volcanic-dominated environment which was characterized by development of very thick andesitic stratovolcano-type successions (proximal facies) which, in turn, have been laterally rapidly replaced by more distal volcanic facies and successively thinner units (cf. Alldrick, 1985).

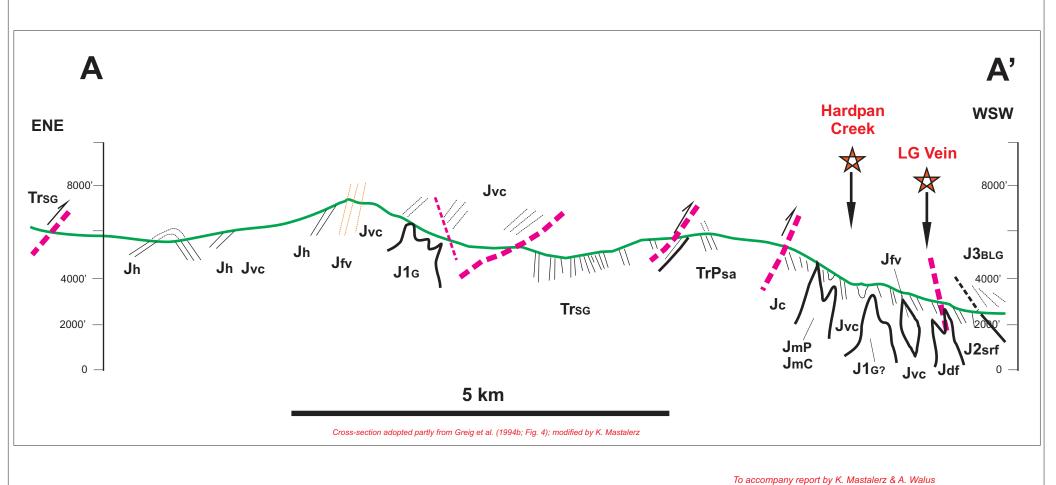
Apart of the layered rocks (volcanogenic and sedimentary) there occur several types of intrusive rocks which have invaded and locally altered the succession of the Hazelton Group in the Stewart complex area (see Alldrick, 1985). The oldest of these rocks is probably the Texas Creek granodiorite (Triassic/Jurassic), a hornblende, coarse-crystalline, locally feldspar-porphyritic rock. Similar in composition is the Summit Lake granodiorite, but this rock is generally finer-crystalline than the Texas Creek intrusive. The Premier Porphyry dykes, most likely of Early Jurassic age, are fine-grained feldspar-porphyritic diorite/andesite rocks. Similar in composition and texture are Mill porphyry dykes (Alldrick, 1985). The Boundary granodiorite is a biotite-minor hornblende, medium grained massive intrusive of Eocene, most likely, age. Of similar age is the Hyder quartz monzonite, as well as the Portland Canal dyke swarm (granodiorite, microdiorite and lamprophyre in composition).

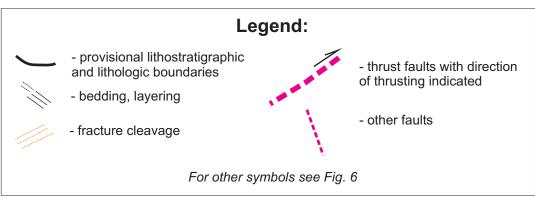
The Middle-to-Upper Jurassic Bowser Lake Group comprises the youngest Mesozoic sediments of the Stewart area and also of the Del Norte project. Generally, the Bowser Lake Group consists of thick, monotonous sequences of dark-grey to black, interbedded turbiditic sandstone (greywacke) and shale, and minor conglomerate. The latter end-member is locally characterized by abundance of well-rounded cherty clasts. However, it is still difficult to distinguish between lithologically similar sediments of the Salmon River Formation and the Bowser Lake Group. Lack of clear field criteria frequently causes inconsistent usage of both lithostratigraphic names and sometimes results in lumping of the strata of both units together (see discussion by Lewis et al., 2001).

The Salmon River Formation has been defined a bit more precisely in the Iskut-Eskay area where it appears obvious that its deposits are coeval with products of syndepositional volcanism (pillowed lava flows) and the Salmon River is topped with a characteristic "pyjama beds" of the Troy Ridge facies (Anderson and Thorkelson, 1990). In the Del Norte area, however, the distinguishing between these two lithostratigraphic units is very ambiguous. Greig et al., (1994a, b) are vaguely suggesting that their subdivision labelled "MUJ" roughly corresponds to the Salmon River Formation over the Cambria Icefield area. Locally this unit (MUJ) comprises relatively thin bedded to laminated, strongly siliceous siltstones/mudstones which are likely equivalent of the Troy Ridge facies (Evenchik et al. 1992 – the "Surpise Creek facies"). Evenchick et al. (1992) suggest that the contact between the Salmon River Formation and the Bowser Lake strata is gradational in the Surprise Creek – Yvonne Peak – Entrance Peak area (north of the Del Norte Property).

The area of the Del Norte project has never been systematically mapped. The early version of the map by E.W. Grove (1986; see also Devonck and Hardy, 1989) is far too generalized and inadequate. The later geological map completed by Greig et al. (1994a) has to be considered as the most comprehensive map for this area, so far. C.J. Greig and co-authors (1994a, b) provided a unified version of the geological map for the entire Cambria Icefield area (between Mt. Patullo in the north and Kitsault Lake in the south, and between the Portland Canal-Bear River valley in the west and the western edge of the Bowser Basin in the east. However, the lithostratigraphic nomenclature applied by these authors is not formalized and differs from the proposed formal one presented by Alldrick 1989, or other authors (e.g., Grove, 1986; Lewis 2001). A modified version of the map by Greig et al. (1994) became a choice selected by the authors of this report (Fig. 6).







DECADE RESOURCES LTD. DEL NORTE PROPERTY SKEENA MINING DIVISION Simplified Cross-Section NI 43-101 Report, 2020 Figure 7 Date: March 2020 Scale 1: 50,000

7.3 STRUCTURE AND PRINCIPAL ZONE

The most important mineral occurrences of the Del Norte property grossly cluster into two principal zones: the LG vein area and Hardpan Creek area (Fig. 3). Both of these zones are characterized by close spatial association with the contact between the Bowser Lake sediments and the upper part of the Hazelton Group volcanogenic rocks (Fig. 6). The LG vein itself locally follows the tectonic (thrust fault) contact between these two lithostratigraphic units almost perfectly. The Hardpan Creek area appears to be hosted entirely within somewhat deeper stratigraphic horizons, but remains still within the strata which tentatively have to be assigned to the upper part of the Hazelton Group (Fig. 7).

Although no systematic geological mapping was ever carried out over the LG vein area, the 2004 field work and drilling results (Mastalerz and Cremonese, 2004) allowed identification of a package of felsic volcanics which usually occurs between a thick succession of the intermediate volcanics (probably the Betty Creek Formation?) in the west and the thick succession of siliciclastic sedimentary rocks (Salmon River and/or Bowser Lake) to the east (Fig. 8). This felsic package has been tentatively assigned as a stratigraphic equivalent of Mt. Dilworth Formation (cf. Mastalerz and Cremonese, 2004) which previously became a regionally significant marker associated with the VMS deposit at Eskay Creek. The true thickness of this unit in Del Norte area varies between a few metres and few tens of metres. Predominant components of this unit are lapilli tuffs, felsic tuffs and coarser-grained tuff breccias. Layers of intermediate volcaniclastic rocks and tuffaceous sediments occur locally within this package. These rocks are locally rich in black, muddy sedimentary matrix. Layered rocks of the felsic package are locally cut by small-scale aphanitic to fine-crystalline intrusive rocks and hornblende-phyric dykes.

It appears that the felsic package lies conformably over the top of the succession of intermediate volcanic rocks and forms its stratigraphic continuation (Mastalerz and Cremonese, 2004). The concordant, lensoidal bodies of matrix-supported felsic lapilli tuff were observed between the intermediate volcanics and the uppermost portion of the Hazelton Group in the (Fig. 8). The felsic layers become progressively thicker up the stratigraphy until they become the dominant lithology. A few drill-hole intersections provide some evidence of a quite similar, transitional character of the upper contact of the felsic package. Locally lapilli tuffs are replaced upwards by their progressively more matrix-enriched varieties, redeposited volcaniclastics (with admixed epiclastic material) and finally texturally diversified, dark grey tuffaceous sediments. Thin layers of fine-grained siliciclastics appear only rarely. The compositional variability of this package suggests its transitional character as a lithostratigraphic member following the predominant accumulation of intermediate-composition volcanics and preceding a period of deposition of predominant siliciclastic material.

Volcaniclastic rocks of the Hazelton Group are overlain by a thick sedimentary succession which has been provisionally assigned to the Bowser Lake Group (cf. Mastalerz and Cremonese, 2004). This succession includes predominantly thin to medium bedded, black to dark grey siltstone, mudstone and argillite, and less common, somewhat thicker beds of sandstone and sedimentary breccia. The sedimentary strata strike N-S and steeply dip

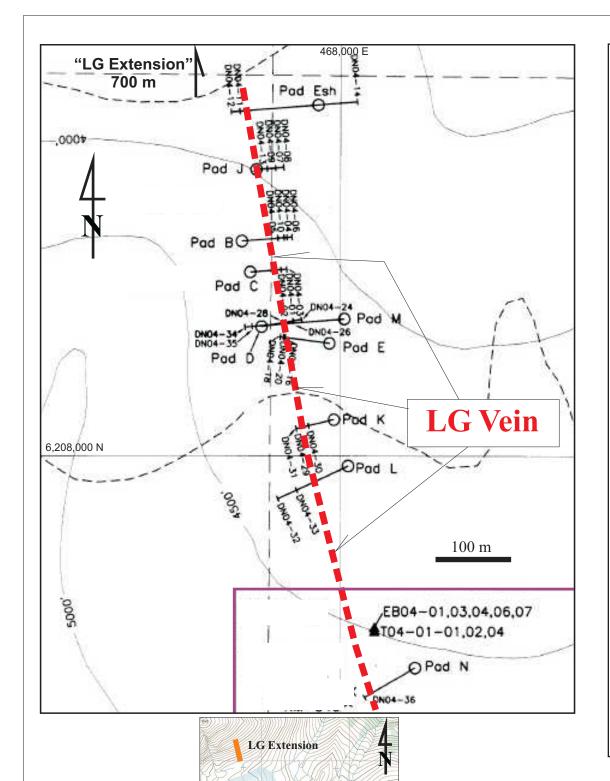
eastward but become subvertical and even slightly overturned at the very thrust fault contact with the volcanogenic package (Figs. 7 and 8). Further eastward, and upwards, within the Bowser Lake sedimentary succession, bedding becomes progressively shallower down to 20-30 degrees within a distance of 1-2 km east of the contact.

The contact between the Hazelton and Bowser Lake Groups locally has an evident tectonic character in the LG vein area. The contact trends roughly from north to south (Figs. 6 and 8) and locally displays a character of a steep, composite thrust fault zone with minor slices and duplexed tectonic slabs which line-up along the wide, sub-vertical thrust zone. Gentle to intense folding locally accompanies some faulted slabs and packages. Localized evidence of tectonic ductile and/or brittle style of deformation, as well as numerous slickensides, can be observed along the contact zone. The Kosciuszko Zone which appears to be distinct from the LG vein, displays especially strong evidence of such localized tectonic disharmonic folding, which is accompanied by strong shearing and slickensides.

Geometry of the contact zone is quite complex and the zone itself is apparently of multistorey character. Felsic volcaniclastics of the contact zone, as well as locally adjacent black sediments display evidence of ductile-to-brittle shearing which includes locally recognized CS structures, narrow zones of mylonitization, local development of incipient foliation, tectonic breccias and various degree of rock fracturing. All the tectonic planar structures strike meridionally-to-submeridionally and display vertical-to-subvertical dips. This deformation style appears to be almost penetrative within the felsic package and dies out away from both contacts of this unit.

Locally, geometry of the thrust zone involves segments of overturned stratigraphic footwall (the felsic package of the top Hazelton Group becomes the geometric hanging-wall of the thrust zone). Some features which accompany the thrust zone (geometry of the associate veins, fractures and component small-scale displacements) point to a hypothesis of inversely reactivated character of originally, probably just a prominent normal fault zone. This tectonic zone originated, most likely, in an extensional regime as a western, steep normal-faulted boundary of an incipient Bowser Basin (half-graben in character?). Subsequently, this deep rupture zone must have been reactivated in a compressional regime and followed by thrusting which led to a substantial inversion of the western part of the Bowser basin (cf. Mastalerz and Cremonese, 2004).

The prominent structural features mentioned above have been overprinted by subsequent and partly also synchronous deformations. The latter ones include NE-SW to W-E trending steep faults of, probably, strike-slip to oblique-slip character, which have developed, most likely, as deformations complementary to thrusting. These faults may be responsible for apparent displacement of some detached segments of the LG vein. It is not clear if the mineral occurrence originally described as the "LG Extension" (Cremonese 2004, see also Fig. 9) is such a detached segment of the proper LG vein or it is a distinct zone of mineralization in its stratigraphic footwall? The 2004 drill program provided compelling evidence that the "LG Extension" is hosted partly in the intermediate volcanics, significantly away from their contact with black sediments of the Salmon River Formation? (Fig. 9).

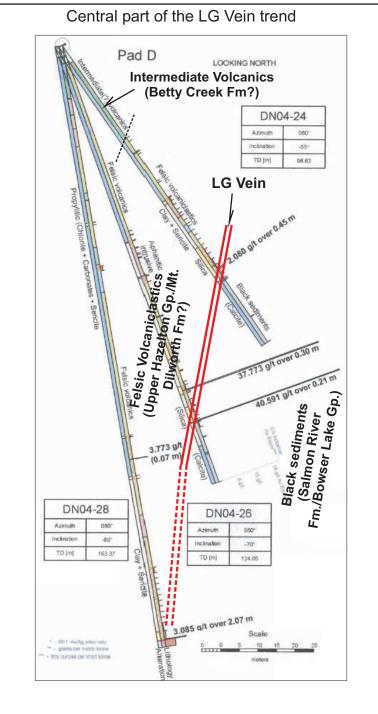


LG Vein

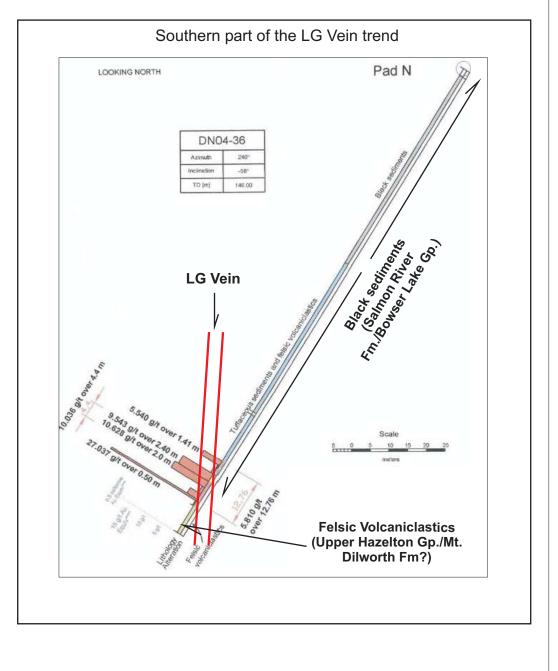
6,209,000 N

6,208,000 N

O Kosciuszko (K Zone)

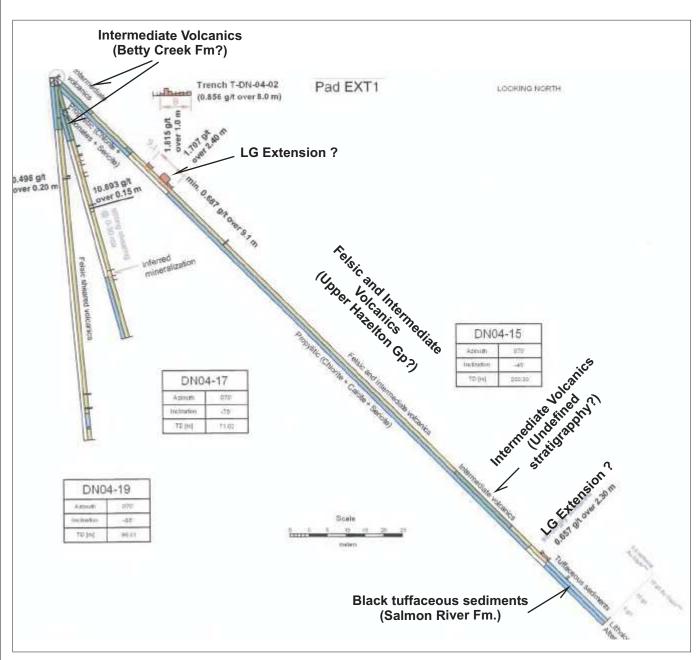






To accompany report by K. Mastalerz & A. Walus

DECADE RESOURCES LTD. DEL NORTE PROPERTY SKEENA MINING DIVISION LG VEIN 2004 Drill Hole Plan and Selected Drill Hole Cross-Sections NI 43-101 Report, 2020 Date: March 2020 Scale 1: 1,000





To accompany report by K. Mastalerz & A. Walus

DECADE RESOURCES LTD. DEL NORTE PROPERTY SKEENA MINING DIVISION "LG Vein Extension" Drill Hole Cross-Sections - Pad Ext 1 (2004) NI 43-101 Report, 2020 Figure 9 Date: March 2020 Scale 1: 1,000

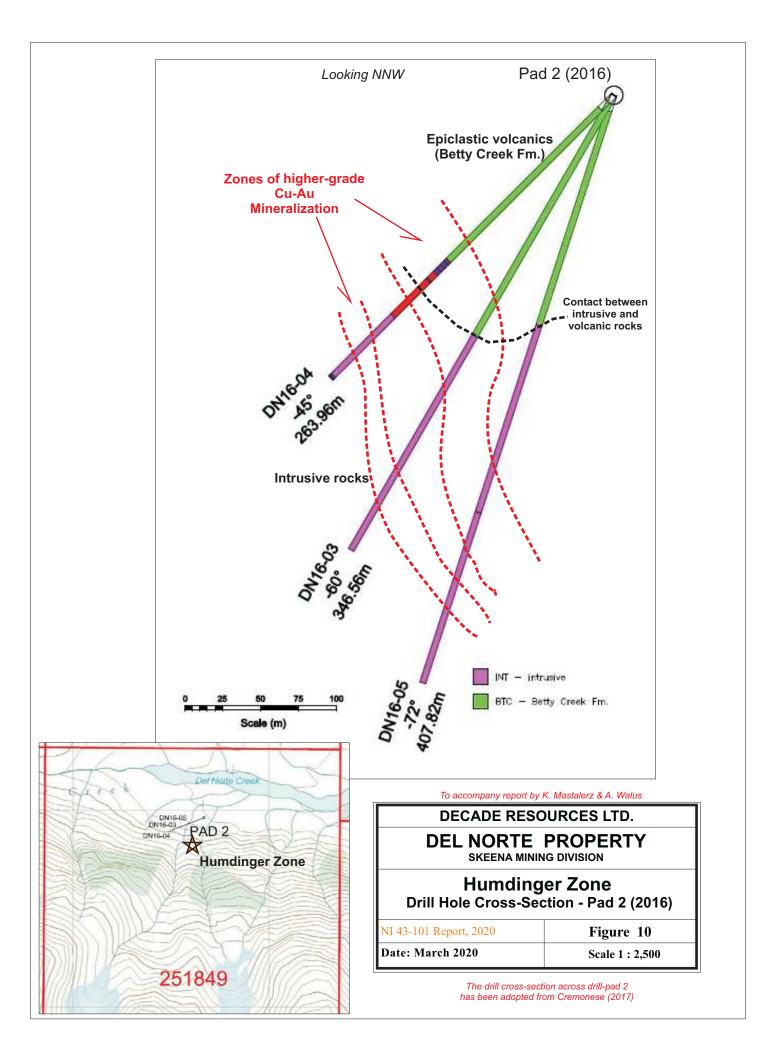
The drill cross-section across drill-pad Ext 1 has been adopted from Mastalerz & Cremonese (2004) The LG vein itself provides numerous evidence of its multistorey origin. It comprises several segments which display diversified brecciated textures and/or shear-related fabric (Mastalerz and Cremonese, 2004). Some of such breccia segments have apparently originated in the early stages of the vein emplacement and the angular vein fragments have been cemented by a quartz/carbonate precipitate still carrying some sulphide mineralization. The other segments display a typical, almost "dry" cataclastic texture which appear to be post-mineralization feature. A ductile-style shear fabric of the vein itself appears to be at least partly related to synmineralization stages. However, the very common overall shear-style deformations of the wall-rocks, especially the argillaceous rocks, were, most likely, induced and/or affected by post-vein emplacement tectonic deformations.

One of the latest-stage features in the LG area are NW-SE trending faults and open fissures, that have been locally used and filled by hornblede-rich dykes. This late stage of deformation can be assigned to up-warping of the Stewart Complex area, which was roughly coincidental with the emplacement of the intrusive rocks of the Coast Plutonic Complex.

Structural and stratigraphic features of the Hardpan Creek area are far less deciphered. The stratigraphic position of the mineral showings of this group may appear generally slightly lower than the Del Norte group showings (Figs. 6 and 7). Most of the Hardpan Creek showings are located roughly along the boundary between the generalized (undivided) packages labelled J_c (sedimentary rocks) and J_{vc} (volcaniclastics, epiclastics), that follows the original Greig's et al. (1994a, b) designation (Fig. 6). It would be probably correct if one includes the J_{vc} package (Figs. 6 and 7) to the Betty Creek Formation because of its thick, partly epiclastic component. Such an interpretation is certainly supported by presence of the adjoining packages labelled J_{mP} and J_{mC} (maroon pyroclastics and maroon epiclastics, respectively) which both fit perfectly to the classic descriptions of the Betty Creek Formation (see Chapters on Regional and Local Geology).

According to the original, although only tentative Greig's et al. (1994a, b) interpretation, the package J_c comprises predominantly "thin-bedded silty mudstone and siltstone, with local thin bedded sandstone and carbonate mudstone lenses" (op. cit. Greig et al., 1994b, pg. 53) and it was included by those authors in a more general category of "Cover rocks", i.e. rather younger than the Betty Creek, most likely of the Salmon River Formation. Quite a similar interpretation has been presented by the authors of the assessment report which summarized the extensive exploration program conducted on the Hardpan Creek area in 1990 (Bishop and Gal, 1991). These authors interpreted that the Del Norte property covers a northerly trending anticlinal structure with its core (central belt) build out of dominantly volcanic rocks of the Betty Creek Formation and flanked on both sides (west and east) by the sediments of the Salmon River Formation.

However, as it has been already stressed earlier the equivocal distinguishing of the Salmon River Formation can be very difficult and requires fulfillment of several criteria (see chapters on Regional and Local Geology). Some of the historic interpretations concerning recognition and mapping of this stratigraphic package have been considered controversial and sometimes



plain wrong where sedimentary packages which belong either to Unuk River and/or Betty Creek Formations were assigned to Salmon River Formation (see discussion in Lewis et al., 2001).

A generalized, interpretive cross section presented on Fig. 7 may shed some light on the general structural-stratigraphic relationships of the southern Del Norte property. This cross section is a compilation based, in great part, on the Greig's et al. (1994b) work but also includes information from various field projects conducted on the Del Norte property (Bishop and Gal, 1991, Mastalerz and Cremonese 2004, Cremonese, 2014, 2019). It demonstrates some effects of the late accretion-stage thrusting of the Stikine volcanic-arc succession over its, probably, late back-arc sedimentary succession of the Bowser basin.

Important features which obviously have influenced mineralization of the Hardpan Creek area are smaller-scale intrusive bodies which have been roughly mapped and/or intersected by numerous historic drill holes. A "potassium feldspar megacrystic monzonite sills" have been identified and mapped to the west of the O-Zone and Humdinger Zone (Bishop and Gal 1991). These sills were interpreted to dip moderately to the west and are mineralogically similar to the Texas Creek intrusives encountered frequently northwestward of Stewart. Several smaller-scale dikes have been intersected and identified in drill holes, similarly as the other intrusive bodies of diorite and microdiorite and of unspecified geometry (Bishop and Gal 1991). All five drill holes completed on the Copper-Gold Zone in 2014 (Cremonese, 2014) entered into prolonged intersection of Cu-Au mineralized, unspecified petrographically, sericite-altered intermediate intrusive. Also, several other drill holes completed in the Hardpan Creek area intersected contacts between intrusive and volcanogenic rocks, some of them with low-grade Cu-Au mineralization spanning the entire sections, and the best mineralization apparently crossing the contact (e.g. Fig. 10).

Bishop and Gal in their assessment report (1991) have identified some large-scale NNW-SSE to N-S striking steep faults which are sub-parallel to foliation and shear fabric in the Hardpan Creek area. They also noticed that some of these faults are accompanied by alteration of the wall rocks, which may be important for mineralization. Besides, the Hardpan Creek area features also a few steep, north-northeast trending faults which are post-alteration (most likely also post-mineralization) features and offset all older structures including mineralization zones (Bishop and Gal, 1991).

Mineralization of the Hardpan Creek area differs quite significantly from the one encountered in the proper LG vein area. Apart from several encounters of heavily Au-Ag-Zn-Pb mineralized veins (e.g. Bullion, Twighlight, Grizzly zones), the Hardpan area contains also wider areas of more uniform, disseminated, fracture-controlled, potentially bulk-type Cu-Au-Ag mineralization (e.g. O-Zone, Humdinger, Copper-Gold Zone).

Weak to moderate propylitic alteration is a common feature throughout the property including both main areas of mineralization mentioned above (Bishop and Gal, 1991, Mastalerz and Cremonese, 2004, Dickenson 2007). It occurs as common quartz and carbonate vein and veinlets, as well as penetrative chloritization, and minor carbonate-pyrite replacement of the

volcanic rocks. It is not clear if quite common epidote alteration is also a regional feature, but stronger epidotization is obviously related to local contact metamorphism (e.g. Gal and Sampson, 1991).

Zones affected by stronger propylitic alteration often display also sericitization and argillitization, which are commonly accompanied by finely disseminated pyrite and variable degree of silicification. Such alteration zones additionally affected by slow-rate, high-strain displacements resulted in development of sericite schist zones (cf. Gal and Sampson, 1991).

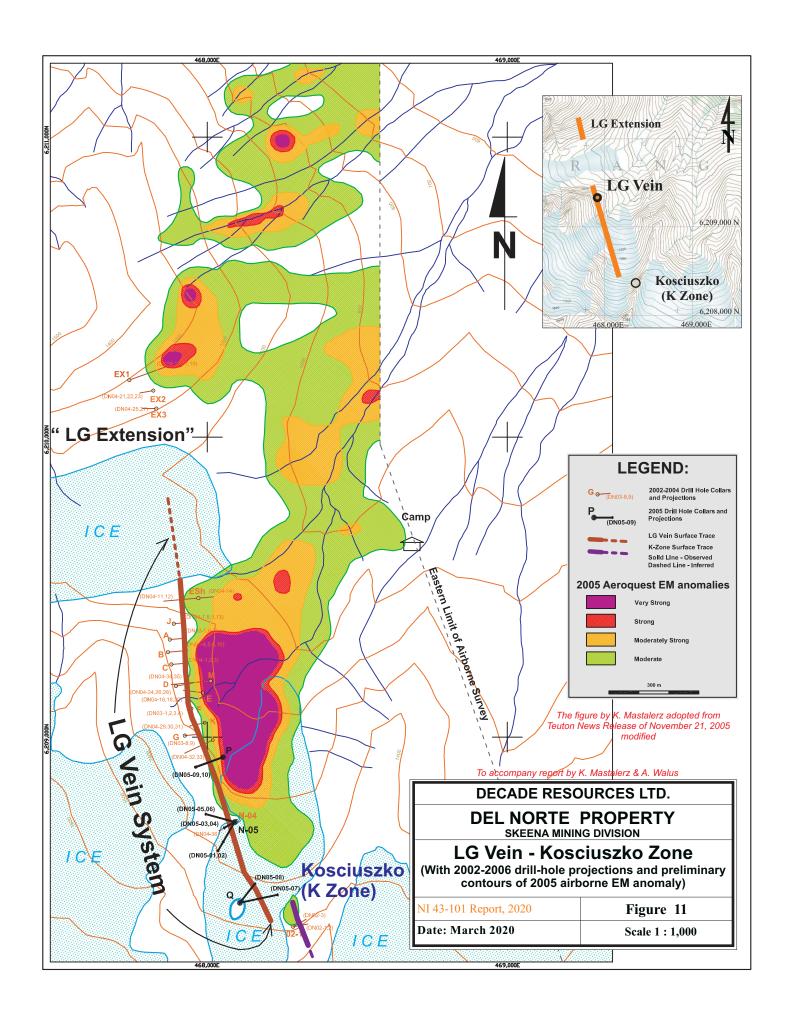
7.4 MINERALIZATION

The Del Norte property contains several mineralized zones. General location of these zones is shown on Figure 3. Their brief description is presented below:

LG vein, MINFILE No. 104A-161

The LG vein is a gold-silver bearing quartz-calcite vein which has been traced by drilling for 1.2 km and vertical extent of 240 metres (Fig. 11; see also Fig. 8). The true width ranges from 0.5 to 1.25 metres. Locally, the vein could be thicker as indicated by a 15.5 metre intercept reported in drillhole DN05-02 that assayed 6.58 grams per tonne gold and 350 grams per tonne silver (Exploration and Mining in BC 2005, page 34). However, the true thickness of this intersection is not known. The vein displays more regular character in the central part but usually thins and splits into swarms of thinner veins toward the north and south. Also, the vein displays numerous evidences of tectonic deformations; predominantly slickensides, brecciation and, locally, shearing and localized small-scale folding. It also appears that it is locally discontinuous what is believed to be caused by the tectonic displacements (Mastalerz and Cremonese, 2004).

The vein is spatially restricted to a narrow package of felsic volcaniclastic rocks and the footwall contact of the trust zone which are hosted within the contact zone between intermediate volcanic/volcaniclastic rocks of the Hazelton Group and argillites/mudstones of the Salmon River Formation. Sulphides include mostly pyrite, galena and sphalerite with trace to minor chalcopyrite, tetrahedrite, silver-sulphosalts and native gold. The latter two minerals often occur in black graphitic (?) matrix filling interstices between chalcedonic breccia fragments. Sulphide content is usually in the range of 1 to 5%. The vein is enveloped by distinct alteration halo (more pronounced on the west side of the vein) comprised of chlorite, pyrite, sericite and pervasive carbonate impregnations. Microscopic examination of several samples derived from LG vein hanging wall indicate that the rock represents plagioclase-quartz-biotite porphyry which could be the source of mineralization. Dyke of a similar rock was spotted along the bottom of a steep gully directly above Kosciuszko zone (Walus, 2003).



LG Vein Extension? (No MINFILE designation)

The showing represents a 10 cm wide quartz vein with minor galena, it strikes 300 degrees and has subvertical dip. Grab sample A04-154 collected from the vein in 2004 yielded 186.2 g/t gold, 1002.1 g/t silver and 0.65% lead (Mastalerz and Cremonese, 2004).

The vein, located some 800-900 metres toward NNW from the LG Vein was initially thought to represent the northern extension of the LG vein (Figs. 9 and 11). Eight holes from 3 different pads have tested the vein but returned rather low-grade gold and silver values (Mastalerz and Cremonese, 2004). However, the showing may not be a part of the LG vein as indicated by the fact that all mineralization encountered in these holes was not located within volcanogenic/sedimentary contact which hosts the LG vein.

Kosciuszko (Del Norte) Zone, MINFILE No. 104A 176

In the MINFILE as well as in many assessment reports Kosciuszko zone is also referred to as Del Norte and K zone (Figs. 9 and 11). The zone is not a part of LG vein, it is located some 60-70 metres NE from the presumed location of LG vein (Cremonese, 2003 – Fig 4). The zone is striking NNW-SSE and dipping 60 degrees to the northeast, what is roughly conformable with the bedding of the surrounding argillites/ siltstones. The zone is 50 metres long and from 3 to 10 metres wide. It is hosted in argillites of the Salmon River Formation approximately 40-50 metres from the contact between intermediate volcanic rocks of Hazelton Group and sedimentary rocks of Salmon River Formation (Cremonese, 2002). The zone is comprised of quartz-carbonate-sulphide cemented breccias, replacement zones and veins. Quartz- carbonates-sulphide material, which accounts for approximately 40-45% of the zone is intimately intercalated with short (10-70 cm) sections of unmineralized host rock. Breccia fragments, composed mostly of argillite are very angular ranging in size from less than 1 cm to 10 cm across. Occasionally there are vuggy cavities. Ganque minerals include quartz, carbonates and minor amounts of zoisite/clinozoisite. Carbonates include white siderite and possibly ankerite, the white colour changes to orange while exposed to the air. Sulphides include pyrite, sphalerite, galena and tetrahedrite along with trace amount of realgar, arsenopyrite and electrum (Walus, 2002). Photo of mineralized sample from Kosciuszko zone taken under petrographic microscope is presented on Photo 1. Part of pyrite occurs in host rock as disseminated grains comprising up to 5% of the rock. Content of sulphides vary widely from trace to 10%, but usually do not exceed 3%. The zone is cut by barren veins of coursegrained calcite, which occasionally contain minor pyrite. Host rocks include argillite, fine to medium grained tonalite, aphanitic dacite (?), and aphanitic to fine-grained andesite. All host rocks except argillite show weak to moderate sericitization and locally silicification. The zone also contains a few small areas of very strong clay-sericite alteration. A 2002 chip sample across the northern end of the zone returned a weighted average of 6.14 grams per tonne gold and 630.86 grams per tonne silver over a 10 metres interval (Teuton Resources Corp. News Release, September 26, 2002). Three holes drilled in 2002 from a single location all intersected the zone at depth. The best drilling result came from hole DN02-3 which returned a weighted average of 7.64 g/t gold and 277 g/t silver over 23.4 m (see also Fig. 18).

Au - native gold

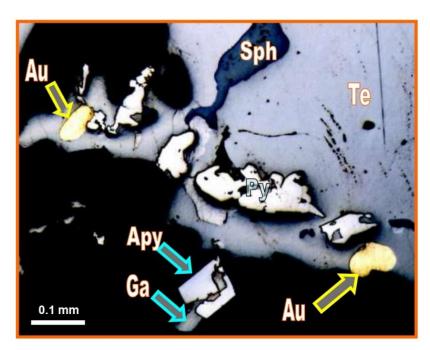
Apy – arsenopyrite

Ga – galena

Py – pyrite

Sph - sphalerite

Te – tetrahedrite



Photomicrograph, reflected light; Kosciuszko Zone

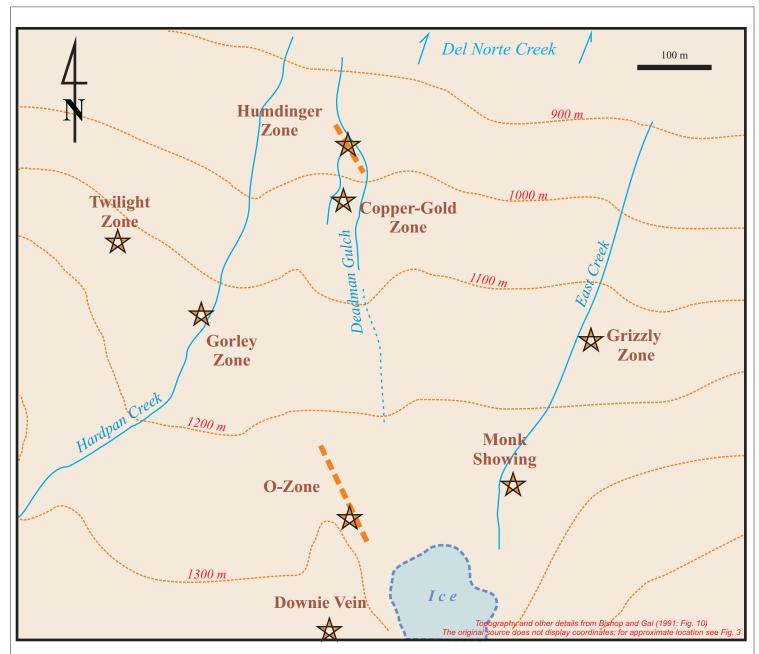
Photo 1. Microscope photomicrograph of the Kosciuszko Zone mineralization in reflected light (a core sample).

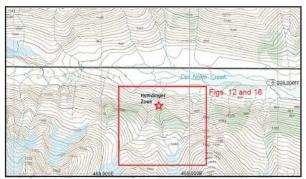
Hardpan, MINFILE No. 103P 318

The Hardpan Creek area contains eight mineral showings which were first evaluated in 1990 (Figs. 3 and 12). The mineralization is primarily fissure vein and replacement type. Geologic setting suggests the potential for vein, gold-copper porphyry and possible volcanogenic massive sulphide mineralization. Two types of mineralization are prominent, base metal and/or precious metal rich fissure type veins (+/- stringer stockwork) and gold-copper rich replacement type horizons. Mineralized zones are accompanied by strong alteration. Fissure type mineralization is by far the most commonly found and can be further subdivided into zinc-lead rich or coppergold rich veins or stringer-stockwork zones. The orientation of the veins and stringers coincides with the main fracture and fault systems, trending towards the northwest, northeast and rarely east. Description of the most prominent zones at this location is provided below:

O-zone

In this zone, disseminated and fracture filling chalcopyrite and pyrite (1 to 5 per cent) pervade a phyllic and +/- argillic altered horizon which has been overprinted by carbonate alteration probably related to faulting and fracturing. The horizon ranges between 6 and 15 metres in width and has been delineated for a strike length of 176 metres, open to the north and south, and a dip length of 35 metres, open to depth. The zone was targeted by 11 trenches and 4 drill holes. The best result from the trenches was 10.53 g/t gold and 0.23% copper over a 5 m





LEGEND:

Humdinger - Mineralized zone with label Zone

To accompany report by K. Mastalerz & A. Walus

DECADE RESOURCES LTD.

DEL NORTE PROPERTY

SKEENA MINING DIVISION

HARDPAN CREEK AREA Mineral Showings/Occurrences

NI 43-101 Report, 2020	Figure 12
Date: March 2020	Scale 1 : 5,000

interval. The best result from drilling was 3.67 g/t gold and 0.41% copper over 15 m interval (Assessment Report 21535). Mineralized halos often extend into carbonate altered hostrock and fracture systems.

Humdinger and Downie Zones

Gold-copper rich fissure style mineralization is associated with more obvious quartz veining and stringer networks in generally more felsic volcanics. Sulphides consist of chalcopyrite and pyrite, occurring mainly in the veins andd stringers. These "veins" range from discrete pods of quartz of less than 2 metres in length (Monk, East Cliffs, Copper-Gold) to more substantial vein systems up to 50 metres wide (Humdinger, Downie). Gold and copper values are relatively erratic within the veins. Concentrations range from 0.62 gram per tonne gold and 0.20 per cent copper over 11 metres at the Monk Showing (stringer stockwork) to 12.06 grams per tonne gold, 63.75 grams per tonne silver, and 400 parts per million copper over 6 metres at the Humdinger zone (Assessment Report 21535).

Twilight and Gorley zones

These lead-zinc rich zones occur as stockwork (Twilight, Gorley) as well as massive and banded zones (Grizzly). Disseminated, fracture filling, vein and banded sulphides and oxides consist of specular hematite, sphalerite, galena, pyrite and traces chalcopyrite. Stockwork systems average 1 to 1.5 per cent of combined zinc-lead over widths ranging from 3 to 75 metres. Grab samples of the less than 5 centimetre wide vein material can run as high as 27.54 per cent zinc, 13.2 per cent lead, 1.55 grams per tonne gold and 125.3 grams per tonne silver (Bishop and Gal, 1991).

Bullion, MINFILE No. 104A 077

The zone occurs in intermediate ash-lapilli and crystal tuffs near a belt of strongly phyllic and argillically altered volcanics and plagioclase-hornblende porphyry dikes. Float at the Bullion zone contained pyrite, chalcopyrite and galena. A one metre chip sample taken from the Bullion area in 1991 assayed 1.14 grams per tonne gold, 54.85 grams per tonne silver, 0.13 per cent lead, 0.012 per cent zinc and 0.003 per cent copper (sample DNLGR-306, Assessment Report 21535). The sample came from a fault zone with abundant quartz stringers.

In 1991, a quartz-sulphide vein, the NMG vein, was discovered about 150 metres north of the Bullion showing. The vein outcrops intermittently over a distance of 225 metres and lies along the same sedimentary-volcanic contact which host LG vein (104A 161) discovered in 1990. The highest assay came from a one metre chip sample taken across the vein at its southern exposure. The sample assayed 10.6 grams per tonne gold and 571.45 grams per tonne silver (sample DM-MR-82646, Assessment Report 22103).

Crackle zone (No MINFILE designation)

This zone, located on the northern side of the Del Norte glacier (Fig. 3) contains a network of quartz stringers /veins (approximately 6 per 3 m section) hosted in a series of blood-red

coloured, resistive knobs comprised of Betty Creek Formation rocks. Stringers/veins are 1 to 15 cm wide and contain medium to coarse grained chalcopyrite, pyrite with lesser amounts of magnetite and arsenopyrite. They have north-south strike with moderate (40-50 degrees) dip to the west. The zone is 50 by 100 metres in size but is partially obscured by glacial overburden. Soil sampling by Teuton Resources in 2017 extended the zone by 300 metres (Assessment Report 37704).

Nelson 2, MINFILE No. 104A 151

The showing is hosted in sericitized andesitic tuff containing 2 per cent pyrite and traces of malachite. A grab sample from the mineralized outcrop assayed 1.93 grams per tonne gold, 0.60 per cent copper and 3.4 grams per tonne silver (Sample #17105, AR 19424).

Just to the south of this showing eight float samples, containing semi-massive to massive pyrite and chalcopyrite, were taken. Seven of these samples were anomalous in gold ranging from 0.160 to 0.835 gram per tonne (Assessment Report 21813).

Nelson 2 South, MINFILE No. 104A 167

The showing is a north-south trending quartz vein hosted in a 1.5-metre wide zone of sheared, chloritic andesite. The vein contained up to 3 per cent coarse-grained galena with trace amounts of chalcopyrite. A 15 centimetres chip sample assayed 4.5 grams per tonne gold, 278.2 grams per tonne silver and 1.95 per cent lead (Sample 39921, Assessment Report 21813).

Nelson 3, MINFILE No. 104A 152

The showing in located on Lord Nelson 6 claim. An extensive gossanous area is developed in the andesitic volcanics on the south side of the Nelson Glacier. A sample of the limonitic rocks, containing 70 to 85 per cent pyrite and pyrrhotite, assayed 0.15 per cent copper, 0.18 gram per tonne gold and 5.6 grams per tonne silver (Assessment Report 19424).

Meziadin, MINFILE No. 103 P 005

The showing comprises partly silicified zone 0.4 to 0.9 metre-wide hosted in argillite. The quartz is locally copper stained and contains pyrite, galena and sphalerite with high values in gold and silver reported. Two other zones occur in the vicinity, they are 0.9 to 6.1 metres wide. They strike northwest, dip vertically and locally contain small silicified sections which contain small patches of sphalerite and chalcopyrite. A 1.46 metre sample from one of these two other zones, assayed 4.8 grams per tonne gold, 27.4 grams per tonne silver, 1.6 per cent copper, 1.6 per cent zinc and nil lead (Mandy, 1939). This was the best assay result from this location; most samples contained only trace gold and silver.

Croesus, MINFILE No. 104A 163

Few details are available regarding the showing comprised of 15-centimetre-wide quartz vein (Fig. 3). A grab sample collected in 1990 assayed 4.24 per cent zinc, 1.0 gram per tonne gold and 20.8 grams per tonne silver; copper and lead values are negligible (Sample DNMBR-012, Assessment Report 21535).

8. DEPOSIT TYPES

8.1 DEPOSIT TYPE

The presence of chalcedonic quartz, vuggy cavities, realgar, silver sulphosalts and electrum within LG vein and Kosciuszko zone (Walus; 2002, 2003) unequivocally point to a low-sulphidation gold-silver epithermal environment. Description of this deposit type extracted from British Columbia Mineral Deposit Profiles by A. Panteleyev of British Columbia Geological Survey is presented below.

Capsule description:

Quartz veins, stockworks and breccias carrying gold, silver, electrum, argentite and pyrite with lesser and variable amounts of sphalerite, chalcopyrite, galena, rare tetrahedrite and sulphosalt minerals formed in high- level (epizonal) to near-surface environments. The ore commonly exhibits open- space filling textures and is associated with volcanic-related hydrothermal to geothermal systems.

Depositional environment/Geological setting:

High-level hydrothermal systems from depths of ~1 km to surficial hotspring settings. Regional-scale fracture systems related to grabens, (resurgent) calderas, flow-dome complexes and rarely, maar diatremes. Extensional structures in volcanic fields (normal faults, fault splays, ladder veins and cymoid loops, etc.) are common; locally graben or caldera-fill clastic rocks are present.

Host/Associated rock types:

Most types of volcanic rocks; calcalkaline andesitic compositions predominate. Some deposits occur in areas with bimodal volcanism and extensive subaerial ashflow deposits. A less common association is with alkalic intrusive rocks and shoshonitic volcanics.

Deposit form:

Ore zones are typically localized in structures, but may occur in permeable lithologies. Upward-flaring ore zones centred on structurally controlled hydrothermal conduits are typical. Large (> 1 m wide and hundreds of metres in strike length) to small veins and stockworks are common with lesser disseminations and replacements. Vein systems can be laterally extensive but ore shoots have relatively restricted vertical extent. High-grade ores are commonly found in dilational zones in faults at flexures, splays and in cymoid loops.

Texture/Structure:

Open-space filling, symmetrical and other layering, crustification, comb structure, colloform banding and multiple brecciation. Ore mineralogy (Principal and subordinate): Pyrite, electrum, gold, silver, argentite; chalcopyrite, sphalerite, galena, tetrahedrite, silver sulphosalt and/or selenide minerals. Deposits can be strongly zoned along strike and vertically.

Gangue mineralogy:

Quartz, amethyst, chalcedony, quartz pseudomorphs after calcite; adularia, sericite, barite, fluorite, Ca- Mg-Mn-Fe carbonate minerals such as rhodochrosite, hematite and chlorite.

Alteration mineralogy:

Silicification is extensive in ores as multiple generations of quartz and chalcedony are commonly accompanied by adularia and calcite. Pervasive silicification in vein envelopes is flanked by sericite-illite- kaolinite assemblages. Intermediate argillic alteration [kaolinite-illitemontmorillonite (smectite)] formed adjacent to some veins; advanced argillic alteration (kaolinite-alunite) may form along the tops of mineralized zones. Propylitic alteration dominates at depth and peripherally.

Ore controls:

The veins are emplaced within a restricted stratigraphic interval generally within 1 km of the paleosurface. Mineralization near surface takes place in hotspring systems, or the deeper underlying hydrothermal conduits. Normal faults, margins of grabens, coarse clastic caldera moat-fill units, radial and ring dike fracture sets and both hydrothermal and tectonic breccias are all ore fluid channeling structures. Branching, anastomosing and intersecting fracture systems are commonly mineralized. Ore shoots form where dilational openings and cymoid loops develop, typically where the strike or dip of veins change. Hangingwall fractures in mineralized structures are particularly favourable for high-grade ore.

Geochemical signature:

Elevated values in rocks of Au, Ag, Zn, Pb, Cu and As, Sb, Ba, F, Mn; locally Te, Se and Hg.

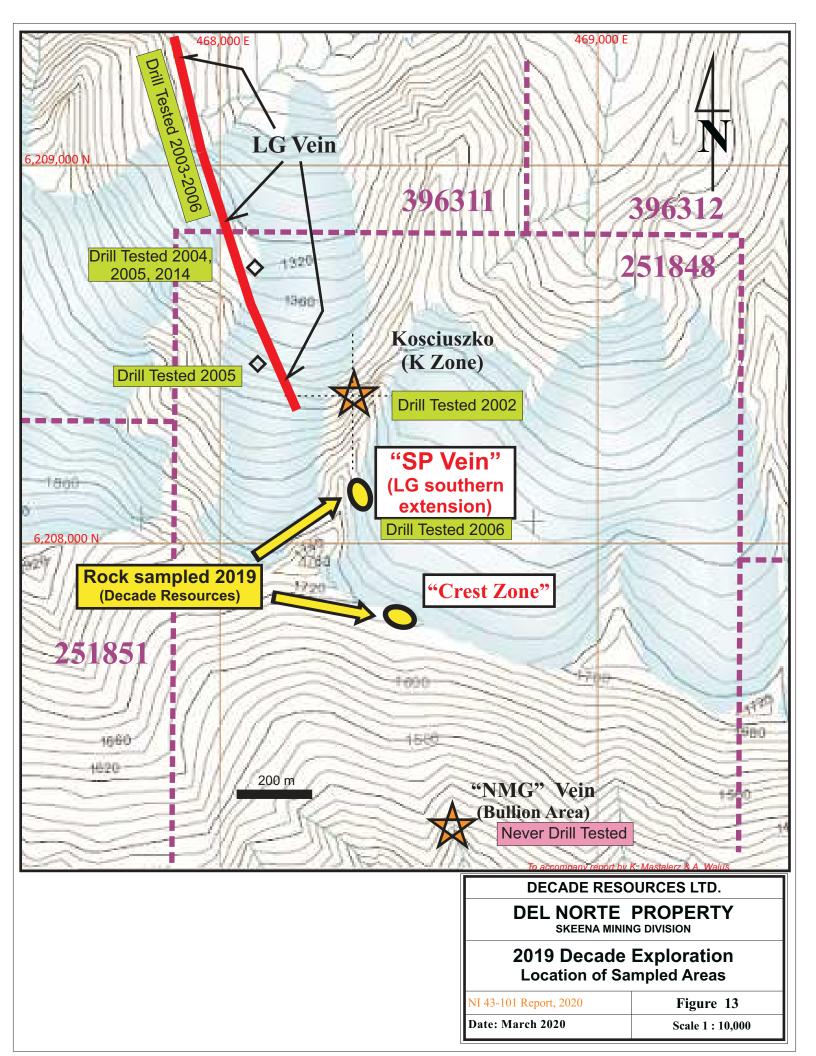
Typical grade and tonnage:

The following data describe the median deposits based on worldwide mines and U.S.A. models:

• Au-Ag deposits (41 Comstock-type 'bonanza' deposits) - 0.77 Mt with 7.5 g/t Au, 110 g/t Ag and minor Cu, Zn and Pb. The highest base metal contents in the top decile of deposits all contain 0.3% Cu; 10 % of the deposits contain, on average, about 0.75% Cu with one having >3.2% Cu.

8.2 GEOLOGICAL MODEL AND CONCEPT USED

The main mineralized zones of the Del Norte property are located within (LG vein) or close to (Kosciuszko zone) a narrow (a few dozen metres wide) horizon of felsic volcanics situated on the contact between intermediate volcanic rocks of Hazelton Group and sedimentary rocks of



Salmon River Formation. This geological model was successfully applied during previous and is planned to be applied in the future drilling campaigns.

9. EXPLORATION

On September 29, 2019 K. Mastalerz and A. Walus acting on behalf of Decade Resources conducted one day of sampling on the southern portion of the LG vein (originally called SP vein/zone) checking also for the potential further, southern extension of the vein as part of the Company due diligence (Fig. 13). The SP Vein was discovered during the 2006 exploration program on the Del Norte Property, which was under the option to Sabina Silver at that time. The surface grab sampling brought about from trace to 19.5g/t (0.57 oz/t) gold and trace to 3910g/t (114 oz/t) silver (Teuton, 2006a). The SP vein was drilled during the same field season and was successfully intersected at two various depths by two drill holes with encouraging results (Teuton 2006b; see also here Chapter "Drilling".)

In 2019, the authors briefly explored and sampled two distinct areas situated few hundred metres south of the K-Zone (Fig. 13, 14). In total, 21 samples were collected from the vein, its direct wall rocks and also from more distant wall-rocks where exposed. The sample locations and descriptions of the entire sampled population are presented in Table 3 while the corresponding laboratory results are collected and presented in Table 4 and Appendix 1. Sample locations are also presented on the maps of Figs. 14 and 15.

The first of the areas sampled during 2019 is located just south of the original SP vein discovery, where ice/firn recession during the last years facilitated local exposition of the SP vein for about 40 metres along its strike from under the relatively thin veneer of talus scree and moraine. The vein has been exposed in three distinct intervals (Fig.15). It attains approximately 1.1-1.2 metre in true width, strikes NNW-SSE to almost N-S and dips very steeply (at angle of 75-85 degrees) toward WSW to W. It is locally heavily mineralized with sulfides (pyrite, galena, sphalerite, arsenopyrite and chalcopyrite). Sulfides are commonly concentrated in moderately narrow (20-30 cm) zones along the contacts of the vein. However, the vein locally shows a brecciated, cataclastic internal texture which penetrates the central portions of the vein and is accompanied by strong sulfide mineralization. Southward, the bedrock exposure dives and is concealed under the significantly thicker cover of loose material of talus scree and/or lateral moraine.

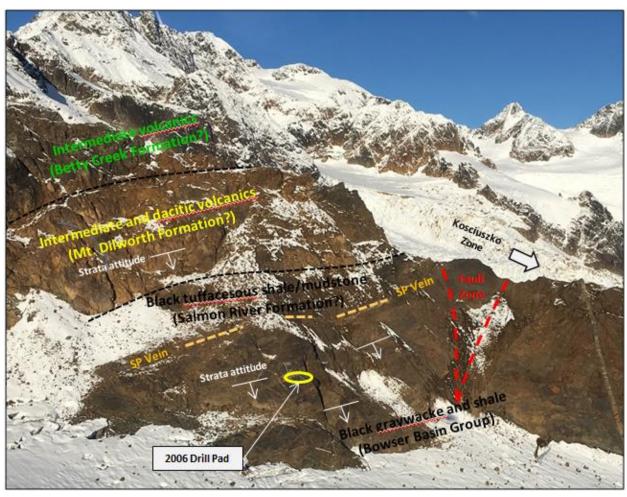


Photo 2. SP vein area sampled by the authors in autumn 2019 (view from the "Crest Zone"; looking northwestward).

Table 3. Decade 2019 exploration – Sample locations and descriptions

Sample		ates (NAD 33)	Sample	Description	Comment
label	Easting	Northing	type	·	
A19-258	468,379	6,208,127	Ch- 110cm	Chip across 1.1 metre wide quartz vein with 2-3% of combined galena, sphalerite and pyrite plus minor malachite stain. Orientation is N-S with vertical dip	SP vein
A19-259	468,376	6,208,142	Ch- 120cm	1.2 metre chip across quartz vein with 2-3% of combined galena, sphalerite, pyrite and malachite. Orientation 350°/80 W. The width of the vein is unknown as its footwall is obscured by talus.	SP vein
A19-260	468,567	6,207,765	Gb	Several fragments collected from suboutcrop of 0.5-1.0 m wide quartz vein with some limonite and wad stain.	SP Vein? (K-Dog)
DNK19- 01	468,374	6,208,120	Talus	Greenish andesite, quartz-carbonate alt'n	Talus
DNK19- 02	468,379	6,208,124	Gb	2-10 cm wide qtz vein in black tuffaceous sed's; 3-7 Py coarse crystalline, Apy 1%, tr. Sph	SP

DNK19- 03	468,379	6,208,123	Gb	Black tuffaceos mudstone (wall rock of DNK19-02); few thin qtz veinlets; Py 2-4, tr. Spy, tr. Ga?	SP
DNK19- 04	468,381	6,208,127	Gb	Quartz stockwork in black tuffaceous mudstone, 2-3 cm veins, irregular; Py 2-5%, tr. Apy, Sph	SP
DNK19- 05	468,382	6,208,124	Ch-30 cm	1-12 cm vein of banded quartz + 20-25 cm zone of quartz vlets in wallrock mudstone; Py 3-5%, Cpy 1-2%, Ga 1-4%, tr. azurite + malachite	SP
DNK19- 06	468,389	6,208,123	Gb	Quartz stockwork in tuffaceous mudstone (far east side); Py 1-4% coarse-crystalline and stringers, tr. Sph	East of SP
DNK19- 07	468,381	6,208,121	Ch-30 cm	Banded quartz vein (SP vein) 30cm chip, W side; abundant Ga, Sph, Py, tr. Cpy	SP vein
DNK19- 08	468,382	6,208,121	Ch-27 cm	Quartz vein to quartz breccia (SP vein) 27cm chip, E side; Py 3-5%, Sph 1, tr. Ga	SP vein
DNK19- 09	468,384	6,208,120	Gb	Quartz stockwork in black tuffaceous seds/greenish andesite; Py 2-3%	SP
DNK19- 10	468,379	6,208,130	Gb	Quartz vein/coarse quartz breccia (SP vein; = A19- 258); relatively abundant Py, Sph, Ga, tr. Cpy, Apy	SP vein
DNK19- 11	468,377	6,208,130	Gb	Quartz stockwork in black tuffaceous seds, west side of the SP vein; Py 1-4%, Sph tr1%, tr. Ga	SP
DNK19- 12	468,369	6,208,141	Talus	Andesitic black-matrix, broken-pillow breccia; matrix of tuffaceous mudstone; tr. Py, Sph	Talus
DNK19- 13	468,362	6,208,140	Talus	Andesitic black-matrix, broken-pillow breccia; matrix of tuffaceous mudstone; tr. Py	Talus
DNK19- 14	468,357	6,208,144	Gb	Incipient quartz-carbonate stockwork in strongly tuffaceous seds (far west from the SP vein); distinctly limonitic, tr. Ga?; strongly weathered	West of SP
DNK19- 15	468,362	6,208,112	Gb	Quartz veinlets scattered in black tuffaceous seds (far west from the SP vein); tr. Py, tr. Ga	West of SP
DNK19- 16	468,662	6,207,760	Gb	Zone of irregular quartz vein/pod (barren) in very strongly weathered andesite(?); no visible sulfides	K-Dog
DNK19- 17	468,611	6,207,750	Gb	Very strongly weathered, fractured andesite? (pinnacle of the ridge), locally chalcedony? Veinlets; tr1% Py	K-Dog
DNK19- 18	468,593	6,207,742	Gb	Strongly weathered andesite(?) with irregular quartz-carbonate pods/impregnations; Py tr1%; strong manganese stain	K-Dog

Abbreviations used: Py – pyrite, Apy – arsenopyrite, Cpy – chalcopyrite, Sph – sphalerite, Ga – galena; tr. – trace amount

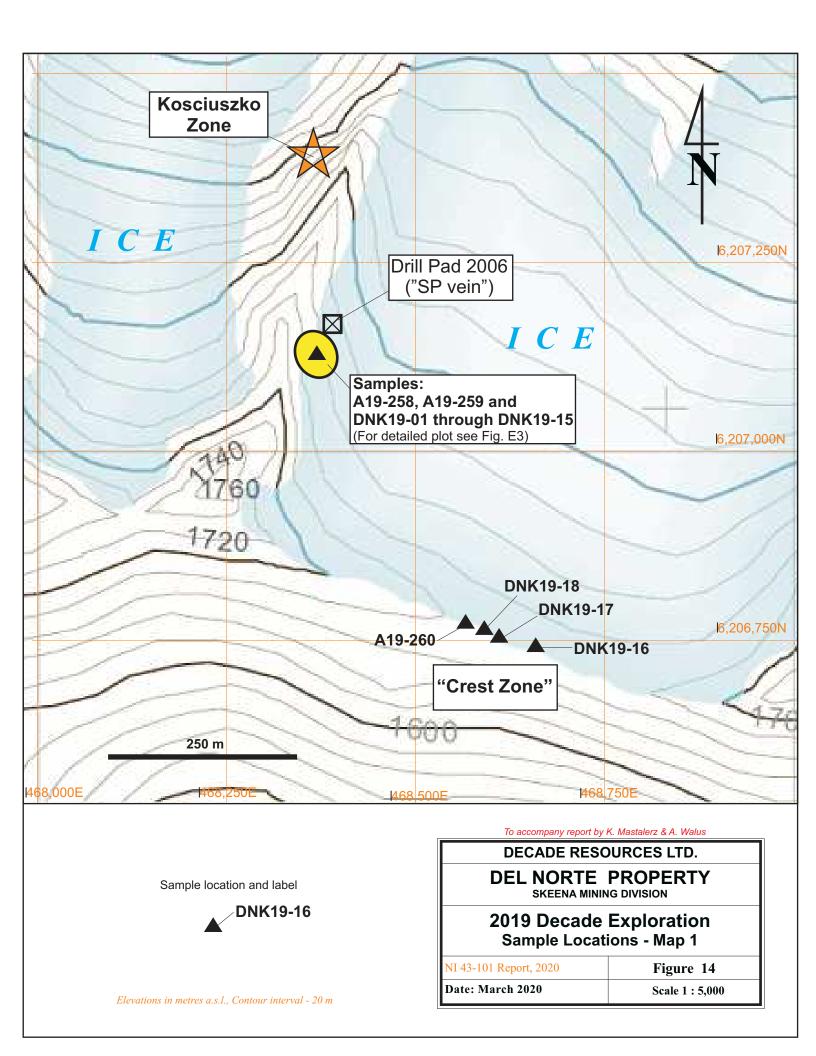
Two chip samples were collected across the entire width of the SP vein. Sample A19-258, a 1.1 m chip returned 6.59 g/t gold, 91.4 g/t silver, 2.06% lead and 2.01% zinc. Sample A19-259, a 1.2 m chip across the vein taken 20 metres north from A19-258 returned 751 ppb gold, 67.7 g/t silver and 0.95% lead (Table 4, Fig 15). Two other, narrower chip samples (DNK19-07 and DNK19-08) which also returned highly elevated concentrations of precious and base metals (Table 4) were taken from two near-wall contacts zones of the vein (Table 3) in the third exposed vein intersection, approximately 10 metres south of the A19-258 (Fig. 15). The fifth, grab sample (DNK19-10) collected from the central part of the vein which displays a strongly brecciated texture, also returned highly elevated amounts of gold, silver and base metals.

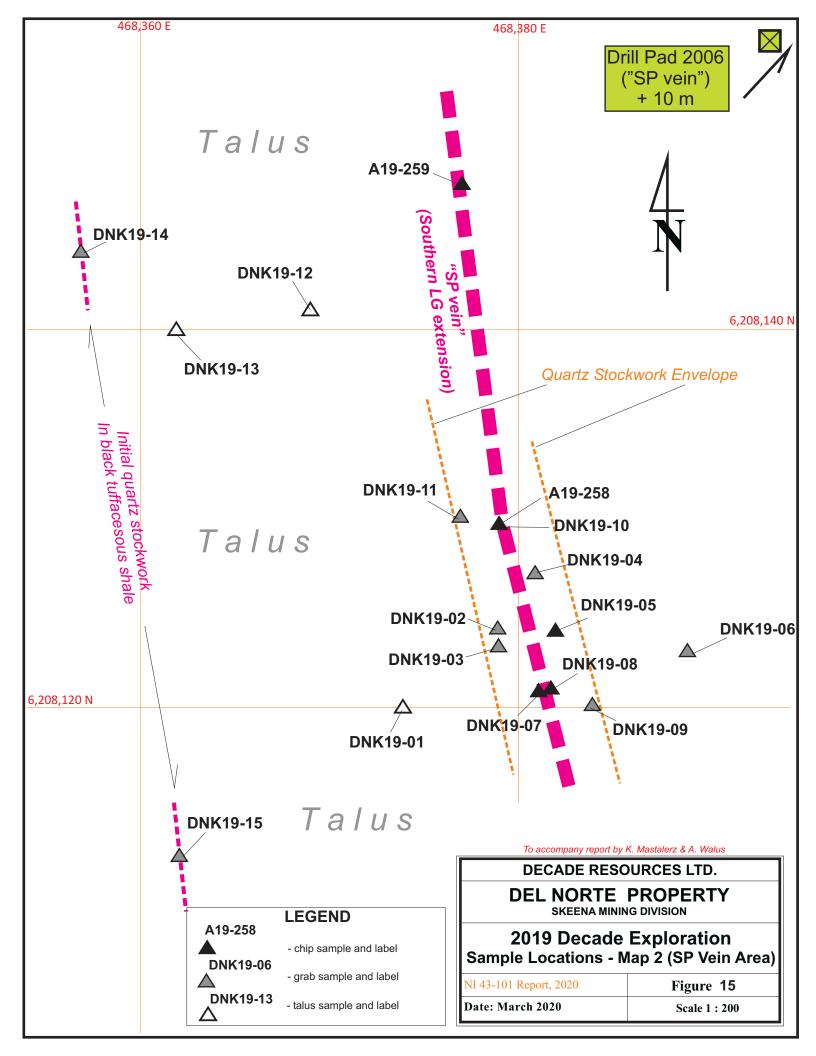
Six samples (DNK19-02, -03, -04, 05, -09 and -11) were taken from minor veins and semi-stockwork domains in near proximity (1-3 metres from the vein contacts) to the main LG (SP) vein on both its sides (Fig. 15). The host rock in these zones is represented predominantly by dark-grey to black tuffaceous mudstone/shale, although some andesite volcanics have been also encountered locally (Table 3). It is worth noting the the exposure of the area is limited and the precise description of relationship between particular lithologies will require additional clearing. All the above mentioned samples also returned moderately to highly anomalous gold, silver and zinc values (Table 4). The highest assays came from the 30-cm wide chip sample DNK19-5 which tested a banded, 10-12 cm wide quartz vein and its wall rocks, and yielded 19.6 g/t Au and 3920 g/t Ag. It appears that the banded vein is subparallel to the SP vein and is located about 2-3 metres east of it (Fig. 15).

Table 4. Decade 2019 exploration - Laboratory results (selected elements)

Sample	Au	Ag	Cd	Cu	Pb	Zn	As	Sb
label	ppb	ppm	ppm	ppm	ppm	ppm	ppm	ppm
A19-258	6.59 g/t	91.4	411.0	812	2.06%	2.01%	1300	793
A19-259	751	67.7	86.6	200	0.95%	4890	217	231
A19-260	462	25.8	20.5	39	1530	967	252	39
DNK19-01	19	2.6	0.8	15	21	107	74	8
DNK19-02	606	6.6	98.6	66	21	4010	2230	6
DNK19-03	72	1.6	12.6	48	37	957	167	7
DNK19-04	243	2.6	8.5	27	24	374	739	5
DNK19-05	19.6 g/t	3920	639.0	5820	1.91%	2.7	654	4670
DNK19-06	146	2.8	15.4	36	47	810	275	35
DNK19-07	6.21 g/t	87.9	531.0	2000	1.91%	2.51%	866	1560
DNK19-08	776	14.9	32.3	43	366	1830	527	25
DNK19-09	86	10.2	11.3	55	102	627	289	15
DNK19-10	7.53 g/t	794	378.0	770	3.17%	1.70%	1860	716
DNK19-11	684	88.2	35	108	882	1030	360	49
DNK19-12	9	0.8	1.3	109	22	157	26	8
DNK19-13	8	1.1	1.2	33	12	120	24	4
DNK19-14	123	15.3	6.3	41	1100	292	117	17
DNK19-15	323	44.6	8.0	118	1030	304	610	46
DNK19-16	7	< 0.2	0.6	4	12	75	30	6
DNK19-17	12	0.2	< 0.5	18	13	77	15	3
DNK19-18	69	< 0.2	1.5	11	12	69	227	6

Three other samples (DNK19-06, -14 and -15) were collected from a more distant wall-rock formation, away from the SP vein (Fig. 15). However, all these samples are represented by semi-quartz-stockwork host by the black tuffaceous sediments, similarly as the previous group. Also these samples displayed weakly, although significantly elevated concentrations of gold and





silver. Especially interesting are samples DNK19-14 and DNK19-15, which came from sites located about 20 metres westward from the SP vein (Fig. 15). These last two samples are also characterized by significantly elevated lead values (Table 4).

Three samples (DNK19-01, -12 and -13) were taken from large-size boulders of the local talus scree (Fig. 15, Table 3). These boulders appear to come from the bedrock outcrops along the steep cliff further east of the SP vein and represent variable, propylitic altered (regional?) andesite. These three samples displayed only very slightly elevated, probably not significant concentrations of gold and silver (Table 4).

The remaining four samples (A19-260, DNK19-16, -17 and -18) were collected in the second area visited during the 2019, a narrow ridge, a watershed between the valleys of Del Norte and Nelson Creeks (the "Crest area" - Fig. 13, 14). This area is situated approximately 350 metres south-southeastward of the SP vein sample A19-258, across a large ice-snow patch. The area is underlain by strongly weathered, commonly also strongly fractured dark-grey tuffaceous sediments, and pale-greenish-to-grey tuff and lapilli tuff. Few narrow patches of irregular, barren-looking quartz veins/pods and/or just a stronger silicification have been encountered along this ridge. Sample A19-260 taken from the "crest area" was the most significant. It assayed 462 ppb Au and 25.8 g/t Ag, which were accompanied by highly anomalous Pb, Zn and As contents. The sample was collected from sub-crop of a 0.5-1.0 m wide quartz vein which most likely represents the southern extension of LG vein. The sample DNK19-18 also displayed weakly, but significantly elevated concentrations of gold and arsenic (Table 4).

The 2019 reconnaissance program proved that in the SP area apart of the LG vein, there occur quite a broad zone of its wall rock (predominantly black tuffaceous sediments of, most likely, the Salmon River Formation), which is significantly mineralized. This zone is probably at least 2-3 metres wide on both sides of the vein. Besides, a few samples brought about evidence that also a bedrock more distant (up to 20 metres apart) from the vein is at least locally significantly mineralized.

The encountered mineralization displays a significant geochemical imprint and specific association of elements. It appears that the most elevated gold displays the most significant statistical association with some elevated indicator elements such as arsenic and antimony but also with copper (Table 4, Fig. 15). Its association with mercury, bismuth and/or tellurium appears to be quite random (see Appendix 1). Elevated values of silver are characteristically associated with high contents of zinc, lead and cadmium.

10. PREVIOUS DRILLING AND GEOPHYSICS

10.1 DRILLING

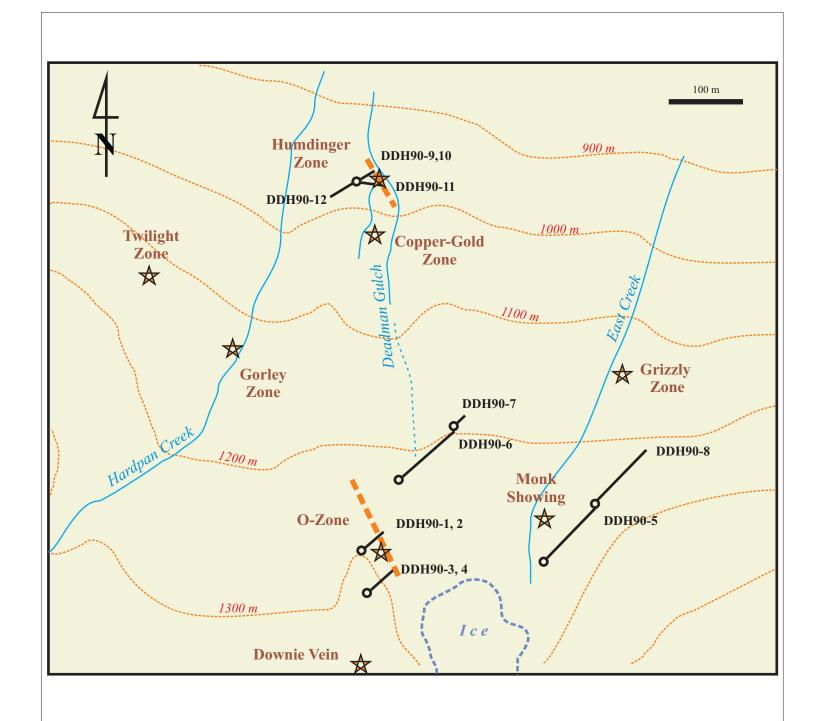
The Del Norte property witnessed several diamond drilling programs. International Kodiak

Resources carried out an exploration program on the property on behalf of Teuton Resources/Goodgold Resources during the 1990 summer season. Twelve diamond drill holes were completed and yielded 1118.3 m of BQTW core from the broad Hardpan Creek area. The holes were drilled from seven set-ups and tested subsurface extensions of the main Humdinger Zone, O Zone, the Monk and East Cliff showings, and two of them were drilled on overlapping geophysical and geochemical anomaly of Deadman's Gulch (Fig. 16). The laboratory testing of the core samples returned several strongly anomalous results (Table. 5). The most encouraging was a 3 feet (0.9144 m) long interval which assayed 0.43 opt (14.74 g/t) gold, 2.65 opt (90.86 g/t) silver and 1.16% copper. Most mineralization encountered and tested in the Hardpan Creek area was of vein-to-fracture filling type with transitional zones to replacement type. Overall, the results strongly indicated a potential for copper-gold porphyry system in the area.

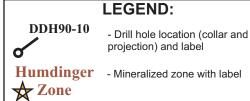
Table 5. Selected significant 1990 drill-hole intercepts (based on original data presented by Bishop and Gal, 1991: converted to metric system)

Distriop artu C	From	То	Interval	Aı		Aį	7	Cu	Comment
Drill Hole	m	m	m	oz/ton	g/t	oz/ton	g/t	%	
DDH90-01	18.29	33.83	15.54	0.107	3.67	0.39	13.37	0.41	O-Zone
incl.	21.64	23.47	1.83	0.236	8.09			0.268	O-Zone
incl.	30.48	32.31	1.83	0.277	9.50			0.872	O-Zone
DDH90-02	18.29	25.51	7.22	0.04	1.37			0.38	O-Zone
DDH90-03	54.86	76.20	21.34	0.016	0.55			0.329	O-Zone
and	80.77	81.99	1.22	0.43	14.74			1.16	O-Zone
DDH90-04	48.46	51.08	2.62	0.156	5.35	0.81	27.77	1.03	O-Zone
and	79.10	79.80	0.70	0.036	1.23			0.747	O-Zone
DDH90-05	9.14	10.97	1.83	0.043	1.47			311 ppm	see Fig. D16
and	91.44	94.06	2.62	0.076	2.61			1.02	see Fig. D16
DDH90-06	65.53	67.36	1.83	0.028	0.96				see Fig. D16
DDH90-07	15.24	18.59	3.35	0.037	1.27			488 ppm	see Fig. D16
DDH90-08				no significa	int results				see Fig. D16
DDH90-09	33.92	37.30	3.38	0.013	0.45			0.29	see Fig. D16
	52.69	56.69	3.99	0.01	0.34			0.15	see Fig. D16
DDH90-10	53.64	61.87	8.23	30 ppb	30 ppb			633 ppm	see Fig. D16
DDH90-11				no significa	int results				see Fig. D16
DDH90-12				no significa	int results				see Fig. D16

The most important drill-hole intersections of this program have been encountered in 4 holes which tested the "O-Zone" (Figs 16, Table 5). Holes DDH90-1 and DDH90-3 returned significantly elevated concentrations of gold and copper over 15.54 and 21.34 metre-long intervals, respectively. These intervals have been interpreted as direct continuations of the O-Zone which has been tested by numerous trenches at surface (Bishop and Gal, 1991; see also Fig. 17). The further westward, subsurface continuation of the Au-Cu mineralization of the O-



Topography and other details from Bishop and Gal (1991: Fig. 10)
The original source does not display coordinates; for approximate location see Fig. 3



To accompany report by K. Mastalerz & A. Walus

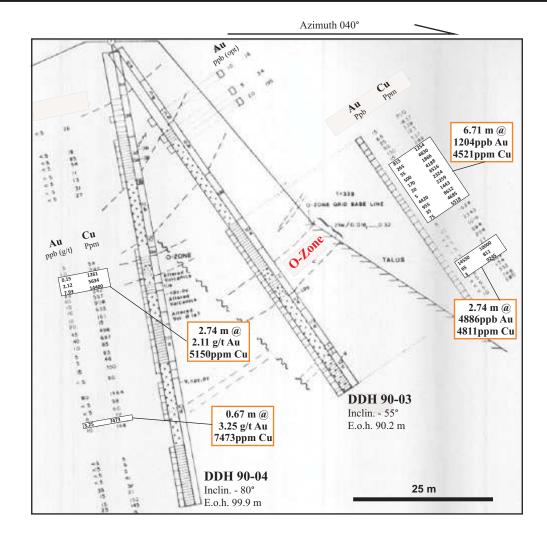
DECADE RESOURCES LTD.

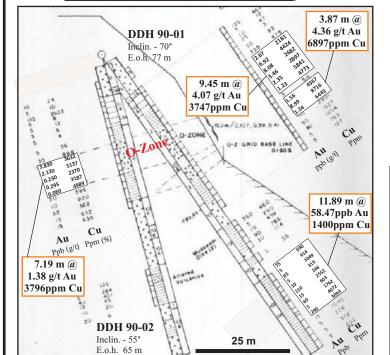
DEL NORTE PROPERTY

SKEENA MINING DIVISION

HARDPAN CREEK AREA Drill Hole Locations (1990)

NI 43-101 Report, 2020	Figure 16
Date: March 2020	Scale 1 : 5,000





Azimuth 040°

Cross sections modified from Bishop and Gal (1991: Fig. 15) (the original source quoted lab results in various units (ppb, ppm, g/t, oz/ton); here the significant lab results are displayed in metric-compatible units (ppm, ppb, g/t)

To accompany report by K. Mastalerz & A. Walus

DECADE RESOURCES LTD. DEL NORTE PROPERTY SKEENA MINING DIVISION HARDPAN CREEK AREA Drill Hole Cross-Sections DDH90-1, 2 and DDH90-3, 4 (1990) NI 43-101 Report, 2020 Figure 17

Scale as shown

Date: March 2020

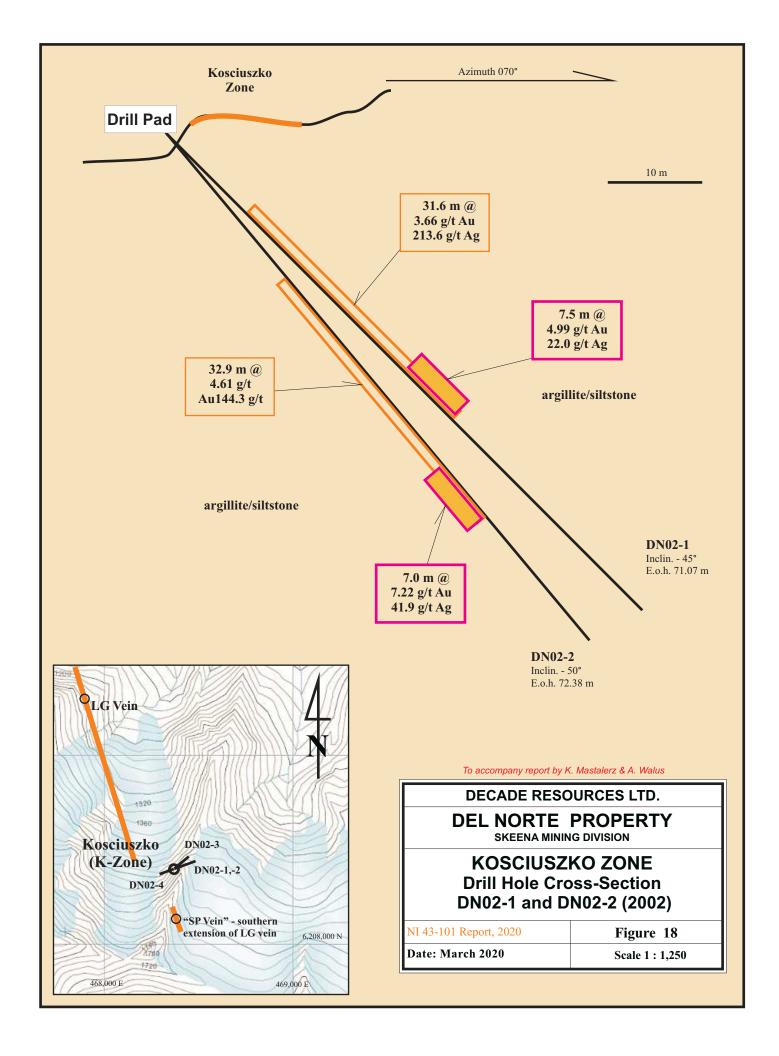
Zone is significantly less evident; the zone appears to be displaced and obscured by faulting (Fig. 17). Interestingly, the anomalous concentrations of gold and copper occur within the limits of a much wider (true width) interval than that was originally concluded from surface observations at O-Zone. Besides, the drill holes DDH90-1 and DDH90-3 intersected additional, approximately 30 metre wide intervals of anomalous concentrations of copper and, minor gold, at greater depth, significantly beneath the proper O-Zone (Fig. 17). The host rocks of these intervals were strongly altered volcanics with K-spar rich monzonite dykes.

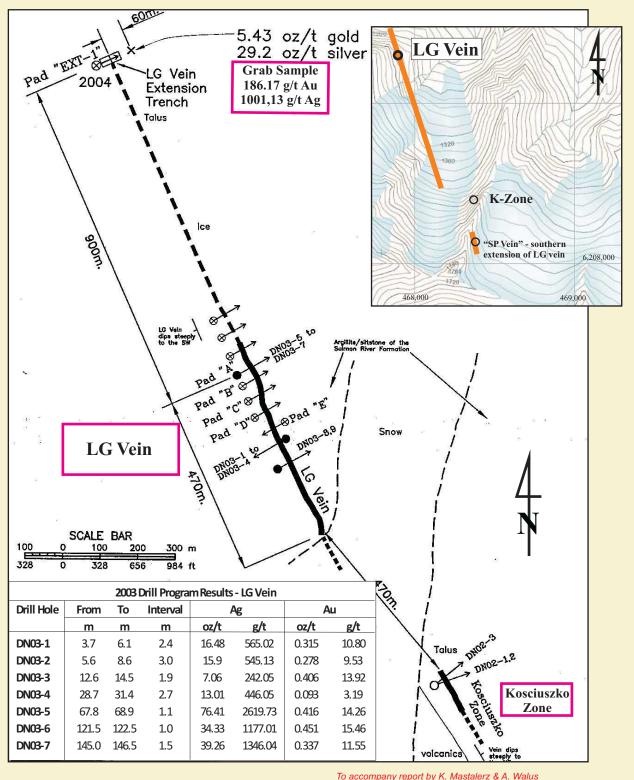
Some of the remaining drill holes of the 1990 exploration program on the Hardpan Creek area returned only narrow intervals with significant copper-gold mineralization, and three of them did not encountered any significant mineralization (Table 5). The drill holes DDH90-5 and DDH90-8 tested the Monk Zone, holes DDH90-6 and DDH90-7, investigated the contact between intermediate and felsic volcanics following the Deadman Gulch, and holes DDH90-9 through DDH90-12, tested the Humdinger Zone (Fig. 16).

Unfortunately, the collars of the drill holes were not surveyed and the location map of the corresponding assessment report does not display UTM or geographic coordinates on it (see Gal and Sampson, 1991 - Fig. 3). The drill collars have been tied to the local grid lines (Bishop and Gal, 1991). The reported results (Bishop and Gal, 1991: Drill Logs) indicate that the rock formations which were intersected at various sites of the Hardpan Creek area displayed more extensive intervals of anomalous concentrations of gold and copper than it has been reported as the "Diamond drill hole significant results" (Bishop and Gal, 1991: Tables IV through VII; see also: Gal and Sampson, 1991: Table 5). It is important to note that, the original laboratory results do include consistently only gold and copper concentrations; only a few selected intervals (core samples) have been assayed for silver and zinc (Bishop and Gal 1991: Drill Logs).

In the summer of 2002, four drill holes have been collared from a single drill pad (Fig. 18; see also Fig. 11) at an elevation of about 1530 m a.s.l.,to test the newly discovered Kosciuszko Zone (K-Zone). The K-Zone was discovered in 2002 along a very steep cliff and as chip-sampled at surface it returned 0.179 oz/ton (6.14 g/t) gold and 18.4 oz/ton (630.86 g/t) silver over an estimated width of 10 metres (Teuton, 2002). The first three drill holes tested potential extension of the zone at shallow inclination toward NE-NNE and all successfully intersected the subsurface continuation of the zone at depth with very encouraging results (Table 6). The cross-section of the first two holes is presented here as Fig. 18. The last drill hole was directed toward WSW and was terminated at depth of 33.24 metre without intersecting any mineralization.

The attempted second drill pad was never completed due to technical difficulties and very steep slope (Cremonese, 2003). The lateral extent of the Kosciuszko Zone has never been evaluated with certainty due to difficult access. As seen from the base of the cliff the zone appears to be involved in strong tectonic deformation including faulting and non-coaxial shearing. The mineralization remains open at depth.





The map with 2003 drill locations modifed from Teuton (2004)
- News Relase of Aug. 19, 2004
Lab results originally displayed in oz/t (troy ounce per short ton (US ton)); provided here are the corresponding results in metric system (g/t)

DECADE RESOURCES LTD. DEL NORTE PROPERTY SKEENA MINING DIVISION LG VEIN 2003 Drill Hole Locations and Results NI 43-101 Report, 2020 Figure 19 Date: March 2020 Scale 1: 10,000

Table 6. Significant results of the 2002 diamond drilling program on the Kosciuszko Zone (according to: Cremonese, 2003).

Drill	Interval	Width	Αι	ı	,	Ag
Hole	m	m	g/t	oz/ton*	g/t	oz/ton*
DN02-1	11.9-43.5	31.6	3.66	0.104	213.6	5.61
incl.	36.0-43.5	7.5	4.99	0.133	622.0	15.96
DN02-2	19.8-52.7	32.9	4.61**	0.134	144.3**	5.22
incl.	33.0-40.0	7.0	7.22**	0.21	451.9**	13.18
DN02-3	1.3-24.7	23.4	7.46***	0.223	277.3	8.09
incl.	16.0-24.7	8.7	7.49***	0.219	508.4	14.82
DN02-4			no significa	ant results		

Remarks: * - results as reported originally by Cremonese (2003); ** - original lab certificates not available, values calculated from the results quoted in drill logs (Cremonese, 2003); *** - gold results provided only in ppb in original lab certificates (Cremonese, 2003)

The attempted second drill pad was never completed due to technical difficulties and very steep slope (Cremonese, 2003). The lateral extent of the Kosciuszko Zone has never been evaluated with certainty due to difficult access. As seen from the base of the cliff the zone appears to be involved in strong tectonic deformation including faulting and non-coaxial shearing. The mineralization remains open at depth.

During the summer, 2003, Lateegra Resources Corp. which optioned the Del Norte property from Teuton, completed a program of 9 diamond drill holes which tested the LG Vein exposed at a moderately high cliff some 700 metres NNW from the Kosciuszko Zone (Figs. 11 and 19). Seven of the completed holes successfully intersected the vein for a 1.0 to 3.0 metre-long interval (apparent width) at variable depths, and returned significantly elevated, very encouraging concentrations of silver and gold (Fig. 19). The most significant results included 3.05 metres interval of 15.90 oz/ton (545.13 g/t) silver and 0.278 oz/ton (9.53 g/t) gold from the hole DN03-2 and 1.1 metre-long interval of 76.41 oz/ton (2619.73 g/t) silver and 0.416 oz/ton (14.26 g/t) gold from DN03-5. The results of this campaign were only reported in a form of Company's News Release (Teuton, 2004a, b) and are not supported by a standard assessment report. The following table summarizes the laboratory results of the 2003 drill program.

Table 7. Significant results of the 2003 diamond drilling program on the LG vein (*according to*: Cremonese, 2004).

			LG VEIN	DRILL HO	DLE SUMI	MARY (20	003)			
Drill	Distance	Drill Hole	Dip Angle	From	То	Interval	Silv	ver	Go	old
Pad	to K Zone	Number	degrees	m	m	m	oz/ton	g/t	oz/ton	g/t
		2003-01	-45	3.72	6.10	2.38	16.48	565.0	0.315	10.80
_	671 m	2003-02	-55	5.61	8.66	3.05	15.9	545.1	0.278	9.53
F	671 m	2003-03	-65	12.65	14.51	1.86	7.06	242.1	0.406	13.92
		2003-04	-77.5	28.68	31.42	2.74	13.01	446.1	0.093	3.19

		2003-05	-60	67.79	68.88	1.10	76.41	2619.7	0.416	14.26
Α	911 m	2003-06	-70	121.49	122.50	1.01	34.33	1177.0	0.451	15.46
		2003-07	-75	144.99	146.49	1.49	39.26	1346.0	0.337	11.55
G	600 m	2003-08	-56			No sig	gnificant r	esults		
l G	000111	2003-09	-70			No sig	gnificant r	esults		

In the summer of 2004 Teuton and Lateegra Resources mounted a much more extensive exploration program which included 36 diamond drill holes (two drilling teams were hired for that contract) designed to test the LG Vein in details (Mastalerz and Cremonese, 2004; Figs. 11 and 20). The holes yielded 4,815 metres of BQ core and 771 core samples which were assayed. The program provided significant amount of data which has allowed for better understanding of the extent and character of mineralization. Strikingly, numerous drill holes intersected the LG vein following more or less the prominent lithological contact between predominant intermediate and felsic volcanic and volcanioclastic rocks in the west with black tuffaceous mudstone and argillites in the east (Fig. 21). This contact is inferred to coincide with the prominent lithostratigraphic boundary between the older, predominant volcanogenic rocks of the Betty Creek-to-Mt. Dilworth Formations with the younger sediments and tuffaceous sediments of the Salmon River Formation. The contact (and the vein) is steeply dipping westwards (Fig. 21), plunging under the older stratigraphic formations. The contact zone carries abundant evidence of variable tectonic displacements including shear surfaces, fracture cleavage, tectonic breccia, tectonic groves and slickensides.

The set of the significant laboratory results from the 2004 drill program is presented in the table below.

Table 8. Significant results of the 2004 diamond drilling program on the LG vein (*according to*: Mastalerz and Cremonese, 2004).

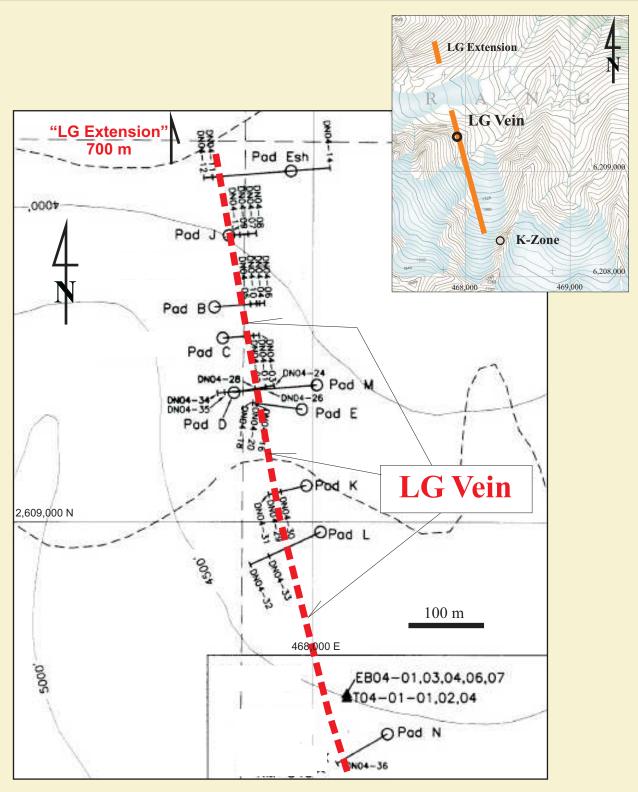
Distance to K Zone –Pad Name	Drill Hole	Dip Angle	From	То	Interval	Silv	ver	G	old
m		degrees	m	m	m	oz/ton	g/t	oz/ton	g/t
325N	2004-36		127.59	136.18	8.60	7.89	270.5	0.088	3.02
	2004-32	-50	101.44	103.78	2.35	28.96	992.9	0.24	8.23
545L	2004-33	-60	142.43	143.80	1.37	27.56	944.9	1.051	36.03
	2004-29	-51	70.90	77.05	6.16	1.87	64.1	0.059	2.02
	2004-30	-65	Но	ole did not	reach targ	et depth du	e to technic	al probler	ns
610K	2004-31	-72	147.19	148.10	0.91	32.56	1116.3	0.201	6.89
	2004-16	-45	66.14	66.63	0.49	53.73	1842.1	0.49	16.80
	2004-18	-55	83.00	84.09	1.10	16.29	558.5	0.43	14.74
	2004-20	-62.5	118.35	127.10	8.75	14.86	509.5	0.206	7.06
705E	Includes		118.35	121.13	2.77	41.85	1434.8	0.571	19.58
735M	2004-34	-55	220.19	221.28	1.10	1.17	40.1	0.036	1.23

	2004-35	-65	40.57	43.01	2.44			0.152	5.21
	2004-24	-55			No s	ignificant re	sults		
	2004-26	-70	97.78	98.08	0.30	55.01	1886.0	0.185	6.34
	and		106.60	106.80	0.20	0.7	24.0	56.1	1923.40
760D	2004-28	-80	161.30	163.37	2.07	1.81	62.1	0.06	2.06
	2004-01	-65	96.62	97.35	0.73	27.95	958.3	0.27	9.26
	2004-02	-75	148.96	152.55	3.60	8.67	297.3	0.244	8.37
830C	2004-03	-60		Le	ss than 20	% core reco	very in zon	e	
	2004-04	-60	103.91	107.05	3.14	18.69	640.8	0.358	12.27
	2004-05	-70		Le	ss than 20	% core reco	very in zon	e	
	2004-06	-75		Le	ss than 20	% core reco	very in zon	e	
2,870B	2004-10	-65	113.39	114.79	1.40	5.72	196.1	0.143	4.90
	2004-07	-60	63.98	65.04	1.07	9.21	315.8	0.262	8.98
	2004-08	-70	79.89	81.29	1.40	18.67	640.1	0.201	6.89
	2004-09	-77.5							
875J	2004-13	-85	95.01	96.19	1.19	17.1	586.3	0.146	5.01
1010—	2004-11	,	\			f-:11+		: /£	.11
E-Sh	2004-12	<i>,</i>	ali noies tro	in this pad		railed to rea ements?)	ich target h	orizon (tai	uit
	2004-14				aispiac				

The 2004 diamond drilling program included also several drill holes completed on the "LG Extension Zone" which was discovered in 2003 some 700-1000 metres toward NNW from the central part of the LG Vein showing (Figs. 11, 19 and 20). The grab sample of 20 cm wide quartz vein collected near a trench which was excavated in that area returned 5.43 oz/ton (186.17 g/t) gold and 29.2 oz/ton (1001.13 g/t) silver. Eight holes which were completed in that area from 3 different pads have returned rather low-grade gold and silver over variable-length intervals with a few thin intervals (0.1-0.2 m) of quartz vein intersections (e.g. Fig. 10) which displayed high grade of precious metals (Mastalerz and Cremonese, 2004). The best, 14.95 metre-long interval which was intersected in drill hole DN04-25, has returned 0.554 g/t Au equivalent (using 60/1 gold/silver price ratio at that time). It is worth noting that all the mineralization intersected on the LG Extension zone is host by intermediate-to-felsic volcanogenic rocks, substantially far from the contact with black tuffaceous sediments to the east (Mastalerz and Cremonese, 2004: Figs. 11, 13 and 15).

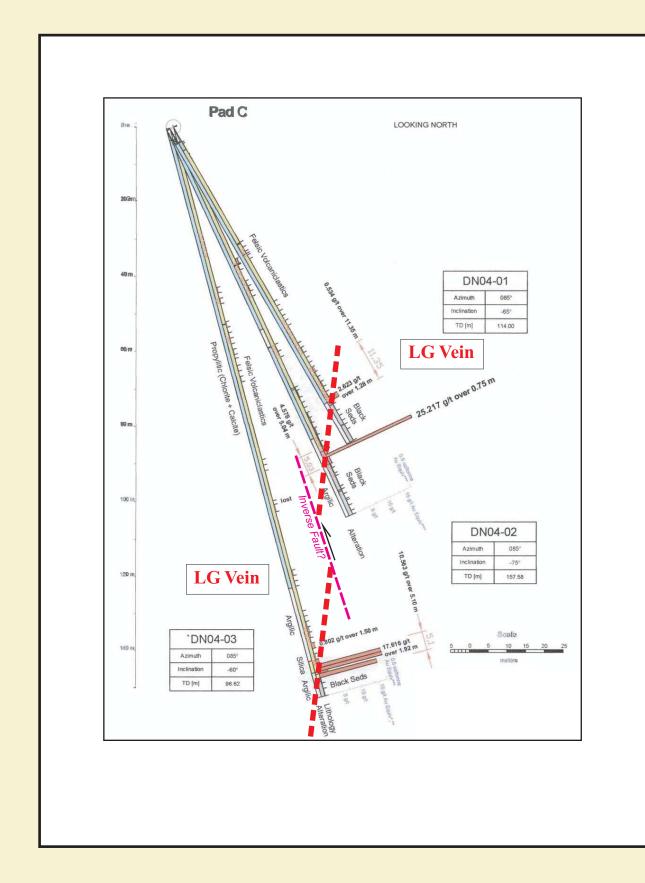
In 2005 Sabina Silver Corporation entered into an option agreement with Teuton and funded three other summer drill programs on the greater Del Norte property. However, the publicly available documentation from these three programs is very limited. The following comments concerning these drilling programs have been excerpted from the subsequent assessment report by D. Cremonese, (2015):

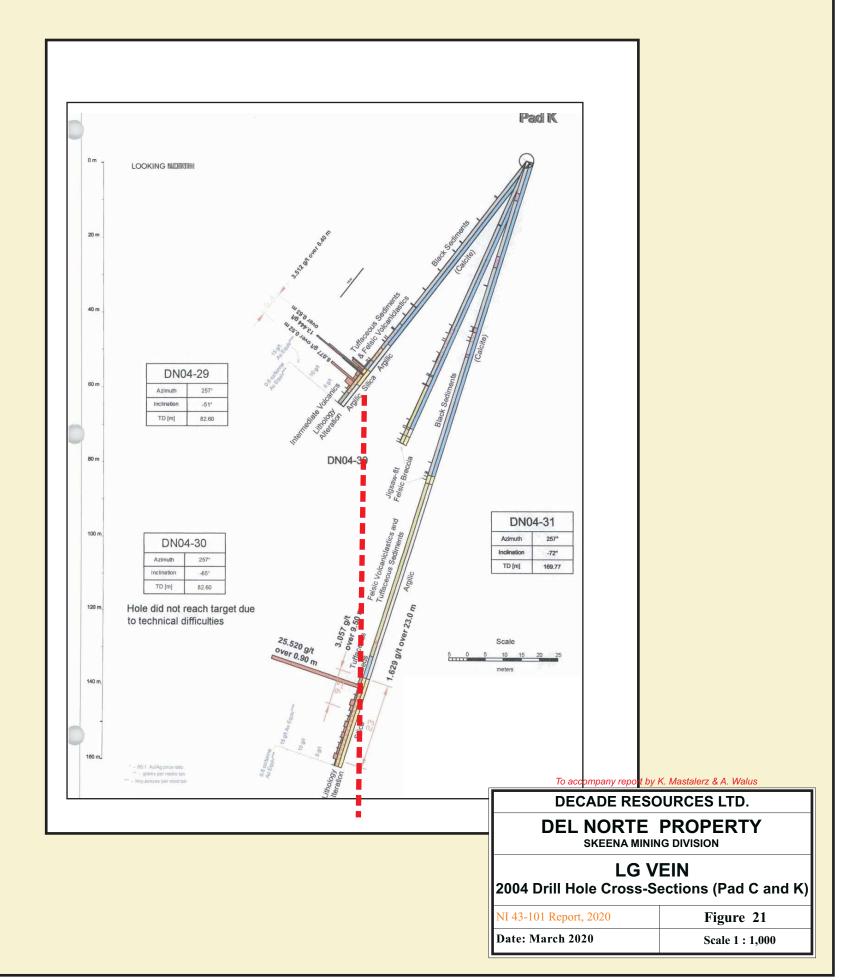
[&]quot;Sabina focused primarily on the LG Vein shear/breccia zone in 2005, drilling ten holes to test the strike and dip of the LG Zone. Late in 2005,



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The map is modified from Mastalerz & Cremonese (2004), Fig. 29





an Aeroquest helicopter borne EM/magnetic survey was flown over most of the property which identified a series of parallel EM conductors situated just east and parallel to the LG vein.

In 2006, Sabina completed an extensive prospecting program covering numerous areas on the Del Norte property as follow up to the airborne EM survey. Grab samples from the "3 Oz" Vein returned encouraging anomalous gold values up to 29.65 g/t Au. Sabina's 2006 drill program at the "3 Oz" Vein gold showing consisted of 3 short holes, totaling 659.59 metres.

In 2007, Sabina followed up on the high grade gold and silver results from the "3 Oz Vein" with a nine hole drill program, totaling 1,600 metres."

It has to be keep in mind that the "3 Oz Vein/Zone" area is located on the mineral claim Midas 3 (Minfile #396295) which does not belong to the Del Norte property optioned by Decade (see Fig. 3).

The 2005 drilling program included 10 diamond drill holes which totaled in 4600 ft (1402.08 m) of core and brought about additional intersections which allowed for better understanding of geometry, continuity and grades of the vein. The following information concerning the results of the 2005 drill program comes from the unpublished excel file made available to authors by Mr. D. Cremonese, CEO of Teuton Resources Inc. for the purpose of completion of this report. The most significant results of this drill program are summarized in the following table (Table 9). The results of this program were also briefly reported in the corresponding issue of News Release (Teuton, 2005).

Table 9. Significant results of the 2005 diamond drill program on the LG Vein area (Cremonese, 2020; unpublished file).

Drill	Inclination	From	То	Interval	Ag g/t	Au g/t
Hole	(deg)	m	m	m		
DN05-01	-45	20.42	28.04	7.62	31.12	0.25
DN05-02	-60	184.10	199.64	15.54	346.02	6.47
including		190.78	191.64	0.86	3267.9	71.5
DN05-03	-47	17.37	19.66	2.29	419.11	2.54
DN05-04	-58	64.57	69.19	4.62	214.07	4.32
DN05-05	-60	180.44	184.71	4.27	476.45	4.48
DN05-06	-50	34.44	37.03	2.59	734.65	5.42
DN05-07	-53	168.30	170.09	1.79	129.12	6.42
DN05-08	-67	159.82	160.74	0.93	133.21	0.72
DN05-09	-45	124.97	127.92	2.95	53.43	0.91
DN05-10	-58	168.65	172.82	4.17	142.12	3.14

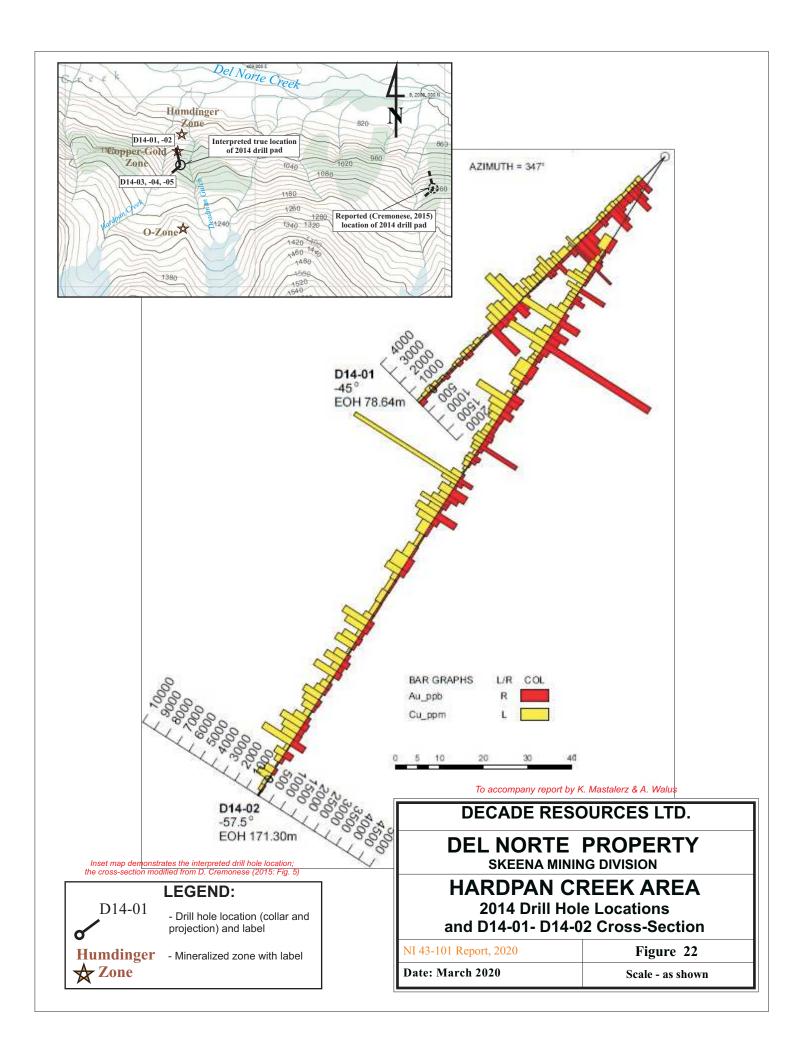
The 2005 drill holes included a few holes collared at the pad N, the southernmost pad from the previous drilling season. The hole DN05-02 collared at this pad intersected a short interval (0.86 m long) which returned 3267.9 g.t silver and 71.5 g/t gold, the intersection labelled the "high-grade zone" (Teuton, 2005). The 2005 program also included two holes drilled from the pad constructed over a nunatak within the middle of the glacier some 270 metres south of the pad N. Both these holes (DN05-07 and -06) drilled at distinct azimuths successfully intersected the LG vein at moderate depths (Table 9).

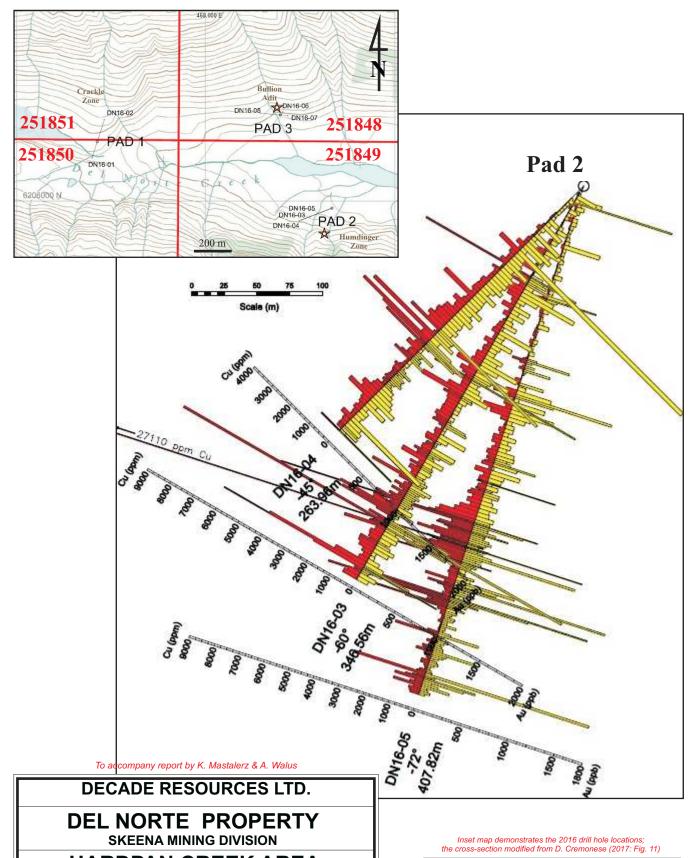
The following, 2006 drill program consisted of a few drill holes which tested the LG vein (4 holes), Kosciuszko Zone (K-zone) – 2 holes and the recently discovered "SP Vein" (Teuton, 2006a) – the southern extension of LG Vein (2 holes). The SP vein was discovered a few hundred metres south of the Kosciuszko Zone beyond the crest of the Kosciuszko cliff (Figs. 4 and 17). Most of these holes successfully intersected the mineralization and the significant results are presented in the Table 10.

Table 10. Significant results of the 2006 drill program on the Del Norte property (according to Teuton News Release of December 18th, 2006.

D.:III.I.a.la	Target	From	То	Interval	Silver	Gold	True width
Drill Hole		m	m	m	g/t	g/t	factor
D06-05	LG Vein	79.7	83.6	3.9	0.5	0.43	unknown
and		194.46	195.95	1.49	57.8	1.35	
and		231.82	233.65	1.83	94.4	0.86	
D06-06	LG Vein	not o	completed	due to tech	nnical prob	lems	
D06-07	LG Vein	39	51.7	12.7	1.5	0.17	unknown
and		199.89	200.25	0.36	372.0	6.81	
and		264.45	267.15	2.7	6.0	0.43	
and		280.95	288.5	7.55	13.7	0.47	
D06-08	LG Vein	133.2	134.85	1.65	1.9	0.52	unknown
D06-09	K-Zone	not completed due to technical problems					
D06-10	K-Zone	175.7	176.7	1	131.0	0.62	unknown
and		200.25 201.65 1.4 30.2 0.8					
D06-11	SP Vein	73.25	77.3	4.05	4.9	0.61	0.53
including		76.3	77.3	1	1208.0	26.54	
D06-12	SP Vein	135.2	143.8	8.6	14.6	0.9	0.24
including		138.83	142.8	3.97	20.7	1.35	
D06-13 HG Zone		abandoned					
D06-14 HG Zone		no significant values					
D06-15 HG Zone		abandoned					

Comment: the above results are quoted strictly according to source; however, some of the "included" intervals do not correspond mathematically to the ones quoted for the corresponding entire intervals.





HARDPAN CREEK AREA

2016 Drill Hole Locations and DN16-03, DN16-04, DN16-05 Cross-Section

NI 43-101 Report, 2020	Figure 23
Date: March 2020	Scale - as shown

DN16-04 Humdinger Zone

- Drill hole location (collar and projection) and label

projection, and laze.

LEGEND:

Humdinger - Mineralized zone with label

Both drill holes which tested the southern extension of the LG vein (SP vein) encountered it at significant depths and of consistently similar width (1.1-1.2 metre true width) as it was encountered and measured at surface (Teuton, 2005). Three additional holes which attempted to test the "high grade zone" discovered the previous year, were abandoned before reaching target depth due to technical problems related to glacier-borne drilling pad stability.

Unfortunately, the pad coordinates for all the drill holes completed in 2006 were never published and they are not know to the authors, similarly as the holes' azimuths and inclinations. The authors were able to verify only the common pad location of the holes D06-11 and D06-12 which tested the SP vein, during their reconnaissance visit to the property in 2019 (see chapter "Exploration" and Figs. 14 and 15).

In 2014, Teuton purchased back Sabina's 50% interest in the Del Norte property and organized another exploration program which included 5 diamond drill holes which yielded 591.93 metres of core and provided 371 core samples for assaying (Cremonese, 2014). All drill holes were collared from a single drill pad and were designed to test a broad zone of copper mineralization on the south side of Del Norte Creek (Fig. 22). However, the true location of the drill pad cannot be verified/established at the present time since its description is not precise. D. Cremonese (2014) wrote in his assessment report concerning the 2014 drilling program as follow:

"The 2014 program was aimed at testing a new zone of copper mineralization which had been recently exposed on the south side of Del Norte Creek just west of Deadman's Gulch. The drill pad was built approximately 150 metres southeast of a prominent showing where blue copper oxides are observed precipitating from ground water at surface." (op. cit.; Cremonese, 2014; pg. 11).

However, the reader is finding the contradictory information in the other part of the same report where the UTM coordinates of the drill pad are listed in a table which summarizes the drill-hole information as: 469970 East and 6205502 North (Cremonese, 2014: page 11). The quoted projection has to be located approximately 1.5 km eastward of the Deadman's Gulch (Fig. 22). The authors of this report had no opportunity to verify the location of the quoted drill holes during their 2019" visit to the property. As for this stage it is assumed the drill holes of the 2014 program were testing the historic "Copper-Gold Zone" identified already earlier on the west-side of the Deadman's Gulch (Fig. 16 and 22, see also Bishop and Gal, 1991).

The significant laboratory results from the 2014 drill program are presented in the table below.

Table 11. Significant results of the 2014 drill program (according to Cremonese, 2015)

Drill Hole	Interval (metres)	Length (metres)	Length (feet)	Gold (g/t)	Copper (%)	Silver (g/t)
D14-01	7.32 to 78.64	71.3	234	0.32	0.06	4.03
incl.	7.32 to 55.78	48.5	100	0.42	0.07	5.2

D14-02	21.95 to 162.46	140.5	461	0.26	0.09	1.57
incl.	43.59 to 81.99	38.4	126	0.49	0.1	2.78
D14-03	63.70 to 75.90	12.2	40	0.92	0.16	3.07
D14-04	19.51 to 110.03	90.5	297	0.15	0.08	0.74
incl.	19.51 to 33.83	14.3	47	0.36	0.17	1.12
D14-05	17.68 to 125.58	107.9	354	0.44	0.07	1.96
incl.	17.68 to 60.05	42.4	139	0.71	0.06	2.8

All drill holes were collared from the outcrop of strongly fractured, sericite-chlorite altered intermediate intrusive rock. Drill holes D14-01 and -02 were drilled northwestward at dips of -45 and -57.5 degrees respectively. The holes intersected significant intervals of moderate quartz-carbonate stockwork with fine grained chalcopyrite and additionally (hole D14-02) the sericite altered intermediate volcanics with some fault gauging and discontinuous quartz-carbonate stockwork.

The other three holes were drilled southwestward at dips of -45, -60 and -70 degrees to test the southern extension of the copper zone. The holes intersected a section of chalcopyrite enriched intermediate intrusives and then entered, and were terminated within a long section of sericite altered intermediate volcanics. It appears that the holes D14-01, -02, -03 and -05 were all stopped, unfortunately, still in a low-grade gold-copper mineralization (Fig. 22; see also: Cremonese, 2014; Figs. 4 and 5). It was also reported that the intersected intrusive rocks contained minor fine grained sphalerite and trace of fine grained disseminated galena.

Another drill program has been completed on the Del Norte property in 2016 when Teuton drill tested four distinct target areas from four drill pads. Thirteen drill holes were planned to be drilled, out of which 12 were successfully completed (Cremonese, 2017). The drill holes yielded 1822.45 metres of core and 711 core samples (including standards and blanks) were assayed. The following table summarizes the most significant laboratory results of the 2016 program.

Table 12. Significant results of the 2016 diamond drill program on the Del Norte property (according to Cremonese, 2017).

		Interval	Gold	Copper		
Drill Hole	from	to length		Gold	Ооррсі	
	m	m	m	g/t	%	
D16-03	160.29	172.24	11.95	0.23	0.19	
D16-04	19.51	110.03	90.52	0.15	0.08	
including	19.51	33.83	14.32	0.36	0.17	
D16-05	264.6	338.8	74.2	0.27	0.15	
including	284.9	296.6	11.7	0.37	0.31	

Drill holes D16-01 and D16-02 were drilled from one pad (Pad 1) toward SSW and NNE, respectively, to test the zone known from occurrence of numerous thin sulfide-bearing quartz veins and stringers south of the Crackle Zone (Fig. 23). Neither of these holes intersected significant mineralization (Cremonese, 2017). Negative results might resulted from the

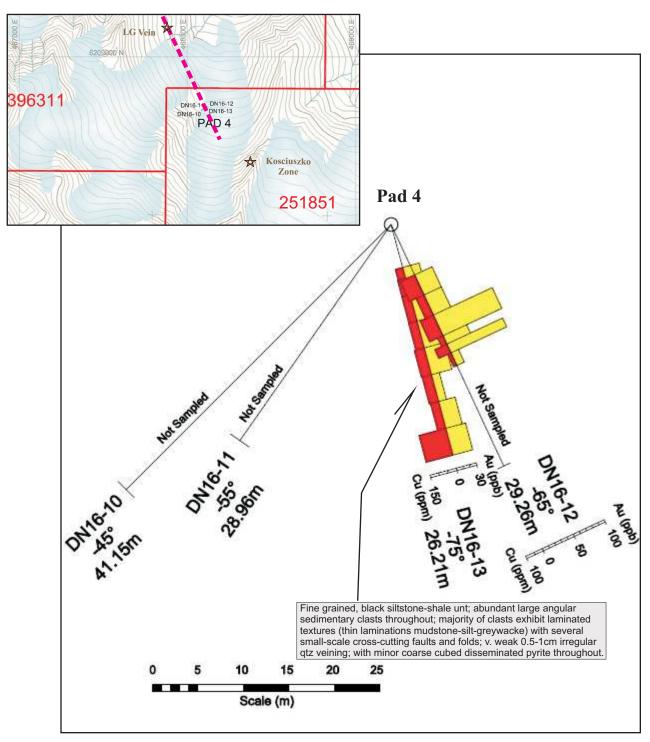
inadequate azimuths of both holes since the Crackle Zone stringers are known to strike approximately N-S and dip moderately to the west (see chapter "Mineralization").

Three holes were drilled from Pad 2 (Fig. 23) northward and downslope from the Humdinger and Copper-Gold zones to test low-grade Cu-Au previously documented from surface work (Bishop and Gal (1991) and intersected by five drill holes in 2014 (Cremonese, 2015). All the holes intersected significant low-grade, fracture-controlled mineralization (Table 12). Long intervals intersected by these holes were characterized by significant K-spar alteration and, locally, sericitization. All three holes encountered a contact between intermediate volcanics and the intrusive rocks at various depths (Fig. 10). It appears that the zones of stronger mineralization are crossing this intrusive boundary and they occur in both, the volcanogenic and intrusive rocks.

Three other drill holes were completed from Pad 3 just below the old adit on the north side of the Del Norte valley (Fig. 23) with an intention to intersect mineralization known from the Bullion Zone and/or NMG vein (Cremonese, 2017). Neither of these holes intersected significant mineralization.

The fourth drill pad was designed to test the LG vein in a spot previously not tested (Fig. 24). The first two short holes drilled westward, toward the vein, did not encounter the target and did not intersect any other visible mineralization. The other two holes were drilled eastward, presumably up-the stratigraphic succession, toward the black sediments (Salmon River Formation?). Both holes intersected short intervals of slightly anomalous gold and copper concentrations host within the sediments (Fig. 24). The last two intersections appear to be of importance in spite of a rather small scale of geochemical anomaly. The intersected rocks were described as "fine-grained, black siltstone-shale unit; abundant large angular sedimentary clasts throughout; majority of clasts exhibit laminated textures (thin laminations mudstone-siltgreywacke)" (op. cit. Cremonese, 2017: Appendix III – Drill Logs). The described rock unit may represent a distal facies of a hydrothermal spring system (elevated gold and copper) which was brecciated (due to syn- and/or partly early post-hydrothermal faulting) and partly re-mobilized, and redeposited after a very short transport. The mentioned above two intervals are also characterized by quite consistently elevated concentrations of silver (up to 1.2-3.0 ppm Ag) and locally zinc (up to over 2081 ppm Zn), the observation that was not, however, reported in the corresponding assessment report (Cremonese, 2017). The observed bedding/lamination cut the core axis at moderate angle to almost perpendicularly, that may indicate the intersected sequence is dipping steeply, potentially sub-vertically.

During the summer of 2019, Teuton completed two longer drill holes in the LG area (Cremonese, 2019). The 2019 drill program was aimed at testing an airborne EM anomaly discovered by a 2005 VTEM airborne survey (Cremonese, 2006) and confirmed, and elaborated by a subsequent ZTEM survey in 2018 (Geotech, 2019). The drill pad was prepared significantly downhill (and northward) from the Kosciuszko Zone and about 200 metres downhill from the LG vein exposure (Fig. 25). The holes were drilled at azimuths of 150 and 130 degrees at moderately to high inclination and reached depths of 342.38 and 349.09 metres. The slightly, although apparently significantly elevated concentrations of zinc and silver have been encountered in both drill holes, in their shallowest intersections (Table 13).



To accompany report by K. Mastalerz & A. Walus

Inset map demonstrates the interpreted drill hole location; the cross-section modified from D. Cremonese (2017: Fig. 14)



LEGEND:

- Drill hole location (collar and projection) and label

DECADE RESOURCES LTD.

DEL NORTE PROPERTY

SKEENA MINING DIVISION

LG Vein Area

2016 Drill Hole Locations and DN16-10 through DN16-13 Cross-Section

NI 43-101 Report, 2020	Figure 24
Date: March 2020	Scale - as shown

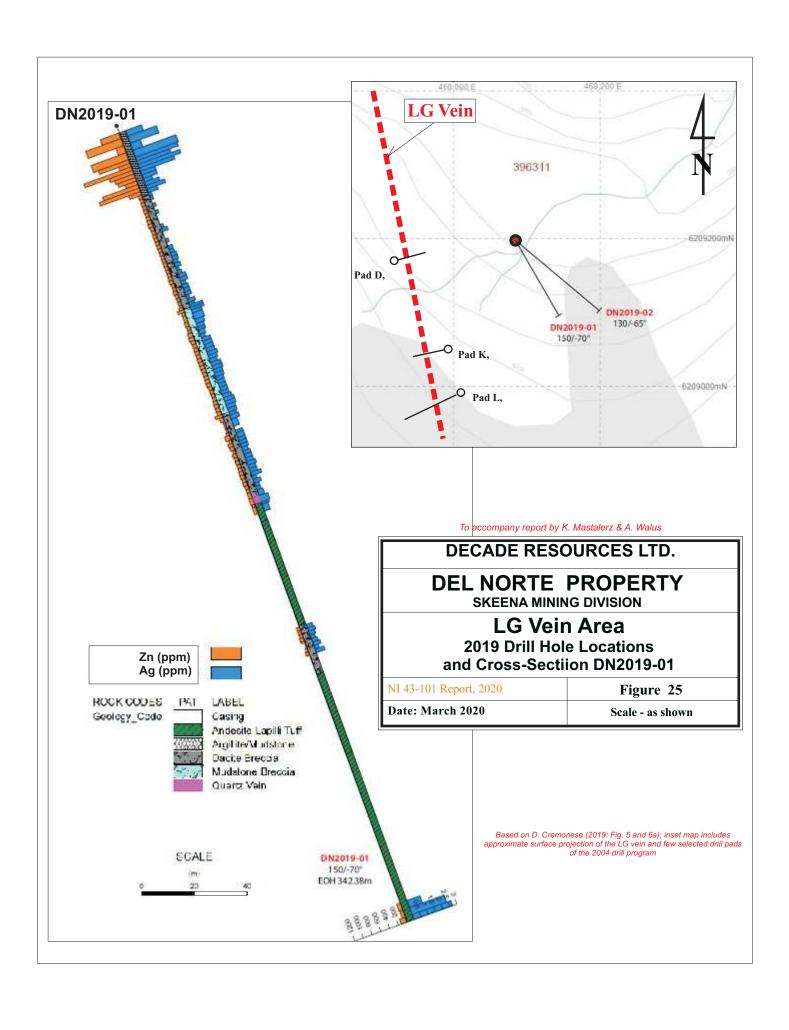


Table 13. Significant intersections of the drill holes DN-2019-1 and DN-2019-2

		Interval		Zinc	Silver
Drill Hole	from	to	length	ZIIIC	Silvei
	m	m	m	ppm	ppm
DN-2019-1	2.44	29.27	26.83	558	1. 68
DN-2019-2	1.52	28.20	26.68	746	1.70

The structural arrangement of the main lithological units intersected in these two holes differs significantly from what supposed to be expected with reference to the classic LG vein area (see chapter on Local Geology). Both holes intersected fine-grained sediments (Salmon River Formation?) at shallow depths (Fig. 10) and then they encountered dacite breccia, mudstone breccia and intermediate volcanics progressively deeper. Such a succession of lithologies appears to imply that the holes intersected successively Mt. Dilworth and then enter into the Betty Creek Formation eastward, hence just an opposite order of lithostratigraphic units as is observed near the LG vein. Either a narrow, tight syncline and/or tectonic faulting had to be held responsible for such a structural inversion.

Independently, both intersections of black argillites with anomalous concentrations of silver and zinc in such a stratigraphic and structural context may be reasonably regarded as a distal portion of the fossil hydrothermal vent complex. It is worth noting that mineralization similar in geochemical character (very strong Ag-Zn anomalies) has been discovered and documented nearby at the BA prospect and Ataman showing in analogous stratigraphic context (Walus 2006; Kruchkowski 2008).

10.2 GEOPHYSICS

In summer of 1989, Prime Exploration funded a helicopterborne, 292 line-kilometer combined magnetic, electromagnetic and VLF-EM survey (flown by Aerodat Ltd. of Mississauga, Ontario) over the both slopes of the Del Northe Creek valley down south to the Willoughby Creek of the Del Norte property. The flight lines were oriented E-W and nominally spaced 100 metres apart. Independently on a very common in all rugged terrains problems related to flight height corrections the Aerodat survey was quite fruitful and cost effective (Mallo and Dvorak, 1989).

The total magnetic field intensity displayed a significant variability over the surveyed area. It appears that the strongly elevated magnetic intensity occurs generally over the areas where bedrock consists of volcanogenic rocks; the lower intensity is mostly restricted to sedimentary bedrock. However, several smaller-size, very strongly magnetic domains have been encountered and they seem to correspond to intrusive bodies. Some domains of the locally depressed magnetic field appear to represent areas characterized by strong alteration. One of such areas can be identified over a significant portion of the Hardpan Creek area, as well as the Crackle Zone (see: Mallo and Dvorak, 1989: Map No: 4). Predominant elongation of variable magnetic intensity domains is from NW to SE and had been interpreted to define lithological variability.

The vertical magnetic derivative (vertical gradient) depicts several prominent linear features which had been interpreted as structurally controlled features such as faults (Mallo and Dvorak, 1989). Their preferred orientation is northwestern but they also display numerous SW-NE offsets. Several of these linear features coincide with the marginal parts of the highly magnetic domains.

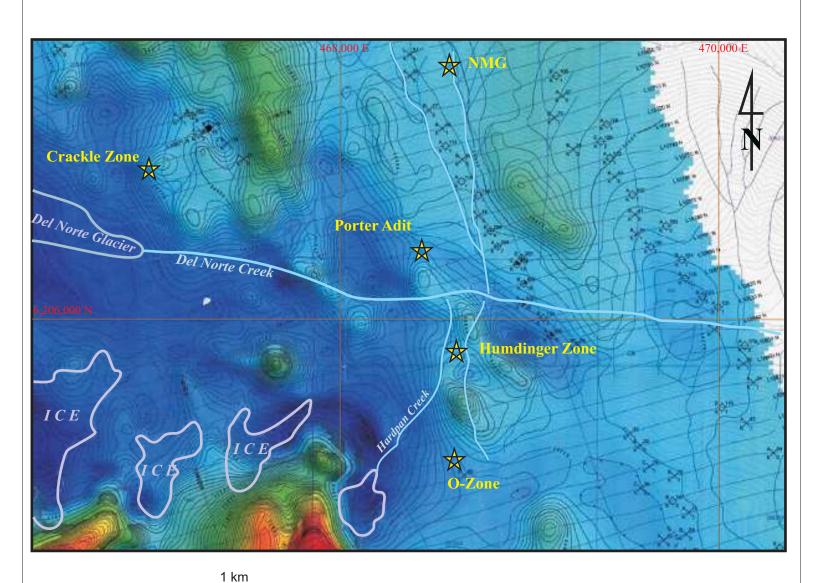
The 1989 electromagnetic survey resulted in a conclusion about generally highly resistive "geological environment" of the area surveyed (Mallo and Dvorak, 1989). However, such type of geologic background usually facilitates delineation of the potential conductor zones/features. Several of linear EM features have been identified within the limits of the survey area and all interpreted conductive zones were regarded to be associated with the flanks of the magnetic anomalies.

Another set of three independent geophysical surveys (VLF-EM, magnetic and UTEM) was completed on the Del Norte property by Goodgold in 1990. The magnetic and VLF surveys were conducted over the entire area of a metric grid established over the Hardpan Creek area of the property (Bishop and Gal, 1991). The UTEM survey was focused on the specific showings encountered at the Hardpan Creek area to test the ability of the method to delineate those types of mineralization. The early results were however not encouraging and the UTEM survey was soon terminated.

The VLF-EM survey outlined numerous weak to moderate conductors which mostly coincide with inferred NW-trending faults in the Hardpan Creek area. Some of these faults were penetrated during the subsequent diamond drilling program. Several weak VLF-EM anomalies were interpreted to represent weak conductors of fault and/or shear zones, or following contacts which may be weakly mineralized (Bishop and Gal, 1991). Results of the magnetometer survey appeared to follow quite closely the pattern of the bedrock lithology. Generally, the volcanic rocks displayed a greater intensity of the total magnetic field than the sedimentary rocks. The intense magnetic highs were interpreted to correspond to the plugs of intrusive rocks, especially their mafic varieties.

In 2005, Teuton funded an extensive 1299.5 line kilometers survey over the combined adjoining areas of the Del Norte, Midas, and Konkin Silver properties. The helicopterborne AeroTEM survey was completed in late summer/autumn by Aeroquest Ltd. of Milton, Ontario, and it covered approximately 65-70% of the area of the present day Del Norte property as optioned to Decade Resources (Cremonese, 2006). The lines were oriented WSW – ENE and were nominally spaced by 100 metres.

The magnetic data gathered during the 2005 survey allowed delineating several domains which differ from each other with respect to intensity of the total magnetic field (Cremonese, 2006; Appendix III; see also Fig. 26). The tilt derivative of the TMI (total magnetic intensity) delineated several NNW-SSE to N-S elongated linear features along the eastern boundary of the coverage area, which are believed to follow the most important lithostratigraphic and tectonic boundaries



Fragment of the TMI map excerpted from
D. Cremonese (2006; Appendix III - Total Magnetic Intensity, Plate 2, Del Norte Block"
by AeroQuest Limited; modified

Relative color scale:

- deep blue low magnetic intensity
- green-to-yellow moderate magnetic intensity orange-tored high magnetic intensity

To accompany report by K. Mastalerz & A. Walus

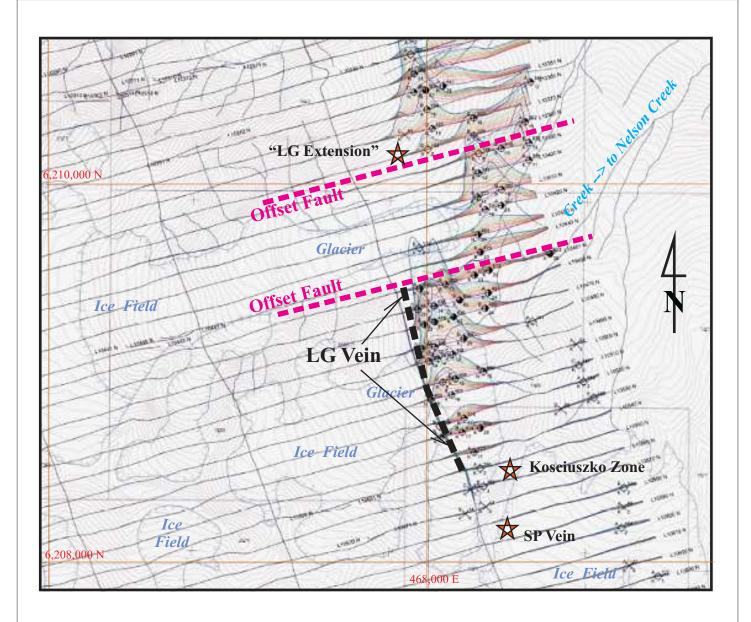
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DEL NORTE PROPERTY

SKEENA MINING DIVISION

Del Norte 2005 Airborne Survey Total Magnetic Field (TMI), Del Norte Valley

NI 43-101 Report, 2020	Figure 26
Date: March 2020	Scale 1 : 20,000



1 km

Fragment of the electromagnetic map excerpted from D. Cremonese (2006; Appendix III - Aerotem Off-Time Profiles, Plate 1, Del Norte Block" by AeroQuest Limited; modified

Legend:



- interpreted surface trace of the LG vein
- interpreted offset faults

To accompany report by K. Mastalerz & A. Walus

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Del Norte 2005 Airborne Survey AeroTEM Off-Time Profiles, LG Vein - Kosciuszko Zone Area

NI 43-101 Report, 2020	Figure 27
Date: March 2020	Scale 1: 20,000

(*ibidem:* Plate 1 and 2). It appears that such a trend is following the broad LG vein zone as delineated from the previous exploration and drill programs (Teuton, 2004, Mastalerz and Cremonese, 2004). One of the significant disturbances in the tilt derivative of TMI of this trend covers the area between the core area of the historic LG showing and its interpreted northern extension ("LG Extension" – compare Figs. 3, 8, 9, 11 and 19), where drilling failed to intersect the proper and/or complete LG vein (Mastalerz and Cremonese, 2004). Distribution of the total magnetic field intensity, as well as its tilt derivative, in the central and western parts of the Del Norte property display much more diversified patterns that appear to mimic a complex distribution of the lithologic varieties of the volcanogenic succession of the Hazelton Group overprinted by their variable alteration.

The data gathered by Aeorquest also displayed a strong, N-S to NNW-SSE trending electromagnetic anomalies along the eastern edge of the survey area (component of the 2005 survey (Cremonese, 2006; Appendix III — Plates 1 and 2). The most important anomaly expressed in stack of derivatives (gradients) of the apparent conductivity of the off-time channels quite closely coincides with the contact between the sedimentary package of the Salmon River Formation and the older volcanogenic rocks of the Hazelton group (*compare chapter*. "Local Geology"), and with the surface exposure of the LG vein. Apparent absence of the LG vein between the drillpad Esh and the "LG Extension" area (Fig. 3, 8, 9, 11 and 18) as documented by drill-hole intersections (Mastalerz and Cremonese, 2004) appears to be caused by a significant tectonic offset of the vein and the Salmon River/Mt. Dilworth contact northeastward (Fig. 27). It has been interpreted that the majority of the electromagnetic sources (anomalies) are steeply dipping eastward in this area and are related to conductive units within the sedimentary package (Cremonese, 2006; Appendix III).

The same EM anomaly continues further south-southeastward, toward the Kosciuszko Zone/SP vein area where it fades (probably due to the rising elevation) and then re-appears toward the NMG-Bullion showings on the northern slopes of the Del Norte valley. The same linear trend fades again while crossing the elevated topography of the ridge between the Del Norte and Willoughby valleys, and re-appears once again at the southern boundary of the property and further south where it coincides the 3 Ounce showing (compare Fig. 3; also: Cremonese, 2006; Appendix III – Plates 1 and 2 and Dickenson, 2008).

In 2016 Simcoe Geoscience conducted a limited magnetotelluric ground geophysical survey in the Del Norte Valley over the low-elevation area between the Hardpan Creek and Crackle zones (Cremonese, 2018). The results were interpreted in terms of the strong K-spar alteration previously encountered in the Hardpan Creek and Crackle zones and its position within a potential porphyry copper type of mineralization. The magnetotelluric stations were designed to aid in interpretation for some historic drill-hole intersections/results which encountered low-grade copper-gold mineralization in the survey area (Bishop and Gal, 1991; Cremonese, 2015, 2017; Figs. 10, 22 and 23). The relatively low and high resistivity domains delineated in the result of the survey were interpreted as predominantly sedimentary and volcanogenic packages of rocks, respectively. The higher-resistivity package, interpreted as a K-spar altered volcanic formations were indicated as a potential host of mineralization (Cremonese, 2018: Appendix III).

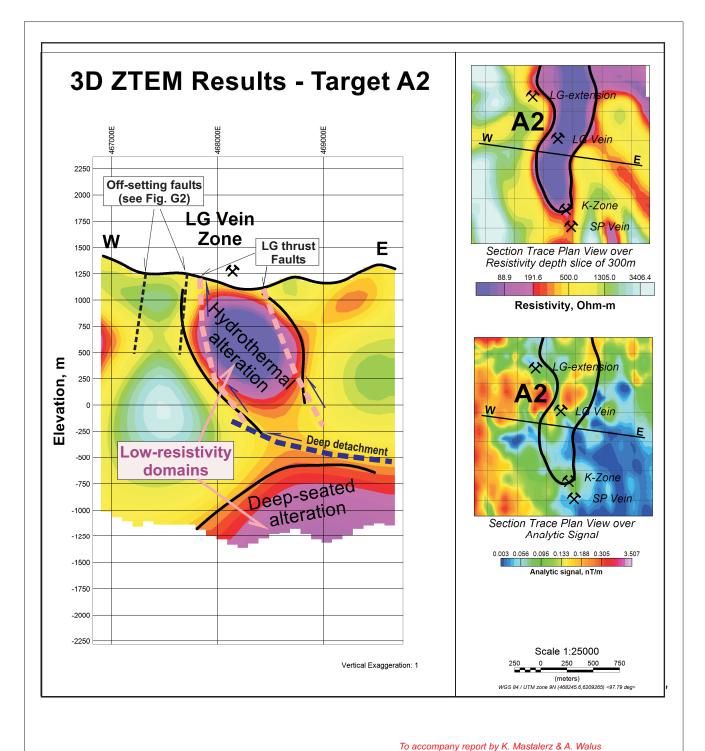


Figure excerpted from Geotech, 2019: Fig 4-2. ZTEM inversion results for Target A2. (Modified)

DECADE RESOURCES LTD. DEL NORTE PROPERTY SKEENA MINING DIVISION Del Norte 2019 Airborne Survey ZTEM Inversion Results by Geotech LG Vein Area

NI 43-101 Report, 2020 Figure 28
Date: March 2020 Scale as shown

The interpreted contact between sedimentary and volcanogenic rocks appears to be subvertical and runs across the Del Norte valley.

In summer of 2018, Geotech Limited carried out a helicopterborne ZTEM survey over the Lord Nelson, Del Norte and Midas properties. The survey included gathering of digital magnetic and electromagnetic data, data inversion, mineralization targeting and 3-D modelling (Geotech, 2019). The magnetic data provided outlines of structural features which differ with respect to magnetic intensity and new insights into a structural interpretation of the Del Norte property. The data analysis allowed distinguishing of a prominent north-south striking feature in the eastern part of the surveyed block, which is interpreted as a significant structural and lithostratigraphic boundary and was suggested to have played a key role in distribution and control of the mesothermal mineralization of the LG vein and related showings. The results also suggested the presence of two main fault systems, NW-SE and SW-NE striking, within the volcanic-rock-dominated central and eastern portions of the surveyed block.

The 3-D inversion of electromagnetic data provided a three-dimensional resistivity model of the subsurface to a depth of approximately 2000 metres. The interpretation of the inversion results allowed identification of numerous conductive and resistive (in relative terms) domains in the subsurface of the surveyed block. The best delineated conductors have been identified along the eastern boundary of the surveyed area (Geotech, 2019). They usually exhibit linear to curvilinear shapes, strike south-to-north and dip very steeply. The authors of the interpretation (Geotech, 2019) found out that most mineral occurrences of the Del Norte property occur coincidentally within those conductor domains which are additionally characterized by low magnetic intensity. It has been suggested that such indicated conductors are associated with graphite-bearing sediments and/or large-scale hydrothermal alteration zones (high sulphidation and argillitization).

The advanced Geotech's interpretation also resulted in identification of eleven potential targets favourable for hosting mesothermal precious metal mineralization and a few targets which displayed some semblance to low-grade porphyry/stockwork model (e.g. Fig. 28).

11. SAMPLE PREPARATION, ANALYSES AND SECURITY

11.1 SAMPLE PREPARATION

All the core obtained from drilling was placed in wooden boxes by the driller helper along with a small wooden block placed at the end of every 10 ft drill run (3.048 m) to mark the depth of the hole. Once full, boxes were covered with a wooden lid and secured for transportation. Next, the core boxes were slung by helicopter for further processing. In 2002, 2014 and 2016 the core was transported to the staging area by Surprise Creek bridge, from there the core was transported by truck to Teuton warehouse in Stewart. During drilling campaigns conducted from 2003 to 2005 the core was transported directly to core shack located on the property.

Upon delivery to the core shack or warehouse the boxes and labelling blocks were inspected for errors, and once it was assured to be correct the footage written on wooden blocks were converted to metres and the ends of the boxes marked with the corresponding metres. Data recorded on drill core include recovery and RQD, geological description and sample intervals. Sample coverage was designed to cover all quartz veins and stockwork as well as adjacent pervasive alteration.

Geological description included rock type, alteration, structures, mineralization and any other feature the geologist considered relevant. Sample breaks were inserted by the geologist, which reflected changes in the rock. Sample intervals ranged from 0.2 to 3.0 metres. For mineralized core the intervals were usually 1.5 m long. For weakly mineralized and/or altered core the intervals were usually 3.0 metres long. After logging, the core samples were sawn in half with a gas powered, diamond saw. Half of the core was placed in numbered plastic bags with corresponding paper tags inside. A sample tag with corresponding number was attached into the core box at the end of every sample interval. Plastic bags were then placed in sealed rice bags and shipped to the laboratory. The other half of the core was placed in racks for reference purposes. Analytical blanks and standards were periodically placed within the core as part of the quality control measures.

11.2 ANALYSES

Companies which conducted exploration work on the property used the following analytical laboratories: Eco Tech Lab of Kamloops, Pioneer Lab of Richmond, ALS Chemex of North Vancouver and Loring Laboratories of Calgary. The authors are not aware if personnel of these laboratories had any relationship to the issuer. These laboratories used standard, proven methods of assaying. ALS Chemex and Loring Laboratories have ISO 9001 certification. Eco-Tech lab was sold (currently it is a part of ALS Chemex) and the authors were not able to obtain information about its certification. A brief description of methods used by these laboratories is presented below.

Eco-Tech lab (1993-94)

All samples were analyzed for 32 elements by Inductively Coupled Argon Plasma of 0.5 g subsample from each analyzed sample. Before analyzing, each 0.5 g sub-sample was digested by 3 ml of Aqua Regia at 95 degrees for one hour followed by dilution to 10 ml by water. The atomic absorption measurements were preceded by subjecting 10 grams samples to standard fire assay preconcentration technique to produce silver bids which were subsequently dissolved. Specific samples were subjected to analysis if values obtained exceeded certain threshold level. High golds were fire assayed using conventional methods followed by paring and weighting of beads.

Pioneer Lab (2002-2004)

All samples were analyzed for 32 elements by Inductively Coupled Argon Plasma of 0.5 g subsample from each analyzed sample. Before analyzing, each 0.5 g sub-sample was digested by

3 ml of Aqua Regia at 95 degrees for one hour followed by dilution to 10 ml by water. The atomic absorption measurements were preceded by subjecting 10 grams samples to standard fire assay preconcentration technique to produce silver bids which were subsequently dissolved. All samples were also analyzed for gold using 30 grams sub-samples by fire assay with Atomic Absorption finish.

ALS Chemex (2006, 2007)

All samples were analyzed for Au, Ag, Cu, Pb, Zn, As and Mo by atomic absorption spectroscopy. A prepared sample is fused with a mixture of lead oxide, sodium carbonate, borax, silica and other reagents as required, inquarted with 6 mg of gold-free silver and then cupelled to yield a precious metal bead. The bead is digested in 0.5 mL dilute nitric acid in the microwave oven, 0.5 mL concentrated hydrochloric acid is then added and the bead is further digested in the microwave at a lower power setting. The digested solution is cooled, diluted to a total volume of 4 mL with de-mineralized water, and analyzed by atomic absorption spectroscopy against matrix-matched standards.

Loring Lab of Calgary (2014, 2016)

All samples were analyzed for 32 elements by Inductively Coupled Argon Plasma of 0.5 g subsample from each analyzed sample. Before analyzing, each 0.5 g sub-sample was digested by 3 ml of Aqua Regia at 95 degrees for one hour followed by dilution to 10 ml by water. The atomic absorption measurements were preceded by subjecting 10 grams samples to standard fire assay preconcentration technique to produce silver bids which were subsequently dissolved. All samples were also analyzed for gold using 30 grams sub-samples by fire assay with Atomic Absorption finish.

11.3 SECURITY

The authors supervised samples preparation during the drilling campaigns from 2002 to 2005 and are not aware of any factors that may have jeopardized samples security. During these programs, no aspect of sample preparation was conducted by an employee, officer, director or associate of any company involved in the property. The authors have no knowledge of the security measures utilized by companies which conducted exploration on the property before and after the 2002-2005 period. It is assumed that the sampling procedure applied by these companies was according to standard industry practices.

The authors are satisfied with the adequacy of sample preparation, security and analytical procedures employed during exploration campaigns on the Del Norte property.

12. DATA VERIFICATION

The authors collected many of the samples as well as logged and supervised the splitting of the core in the period from 2002 to 2005. The authors have no knowledge of the nature and extent of the quality control measures employed by companies which conducted exploration work

before and after 2002-2005 period. It is expected that these companies used standard industry practices in collecting and processing their samples.

The authors consider data from the previous exploration campaigns adequate for the purpose of this report.

ITEMS 13 TO 22

These items are not applicable.

23. ADJACENT PROPERTIES

Important prospects and showings located close to the property (see Fig 3) are described here to give an account of exploration activity in the surrounding areas as well as for the reason of better understanding the mineral potential of the property. The showings described below do not reflect in any manner on mineralization on the Del Norte property. The legal status – current ownership of these properties has not been searched and has no bearing on this technical disclosure by Decade Resources.

BA (Barbara Zone), MINFILE No. 104A 180

BA (Barbara) prospect is located 8 km NW from the NW corner of the Del Norte property (Fig 2). The prospect has been described as being part of an exhalative system with associated zinclead-silver mineralization. The main exhalite horizon is up to 50 metres wide and can be traced for at least one kilometre. It is composed of intercalated centimetre-scale laminae of red, grey, black and green chert. Two types of mineralization have been recognized. Volcanogenic massive sulphide (VMS)-style mineralization is confined to sedimentary horizons immediately below the exhalite. The bulk of the VMS mineralization is contained within a felsic volcanic/sedimentary breccia dominated by strongly silicified semi-angular to angular dacitic clasts ranging in size from 0.1 to over 20 centimetres. Clasts of chert, andesite, mudstone, volcanic tuff, exhalite and rarely of sulphides were also noted in these rocks. The highly brecciated nature of the sediments suggests a fairly active environment proximal to a volcanic vent. Epigenetic mineralization hosted within highly fractured and brecciated dacite with quartz + sulphide veining is also common and may represent either a feeder for the VMS mineralization or post-remobilization. Sulphides include up to 5 per cent pyrite, trace to 3 per cent sphalerite, trace to 3 per cent galena; locally there is also trace to minor chalcopyrite, tetrahedrite and silver sulphosalts. One to three millimetre thick laminae of pyrite and sphalerite intercalated with mudstone laminae and fine felsic tuff are common. In 2007, the best single interval assay was 3.05 metres of 1215 grams per tonne silver, 1.01 per cent lead and 2.26 per cent zinc in hole BA-2007-01. Drilling on the Barbara zone in 2010 (hole BA- 2010-147) intersected 15.24 metres grading 117.5 grams per tonne silver, 0.02 per cent copper, 1.18 per cent lead and 2.81 per cent zinc; this includes 9.15 metres of 150.0 grams per tonne silver, 0.03 per cent copper, 1.75 per cent lead and 3.0 per cent zinc (Assessment Report 32877).

BA North zone, has intersected stockwork mineralization within a subvolcanic andesitic intrusion. Drillhole BA-2010-136 intersected 136 grams per tonne silver, 0.63 per cent lead and 0.53 per cent zinc over 3.05 metres; 82 grams per tonne silver, 1.06 per cent lead and 1.20 per cent zinc over 12.19 metres; and 62 grams per tonne silver, 0.54 per cent lead and 1.72 per cent zinc over 3.05 metres (Assessment Report 32877). Assuming that the North zone is part of the Barbara zone, drilling to date has traced the zone for over 1000 metres along strike. The zone is open along strike as drilling has not as yet determined the boundaries of the system. There is no data between the Barbara and North BA zones as this is covered by glacier.

Three-Ounce Vein, Midas; MINFILE 103 P 262

Three-once vein (Fig. 3) was discovered in 2003 when a float sample of white quartz with no associated sulphides assayed 3 ounces per tonne gold. A follow-up trenching work exposed a quartz vein mineralized with tetrahedrite. The 3 Oz vein is hosted in a west dipping shear zone located at the contact between argillite and volcanic rocks (the same contact which hosts LG vein). The shear zone is comprised of tectonically brecciated argillite and abundant intermittent intervals of fault gouge and quartz/fault gouge breccia. Mineralization consists primarily of acicular arsenopyrite in addition to trace amounts of sphalerite and galena (Dickenson, 2008).

In 2006, Silver Grail Resources Limited in partnership with Teuton Resources Corporation targeted the vein with 3 holes. The best result came from hole SDN-06-02 which yielded 2.52 grams per tonne gold over 32.4 metres. Within this intersection, a sub-interval ran 26.77 grams per tonne gold over 0.7 metre (Property File - Silver Grail Resources Ltd., 2006). In 2007, Sabina Gold & Silver Corporation conducted a nine hole, 1600 metre diamond drilling program to assess the strike and dip extent of mineralization intersected by the three holes drilled in 2006. The 2007 drill program results were not as promising as those obtained in 2006. Assay highlights from the 2007 drill program include 6.75 grams per tonne gold over 0.90 metre and 2810 grams per tonne silver over 1.00 metre, both from hole SDN-07-03. All other assay values were lower. No future exploration work is recommended for the 3 Oz showing (Assessment Report 30156).

Willoughby, MINFILE No. 103 P 006

The Willoughby mineral occurrence located 2.5 km west from the SW corner of the Del Norte property is classified as developed prospect. To date, 11 mineralized occurrences have been located on the Willoughby property. All of the zones are hosted by variable, pervasively sericite +/-carbonate +/-chlorite +/-pyrite altered rocks. Mineralization consisting of pyrite and pyrrhotite along with lesser sphalerite, galena and rare visible gold occurs in veins, stockwork and fracture fillings. In addition, pyrite and pyrrhotite occur as semimassive to massive occurrences in lenses and pods. Several of the zones appear to be intrusion related. At the Willoughby nunatak all zones excluding the North zone occur within andesitic tuffs peripheral to the hornblende feldspar porphyry stock.

Mapping and drilling at the Wilby zone have shown a 20 by 60 metre northwest-trending zone to contain semimassive to massive pyrrhotite and pyrite pods within altered andesitic tuffs. Gold values occur within and immediately peripheral to the sulphides. The zone appears to be flat lying and is open along strike to the northwest. The best drill intercept averages 15.6 grams per tonne gold and 12.3 grams per tonne silver over 4.2 metres (Assessment Report 23674). Drilling at the Upper Icefall zone identified an extensive zone of variably altered andesitic tuffs. Mineralization consists of up to 20 per cent pyrite along with lesser sphalerite, galena and arsenopyrite. One hole tested the zone with the best intercept being a one metre sample assaying 17.8 grams per tonne gold and 44.2 grams per tonne silver (Assessment Report 23674).

Horatio, MINFILE No. 104A 207

Horatio zone found by Teuton crew in 2003 is comprised of several small veins and pods with massive sulphide mineralization. Several grab samples collected from the zone returned high gold assays of up to 171.45 g/t gold.

24. OTHER RELEVANT DATA AND INFORMATION

The authors are not aware of any other relevant data or information on the Del Norte property.

25. INTERPRETATION AND CONCLUSIONS

Historic exploration programs completed on the Del Norte property led to discovery of several significant mineral occurrences. The most prominent mineralized zone discovered on the property to date is the LG vein which has been traced by drilling for 1.2 km and vertical extent of 240 metres (Fig. 11). The vein is open in both directions as well as at depth. The true width of the vein ranges from 0.5 to 1.25 m, however in some places mineralization is not restricted to the vein but occurs also in the surrounding rocks forming breccia and stockwork zones, as well as satellite veins. A good example of such place is the area sampled by the authors in 2019 (Fig14) where several grab samples collected from semi-stockwork and satellite veins close to main LG vein returned up to 19.6 g/t gold and 3920 g/t silver. Another example is the Kosciuszko breccia zone located 60-70 metres from presumed location of the LG vein.

Because of the snow and ice coverage the Kosciuszko zone as well as the stratigraphic contact which host the LG vein located directly to the north and south of the zone were not adequately drill tested (Figs. 14 and 19). This portion of the contact zone which is approximately 700 metres long was tested only by six holes drilled from two set-ups. The LG vein likely continue further south across a large snow patch as indicated by the 2019 sample A19-260 collected 350 metres south from the most southerly located drilling pad used to test the LG vein – SP pad (Fig. 14). The sample, taken from suboutcrop of a vein returned 462 ppb gold and 25.8 g/t silver along with highly anomalous lead, zinc and arsenic values. The NMG vein and Bullion zone which are located within the same stratigraphic contact zone which hosts the LG vein may represent a more distant, southern continuation of the LG vein system. These two mineral occurrences were never drill tested.

The presence of chalcedonic quartz, vuggy cavities, realgar, silver sulphosalts and electrum within LG vein and Kosciuszko zone (Walus, 2002, 2003) unequivocally point to a low-

sulphidation epithermal environment. Considering its epithermal origin, the LG vein shows remarkable horizontal and vertical continuity.

Native gold and silver sulfosalts of the LG vein often occur within a soft, black, graphitic matrix which fills interstices between breccia fragments. In small-diameter (BQ size) drill holes, large part of this soft matrix with accompanying gold and silver mineralization was very likely lost in the drilling process. Significantly, lower core recovery was frequently recorded in numerous historic drill intersections which cut across the strongly fractured vein wall rock and/or across the vein itself. The strong negative correlation between the gold/silver grade of the LG vein and core recovery of the corresponding core intervals has been extensively documented by Mastalerz and Cremonese (2004 – Figs. 22 through 28). An example of such negative correlation is presented on Fig. 29.

According to the heli-borne geophysical survey conducted on the property in 2005 by Aeroquest, the area north of the 2004 Esh drillpad (Fig.20) which include the "LG extension" showing was tectonically shifted 500-600 metres east (Fig. 27). Surprisingly, this important fact was never drill-tested or even acknowledged by previous property operators. Tectonic displacement would certainly explain why several 2004 holes which tested the "LG extension" zone failed to intersect the LG vein (Mastalerz and Cremonese, 2004). It also indicates that the "LG extension" zone is not a part of the LG vein. The broad area of potential true LG northern extension (Fig. 27) becomes the primary targets for the next exploration program.

Another important area which warrants further investigation is the area of so-called Hardpan Creek zone which encompasses 8 mineral occurrences. Some of these occurrences display characteristics of porphyry gold-copper mineralization. However, geology of the entire area is complex and requires a detailed geological mapping as well as compilation of all available data in order to plan further exploration. The results of the historic extensive programs of trenching, rock and soil sampling, diamond drilling and diverse geophysical surveys (*compare*: Mallo and Dvorak, 1989; Bishop and Gal, 1991; Gal and Simpson 1991; Cremonese, 2015, 2017; Geotech, 2019) have never been incorporated into a single comprehensive geological model.

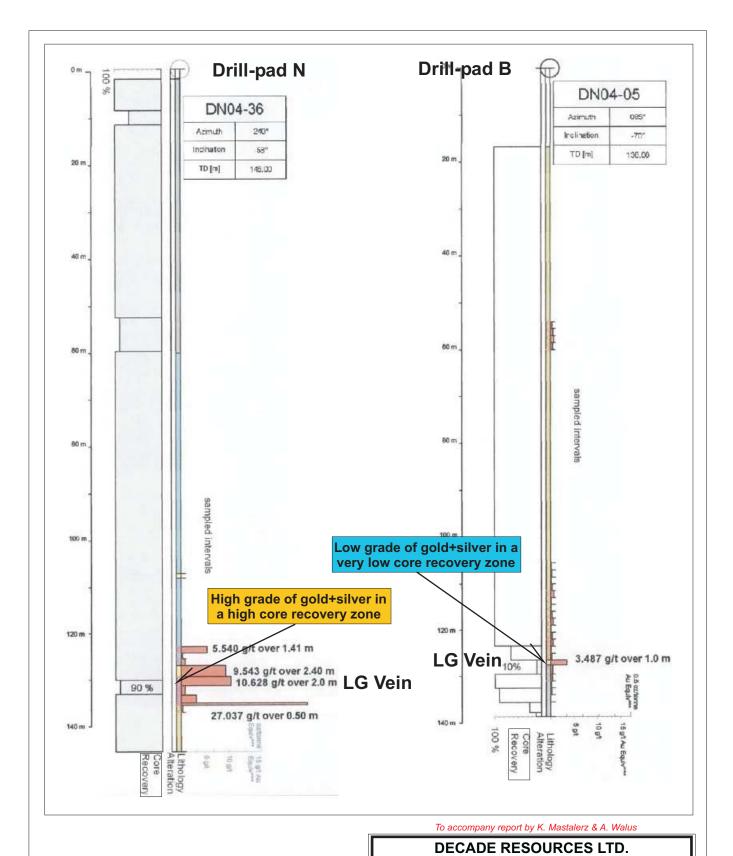
The presence of several large (up to 2.0 metres across) argillite boulders with syngenetic zinc mineralization found in 1994 on Del Norte Glacier (Cremonese, 1995) rises the possibility of finding VMS mineralization similar to one which comprises the nearby BA zone (Fig. 2) which is a major Kuroko style VMS occurrence tested by 178 drillholes (Kushner, 2011).

Large areas of the Del Norte property are weakly explored due to extensive ice and snow coverage. However, the rapidly receding ice enables better access to these areas which may host more mineralized zones.

26 RECOMMENDATIONS

The Del Norte property includes many mineral occurrences, of which several are regarded as especially significant and warrant further exploration. The LG vein-Kosciuszko zone area and the Hardpan Creek group of showings have to be considered the most promising and important.

Several exploration programs completed on the property proved that the LG vein has a more complex structure than originally postulated and requires more exploration work to properly

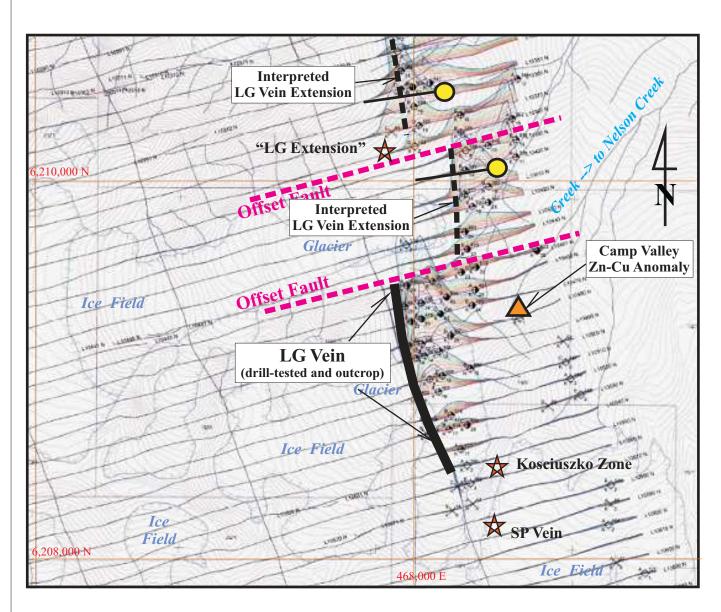


Diagrams adopted and modified from:
Mastalerz And Cremonese (2004 - Figs. 22 and 27).
Gold and silver grades expressed collectively as gold-equivalents according to 2004 relationship between market prices of these metals

DEL NORTE PROPERTY SKEENA MINING DIVISION Core Recovery vs. Gold/Silver Grade 2004 Drill Program - LG Vein NI 43-101 Report, 2020 Figure 29

Scale as shown

Date: March 2020



1 km

Fragment of the electromagnetic map excerpted from D. Cremonese (2006; Appendix III - Aerotem Off-Time Profiles, Plate 1, Del Norte Block" by AeroQuest Limited; modified

Legend:



- LG vein (outcrop and drill tested)
- interpreted northern extension of the LG vein
- interpreted offset faults
- recommended drill pads and drill holes

To accompany report by K. Mastalerz & A. Walus

DECADE RESOURCES LTD.

DEL NORTE PROPERTY

SKEENA MINING DIVISION

Interpreted Northern Extension of LG Vein

(and recommended drill hole locations)

NI 43-101 Report, 2020	Figure 30
Date: March 2020	Scale 1:20,000

evaluate its full potential. Two areas related to the LG vein appear as the most important targets for further investigations. One of them is the potential northern continuation of the vein where it appears to be detached and displaced northeastward by postulated offset faults (Fig. 30). The prior interpretation arises from analysis of the results of the airborne electromagnetic survey completed by Aeroquest in 2005 (Cremonese, 2006; Appendix III). This area requires careful geologic mapping and rock/silt sampling (Fig. 30). However, since a great part of bedrock in this area is covered by loose glacial and fluvial sediments, it appears quite prudent to test the above hypothesis by a dedicated diamond drilling (Fig. 30).

The second area of recommended exploration related to the LG vein is its southern extension (SP vein) and potential further continuation (see chapter "Interpretation and Conclusion"). The "Crest zone" located between the SP glacier and the Del Norte valley (Fig. 14) will require a detailed geological mapping with emphasis on the structural and lithological observations which should shed additional light on the structural and lithostratigraphic context of the vein. The rock formations over this area are strongly weathered and, partly, in sub-crop conditions and will certainly require a reasonable deal of hand trenching and likely, soil sampling (the latter over its southern slope). The "Crest zone" should be also regarded as a preferred location for a drill pad which will allow testing the vein at depth at various azimuths. However, this location will be challenging with respect to the water supply for the drilling operation; the drilling will have to be conducted during the warmest part of summer when the snow/ice-melt water will be available at relatively short distance.

Based on the available data it is assumed that the NMG and Bullion occurrences are very likely, more distant, southern continuations of the LG vein system. These occurrences have never been drill tested so far and two carefully selected drill-pad locations should be designed to serve this purpose. The drilling in this area should be preceded by careful structural mapping and a limited rock sampling. The results of the historic soil and rock sampling (Bishop and Gal, 1991) as well as the results of the airborne geophysical surveys (Mallo and Dvorak, 1989, Cremonese, 2006 and Cremonese, 2018a) should be referred to and be incorporated into the geological interpretation.

It must also be stressed that the results of the authors' reconnaissance rock sampling and geological observations on the SP vein area (LG southern extension) indicate that the LG vein is at least locally accompanied by a much wider zone of a stockwork-type mineralization. This widened zone when approached systematically may bring about a potential economic improvement in evaluation of the vein. A problem of artificially diminished metal grades due to inadequate drill practices can be addressed by drilling a larger-diameter core and using more viscous drilling mad when intersecting the vein or its mineralized wall rocks.

As concluded from the existing data, the Hardpan Creek area is underlain by bedrock formations which are characterized by complex lithostratigraphic and structural relationships, and it demonstrates a reasonable potential for hosting a low-grade, copper-gold, porphyry-type mineralization. However, in spite of the extensive and complex geological documentation the area still requires a complete, comprehensive model which would include all the existing information and provide basis for the further dedicated exploration. Such a cameral, synthesising compilation should be considered a pre-requisite element to any further planned

field exploration. The compilation should be supported by a limited re-logging and re-sampling of the existing core material. The Company should also consider hiring a geophysical consultant who should address the relationship between the physical properties as measured on the core material and the results of the historic geophysical surveys (Mallo and Dvorak, 1989; Cremonese, 2017, Geotech, 2019). It appears that an additional, dedicated geological mapping accompanied by targeted rock sampling, should be regarded as the next exploration step in this area before selecting any new drill targets. The relationship between the intrusive rocks encountered in the area, their host volcanogenic formations and the hydrothermal alteration pattern as expressed in the results of the historic geophysical surveys appears to be of primary significance in selecting further sites for drill testing.

The remaining mineral occurrences of the Del Norte property appear to be of much lesser importance. However, some of them may give reasons for considerations concerning potential for other types of mineralization existing within the limits of the property.

Lithostratigraphy of the rock formations of the Del Norte property, but especially presence of felsic components of the postulated Mt. Dilworth Formation, may allow for hypothesising about a potential for hosting a volcanogenic massive sulphide (VMS) mineralization in this area. Apart of this favourable stratigraphic position however, there is no completely convincing geological documentation supporting such a hypothesis, so far. However, some results of the late drilling programs may provide some preliminary indicators of such an option. Few drill holes completed by Teuton in 2016 and in 2018 in the area eastward from the LG vein returned elevated contents of gold, copper, silver and zinc within the black sediments of the Salmon River Formation. These contents were quoted as "significant" (Cremonese, 2017, 2019). It is hard to definitely ascertain if these, slightly elevated concentrations of precious and base metals represent a geochemical anomaly or they are an inherent, background level characteristic of this formation in the region. The Salmon River Formation was never systematically sampled, so far, being regarded as not having a mineralization potential. The relative proximity to the BA prospect (Fig. 2) where the large-scale VMS-type mineralization is apparently hosted near the Mt. Dilworth/Salmon River boundary (Walus, 2006; Kruchkowski, 2008) may indicate that such elevated concentrations in the Salmon River Formation of the Del Norte property may be just a distant geochemical halo related to the BA prospect.

However, it is worth noting that the results of the last two extensive geophysical surveys completed on the property have indicated some deep "geophysical" targets precisely within the package of the postulated Salmon River Formation and just eastward from the main tectonic/lithostratigraphic contact which is followed by the LG vein (Cremonese, 2006 – Appendix III and Geotech, 2019). It is recommended that the VMS hypothesis is further tested by additional drill holes. Although, it would be prudent that additional, costly drilling is preceded by a more detailed analysis and geophysical modelling of the target. Such modelling should be supported by the direct measurements of physical rock properties as based on the existing core documentation and the bedrock outcrop.

The deeper, older lithostratigraphic elements of the Hazelton Group succession, which are exposed predominantly in the central and western parts of the Del Norte property area, require some more attention than they have received so far. The Hardpan Creek area is a good

example where these older stratigraphic formations host significant mineralization. Similarly, the historic programs of rock sampling brought about several other examples of scattered mineral occurrences in the western parts of the Del Norte property, which are postulated to be underlain by intermediate volcanogenic rocks of the Betty Creek Formation and/or felsic rocks of the deeper horizons of the Mt. Dilworth Formation (e.g. Bishop and Gal, 1991, Cremonese, 1995, Cremonese, 2004, Mastalerz and Cremonese, 2004). However, these small-scale, isolated occurrences were never addressed with sufficient stress to arrive at the model of their genesis and distribution. These occurrences may provide some evidence of the root elements of the mineralization that occurs and/or is expected to appear at shallower stratigraphic levels. It is worth noting that the western parts of the Del Norte property are generally more rugged and in great part covered with ice fields but the progressing glacier ablation provides numerous opportunities to observe several newly exposed areas. Such potential areas should be mapped and rock sampled at the reconnaissance scale.

The 2018 Geotech ZTEM survey revealed also a potential target in the south-eastern part of the property. A prominent, north-south elongated anomaly of low-resistivity crosses the watershed between the Del Norte and Willoughby valleys (Geotech, 2019). This area located just north of the 3 Oz showing (Fig. 3) has been never satisfactorily explored, nor sampled. The anomaly is centered at the "Rhyolite Dome showing" as labelled in 2019 Teuton's news release (Teuton, 2019). It has been ascertained that the all the historic drill holes of the Hardpan Creek area fall outside the northern part of the contour of this anomaly (Teuton, 2019). Interestingly, this ZTEM anomaly appears to coincide with the earlier AeroTEM off-time profiles anomaly (Cremonese, 2006: Appendix III, Plate 2), although the later one fades significantly while crossing the watershed ridge. The "Rhyolite Dome area" requires careful geological mapping and rock sampling, most likely also soil sampling.

It is also recommended to follow up the strong Zn-Cu geochemical anomaly discovered in 2017 while conducting a limited contour soil/talus sampling east of the LG vein (Cremonese, 2018b). The anomaly occurs in the "Camp Valley" (a site of 2004-2006 summer exploration camps) near the base of a very high lateral moraine (Fig. 30). The anomaly coincides with a part of a very strong geophysical apparent conductivity anomaly (Cremonese, 2006; Geotech, 2019) and its nature is unknown, so far.

Lithostratigraphy of the Del Norte property also requires to be seen with much better resolution than it is resolved so far. It may have a critical impact on the evaluation of the property's potential for hosting various types of mineralization and defining future exploration strategy. A reasonable number of carefully selected rock and core samples should be designed for the whole rock analysis. The whole-rock geochemistry will allow for distinguishing between different packages of volcanogenic rocks and ascertain or negate presence of the felsic end members of the Mt. Dilworth Formation. The whole rock geochemistry should be accompanied by the corresponding microscope petrography which together would allow defining the original end-members of the rock formation but also elucidate and aid in defining an impact of post-accumulation processes of the hydrothermal alteration and regional metamorphism and indicate potential sub-surface mineralization packages.

An overall geochemical imprint of the Salmon River Formation also requires detailed addressing. The formation was postulated to carry localized significant geochemistry anomalies (Cremonese, 2017 and 2019) which may indicate the presence of the VMS-type mineralization nearby (sub-surface). These anomalies appear to coincide with some airborne apparent conductivity anomalies. However, the magnitude of these geochemical anomalies (Cu, Zn, Ag) are rather low and the question arises if they fall within the limits of background values for the Salmon River Formation in the area. However, in spite of numerous available outcrops and several drill holes which intersected the rocks of this formation, their geochemistry was never tested systematically. The problem can be resolved in great part by a dedicated rock and core sampling and conducting on them a standard ICP analysis. The problem would be addressed much better if complemented by a limited number of thin-section samples for microscope petrography which should allow defining the petrography of the tuffaceous material usually incorporated in the sedimentary rock of the Salmon River Formation.

The following table summarizes the recommended exploration efforts and provides with a corresponding, generalized financial estimate.

Table 14. Exploration work recommended for the Del Norte property and its estimated costs.

Item	Target area	Specification	Estimated cost*
Project planning, coordination and supervision	Del Norte Project 2020	30 days	24,000.00
Geological mapping	LG north extension	5 field days	3,500.00
Geological mapping	LG - SP - Bullion area	5 field days	3,500.00
Geological mapping	Hardpan Creek	5 field days	3,500.00
Geological mapping	Rhylite Dome	3 field days	2,000.00
Soil/talus Sampling	Camp Valley	2 field days (crew of 2)	1,500.00
Soil Sampling	Rhyolite Dome	4 field days (crew of 2)	3,000.00
Core relogging/target sampling	Historic core	7 days	4,000.00
Diamond drilling	North LG	4 holes - 1000 m	140,000.00
Diamond drilling	SP-Crest-Bullion	4 holes - 1200 m	168,000.00
Diamond drilling	Salmon River - geophysical anomaly	4 holes - 1000 m	140,000.00
Mob/Demob	Drill/Camp/Crew		34,000.00
Geophysical consultsnt		7 days + report	10,000.00
Helicopter support	Field projects	20 days x 1 hour	30,000.00
Helicopter support	Drilling	26 days x 4 hours	140,000.00

Camp costs		28 days (generalized)	100,000.00
Rock sampling	Salmon River, Mt, Dilworth	3 days (geo)	2,000.00
Geochemistry - lab		1000 samples	40,000.00
Compilations/Reporting		12 days (geo + draft)	14,000.00
Microscope petrography		24 samples	7,000.00
Contingency 10%			87,000.00
Total			957,000.00

[•] It is assumed that the above cost estimation will have to be adjusted accordingly to recent and future, global and local market and economical changes

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28. DATE AND SIGNATURE PAGE

This report titled "43-101 F1 Technical Report on Del Norte Property" and dated March 31, 2019 was prepared and signed by the following authors:

Signed "Krzysztof Mastalerz"

Krzysztof Mastalerz, Qualified Person 2005 Bow Drive, Coquitlam, BC

and

Signed "Alojzy A. Walus"

Alojzy A. Walus, Qualified Person 8577 165th Street Surrey, BC

29. CERTIFICATE OF QUALIFIED PERSON

I, Alojzy Aleksander Walus, of 8577-165 Street, Surrey, in the Province of British Columbia, do hereby certify that:

I am a member in good standing of the Association of Professional Engineers and Geoscientists of the Province of British Columbia (License # 24404). I am a graduate (1984) of the University of Wroclaw, Poland and hold M.Sc. Degree in Geology.

I have been practicing my profession continuously since graduation. I have worked in British Columbia from 1988 to 2019 as a geologist with several exploration companies.

This certificate relates to National Instrument 43-101 F-1 Technical Report on the Del Norte property dated March 31, 2019. I am responsible for items 1 to 6, 7.4, 8, 11, 12, 23 and 25 of this report. I worked on the property in 1993, 1994, 2002, 2003 and recently on September 29, 2019.

I am a "qualified person" for the purpose of National Instrument 43-101.

I have read National Instrument 43-101 and the sections of the report for which I am responsible have been prepared in compliance with that instrument.

I am independent of the issuer, Decade Resources Ltd as described in Section 1.5 of the National Instrument 43-101.

As of the effective date of the technical report, to the best of my knowledge, information and belief, the sections of the technical report for which I am responsible contain all scientific and technical information that is required to be disclosed to make those sections of the technical report not misleading.

Submitted this 31 day of March, 2020

Signed "Alojzy A. Walus"

Alojzy A. Walus, Qualified Person 8577 165th Street Surrey, BC I, Krzysztof Mastalerz, of 2005 Bow Drive, Coquitlam, in the Province of British Columbia, do hereby certify that:

I am a member in good standing of the Association of Professional Engineers and Geoscientists of the Province of British Columbia (License # 31243). I am a graduate (1981) of the University of Wroclaw, Poland, and hold M.Sc. Degree in Geology and Ph.D. in Nature Sciences

I have been practicing my profession continuously since graduation as an academic teacher (University of Wrocław and A. Mickiewicz University of Poznań) through 1997, a research associate for the State Geological Survey of Poland (1993-1995), and an independent consulting geologist in Canada, USA and Peru since 1994.

This certificate relates to National Instrument 43-101 F-1 Technical Report on the Del Norte property dated March 31, 2019. I am responsible for items 7.1, 7.2, 7.3, 9, 10, 26, 27, Appendix 1 and all the Figures of this report. I worked on the property in 2004 through 2006 and visited the property on September 29, 2019.

I am a "qualified person" for the purpose of National Instrument 43-101.

I have read National Instrument 43-101 and the sections of the report for which I am responsible have been prepared in compliance with that instrument.

I am independent of the issuer, Decade Resources Ltd as described in Section 1.5 of the National Instrument 43-101.

As of the effective date of the technical report, to the best of my knowledge, information and belief, the sections of the technical report for which I am responsible contain all scientific and technical information that is required to be disclosed to make those sections of the technical report not misleading.

Submitted this 31 day of March, 2020

Signed "Krzysztof Mastalerz"

Krzysztof Mastalerz, Qualified Person 2005 Bow Drive, Coguitlam, BC

APPENDIX 1

Laboratory certificates: analyses of 2019 rock samples collected by the authors on the SP Zone area, Del Norte property

Quality Analysis ...



Innovative Technologies

Report No.: A19-13444
Report Date: 07-Nov-19

Date Submitted: 02-Oct-19

Stewart

Your Reference:

Decade Resources 426 King Street Stewart BC V0T 1W0 Canada

ATTN: Ed Kruchkowski

CERTIFICATE OF ANALYSIS

168 Rock samples were submitted for analysis.

The following analytical package(s) were requested:		Testing Date:
1A2-Kamloops	QOP AA-Au (Au - Fire Assay AA)	2019-10-30 13:20:02
1A3-Kamloops	QOP AA-Au (Au - Fire Assay Gravimetric)	2019-10-31 17:40:03
1E3-Kamloops	QOP AquaGeo (Aqua Regia ICPOES)	2019-10-20 23:50:58
Sieve Report-Kamloops Internal	Sieve Report Internal	2019-10-21 15:55:38

REPORT **A19-13444**

This report may be reproduced without our consent. If only selected portions of the report are reproduced, permission must be obtained. If no instructions were given at time of sample submittal regarding excess material, it will be discarded within 90 days of this report. Our liability is limited solely to the analytical cost of these analyses. Test results are representative only of material submitted for analysis.

Notes:

If value exceeds upper limit we recommend reassay by fire assay gravimetric-Code 1A3

Values which exceed the upper limit should be assayed for accurate numbers.

CERTIFIED BY:

Emmanuel Eseme , Ph.D. Quality Control Coordinator

ACTIVATION LABORATORIES LTD.

9989 Dallas Drive, Kamloops, British Columbia, Canada, V2C 6T4 TELEPHONE +250 573-4484 or +1.888.228.5227 FAX +1.905.648.9613 E-MAIL Kamloops@actlabs.com ACTLABS GROUP WEBSITE www.actlabs.com

Analyte Symbol	Au	Ag	Cd	Cu	Mn	Мо	Ni	Pb	Zn	Al	As	В	Ва	Ве	Bi	Ca	Со	Cr	Fe	Ga	Hg	K	La
Unit Symbol	ppb	ppm	%	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	%	ppm	ppm	%	ppm							
Lower Limit	5	0.2	0.5	1	5	1	1	2	2	0.01	2	10	10	0.5	2	0.01	1	1	0.01	10	1	0.01	10
Method Code	FA-AA	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP	AR-ICP							
GK19-31	6	< 0.2	1.0	3	436	3	2	169	181	0.23	< 2	< 10	34	< 0.5	< 2	0.03	2	25	1.23	< 10	< 1	0.06	< 10
GK19-32	28	7.6	7.1	8	1320	2	1	682	850	0.23	< 2	< 10	29	< 0.5	3	0.02	1	11	2.02	< 10	3	0.05	< 10
GK19-33	1710	> 100	40.6	3170	63	1	4	> 5000	2730	0.23	56	< 10	< 10	< 0.5	21	< 0.01	69	10	9.04	< 10	< 1	0.13	< 10
GK19-34	364	38.1	5.7	629	64	4	2	4490	450	0.29	13	< 10	< 10	< 0.5	13	0.01	42	16	5.38	< 10	< 1	0.17	< 10
GK19-21 ch	8	< 0.2	< 0.5	8	410	1	< 1	46	47	1.03	< 2	< 10	171	0.7	< 2	1.15	5	12	1.14	< 10	< 1	0.56	22
DNK19-01	19	2.6	0.8	15	1870	< 1	3	21	107	1.17	74	< 10	108	< 0.5	4	6.39	3	5	4.70	< 10	< 1	0.32	< 10
DNK19-02	606	6.6	98.6	66	239	10	5	21	4010	0.45	2230	< 10	11	< 0.5	4	0.32	3	19	3.08	< 10	3	0.24	< 10
DNK19-03	72	1.6	12.6	48	2290	7	38	37	957	1.33	167	< 10	22	0.6	< 2	3.42	11	9	5.37	< 10	< 1	0.68	< 10
DNK19-04	243	2.6	8.5	27	1100	2	11	24	374	1.17	739	< 10		< 0.5	4	2.40	4	14	3.09	< 10	< 1	0.56	< 10
DNK19-05	> 5000	> 100	639	5820	10900	2	11	> 5000	> 10000	0.27	654	< 10	21	< 0.5	3	0.47	4	16	4.93	< 10	4	0.14	< 10
DNK19-06	146	2.8	15.4	36	1910	7	10	47	810	0.88	275	< 10	41	< 0.5	< 2	5.98	3		2.39	< 10	< 1	0.46	< 10
DNK19-07	> 5000	87.9	531	2000	892	3			> 10000	0.51	866	< 10	< 10	< 0.5	9	0.63	2	14	2.98	< 10	2	0.26	< 10
DNK19-08	776	14.9	32.3	43	824	5	9	366	1830	0.84	527	< 10	22	< 0.5	3	2.42	5	30	2.76	< 10	< 1	0.42	< 10
DNK19-09	86	10.2	11.3	55	933	8	13	102	627	1.16	289	< 10	19	< 0.5	3	2.01	5	12	3.57	< 10	< 1	0.61	< 10
DNK19-10	> 5000	> 100	378	770	1660	< 1	2	> 5000	> 10000	1.35	1860	< 10	< 10	< 0.5	< 2	1.35	4	13	4.02	< 10	2	0.65	< 10
DNK19-11	684	88.2	35.0	108	1410	4	1	882	1030	0.69	360	< 10	57	< 0.5	< 2	3.83	< 1	14	2.99	< 10	< 1	0.30	< 10
DNK19-11	9	0.8	1.3	109	995	< 1	6	22	157	4.05	26	< 10	130	0.5	< 2	3.36	15	3	8.79	< 10	3	0.39	12
DNK19-12	8	1.1	1.2	33	907	< 1	5		120	2.71	24	< 10	158	0.8	< 2	3.27	15	4	5.67	< 10		0.59	20
DNK19-13	123	15.3	6.3	41	907		9		292	0.72	117	_	67	< 0.5		3.24		13			< 1	0.32	< 10
	323				1430	2	7				610	< 10	55		< 2	4.53	< 1 2	16	3.15	< 10	< 1		< 10
DNK19-15	323 7	44.6	8.0	118			7	1030	304	0.52		< 10		< 0.5	< 2				3.34	< 10	< 1	0.20	
DNK19-16		< 0.2	0.6	4	1650	2	26	12	75	0.36	30	< 10	43	< 0.5	< 2	3.92	2	33	3.00	< 10	< 1	0.17	< 10
DNK19-17	12	0.2	< 0.5	18	1640	< 1	4	13	77	1.75	15	< 10	116	< 0.5	6	4.71	11	9	3.42	< 10	< 1	0.25	< 10
DNK19-18	69	< 0.2	1.5	11	2070	< 1	6	12	69	2.24	227	< 10	172	< 0.5	< 2	8.86	13	2	5.62	< 10	< 1	0.33	10
GTKM19-21	< 5	0.3	7.6	20	4050	8	5	96	1450	2.27	84	13	58	1.0	< 2	8.12	10	8	4.70	< 10	< 1	0.58	< 10
GTKM19-22	< 5	0.8	3.8	21	5690	67	1	22	162	1.49	99	< 10	25	0.5	< 2	7.77	7	3	4.47	< 10	< 1	0.41	12
GTKM19-23	< 5	0.3	6.2	9	5170	2	< 1	11	417	2.76	15	11	99	1.5	< 2	> 10.0	5	4	2.30	< 10	< 1	0.17	20
GTKM19-24	< 5	0.6	1.2	6	4680	3	< 1	45	373	0.38	225	< 10	34	< 0.5	< 2	0.25	2	4	5.77	< 10	< 1	0.31	41
GTKM19-25	< 5	4.9	0.6	36	31100	< 1	81	8	178	2.25	11	< 10	86	< 0.5	12	9.78	15	15	4.54	< 10	1	0.08	15
GTKM19-26	< 5	4.8	3.1	16	850	5	3	83	388	0.49	54	< 10	64	< 0.5	< 2	0.81	8	9	1.94	< 10	< 1	0.34	18
GTKM19-27	< 5	< 0.2	< 0.5	15	660	4	6	< 2	79	1.87	< 2	< 10	101	< 0.5	< 2	0.19	15	9	3.56	< 10	< 1	1.39	15
GTKM19-28	< 5	0.5	9.8	51	2100	< 1	5	37	616	2.11	2	< 10	138	< 0.5	< 2	4.40	15	18	5.01	< 10	< 1	0.17	12
GTKM19-29	< 5	< 0.2	< 0.5	76	1620	2310	19	< 2	105	2.35	4	< 10	103	< 0.5	< 2	4.79	11	20	4.79	< 10	< 1	0.61	< 10
GTKM19-30	< 5	7.4	12.8	23	194	6	2	2450	1310	0.72	22	< 10	76	0.6	< 2	0.30	4	10	1.30	< 10	< 1	0.48	14
GTKM19-31	6	33.4	9.5	367	596	13	1	298	635	0.75	29	< 10	125	< 0.5	< 2	0.30	6	15	1.22	< 10	< 1	0.72	19
GTKM19-32	< 5	3.0	23.3	30	6640	10	4	468	2890	0.78	26	10	122	< 0.5	< 2	7.09	9	3	3.99	< 10	3	0.54	< 10
GTKM19-33	22	1.5	< 0.5	4	222	81	14	24	27	0.51	115	< 10	< 10	< 0.5	3	0.26	< 1	7	15.1	< 10	< 1	0.35	< 10
GTKM19-34	6	13.2	104	54	6000	53	15	1970	> 10000	0.33	325	< 10	37	< 0.5	< 2	8.43	13	7	2.81	< 10	9	0.09	< 10
GTKM19-35	< 5	1.5	2.2	9	88	12	2	175	174	0.88	16	16		< 0.5	< 2	0.07	2	11	0.87	< 10	< 1	0.67	23
GTKM19-36	5	98.1	13.5	22	126	146	< 1	> 5000	804	0.13	269	< 10	11	< 0.5	< 2	0.04	3	10	0.90	< 10	38	0.12	< 10
GTKM19-37	< 5	0.6	< 0.5	21	2810	1	4	44	248	1.75	< 2	19	164	1.2	< 2	4.04	6	8	5.24	< 10	< 1	0.98	18
GTKM19-38	< 5	14.5	15.0	26	788	1	3	290	1160	1.00	15	13	90	0.7	< 2	1.12	9	8	1.49	< 10	< 1	0.54	15
GTKM19-39	< 5	7.6	1.5	29	2350	2	3	1820	285	0.95	29	14	37	0.6	3	4.01	16	11	1.90	< 10	< 1	0.61	< 10
GTKM19-40	< 5	12.6	10.0	19	5130	2	1	> 5000	713	0.33	21	< 10	58	< 0.5	< 2	7.78	9	8	1.24	< 10	2	0.24	< 10
GTKM19-41	6	24.5	11.5	45	363	3	4	2350	1210	1.21	100	17	< 10	0.9	< 2	0.54	5		10.2	< 10	2	0.78	< 10
GTKM19-42	< 5	6.3	14.7	15	2170	< 1	3	700	1570	0.75	33	< 10	91	< 0.5	< 2	1.15	12	8	3.83	< 10	1	0.43	13
GTKM19-43	< 5	9.9	7.6	19	3850	< 1	2	1010	854	0.92	103	< 10	108	0.6	4	1.75	11	8	3.87	< 10	< 1	0.63	20
GTKM19-44	< 5	9.6	8.9	17	783	2	1	936	1010	0.56	66	< 10		< 0.5	3	0.18	12	7	2.43	< 10	< 1	0.39	12
SK19-04	< 5	2.9	12.6	140	5600	17	1	65	810	0.46	35	< 10		< 0.5	2	> 10.0	16	3	2.31	< 10	2	0.32	11
SK19-05	< 5		< 0.5	4	2090	2	3		144	1.41	16	21	115	0.8	< 2	3.32	5		5.61	< 10	< 1	0.81	19
SK19-06	8		22.7	247	1250	25	1	1560	1310	0.33	355	< 10		< 0.5	< 2	0.47	3		5.71	< 10		0.15	< 10
SK19-07	< 5		32.9	88	427	90		> 5000	4030	0.22	488	< 10		< 0.5	< 2	0.02	8		3.36	< 10		0.14	< 10
		i								1	-		1		i e								$\overline{}$

01	Na % 0.001 AR-ICP 0.030 0.054	AR-ICP	S % 0.01 AR-ICP	Sb ppm 2		Sr ppm	Ti %	Th ppm	Te ppm	TI	U	V	W	<u> </u>	Zr ppm	Ag	Pb %	Zn %	Au
01 R-ICP 0.30 0.12 0.02 0.04	0.001 AR-ICP 0.030	0.001 AR-ICP	0.01	_	ppiii	Ibbiii				IT IT ITT			ppm		ILICITI				
0.30 0.12 0.02 0.04	0.030	AR-ICP		_	1	1	0.01	20	1	ppm 2	ppm 10	ppm 1	10	ppm 1	1	ppm 3		0.001	g/tonne 0.03
0.30 0.12 0.02 0.04	0.030			AR-ICP	AR-ICP	AR-ICP	AR-ICP		AR-ICP	AR-ICP	AR-ICP	AR-ICP	_	AR-ICP	AR-ICP	ICP-	ICP-	ICP-	FA-
0.12 0.02 0.04				/ ((10)	7111101	/ 11 101	/ (1 1 1 0 1	/ 11 101	/ 11 101	/ 11 101	/ 11 101	/ 11 101	/ (11 101	/ 11 101	7 11 101	OES	OES	OES	GRA
0.02	0.054	0.002	< 0.01	< 2	< 1	312	< 0.01	< 20	< 1	< 2	< 10	2	< 10	1	< 1				
0.04		0.049	< 0.01	< 2	2	15	< 0.01	< 20	6	< 2	< 10	12	< 10	5	< 1				
	0.065	0.009	0.01	< 2	< 1	6	< 0.01	< 20	< 1	< 2	< 10	5	< 10	< 1	2				
0.01	0.034	0.002	0.08	< 2	< 1	4	< 0.01	< 20	< 1	< 2	< 10	9	< 10	1	2				
0.01	0.021	0.007	10.2	7	< 1	1	< 0.01	< 20	1	< 2	< 10	7	< 10	< 1	4	101	2.15		
0.02	0.034	0.007	5.47	5	< 1	3	< 0.01	< 20	< 1	< 2	< 10	7	< 10	< 1	3				
0.14	0.080	0.041	0.02	< 2	1	137	< 0.01	< 20	2	< 2	< 10	14	< 10	4	2				
1.53	0.044	0.058	0.10	8	3	635	< 0.01	< 20	< 1	< 2	< 10	18	< 10	16	4				
0.09	0.029	0.016	2.21	6	2	28	< 0.01	< 20	8	< 2	< 10	10	< 10	2	2				
0.63	0.048	0.135	2.45	7	10	150	< 0.01	< 20	2	< 2	< 10	44	< 10	12	4				
0.47	0.036	0.057	1.15	5	5	179	< 0.01	< 20	< 1	< 2	< 10	20	< 10	8	2				
0.21	0.043	0.012	1.98	4670	< 1	41	< 0.01	< 20	< 1	< 2	< 10	5	70	4	2	3920	1.91	2.70	19.6
0.55	0.033	0.038	1.17	35	4	555	< 0.01	< 20	< 1	< 2	< 10	30	< 10	7	2				
0.23	0.039	0.022	3.29	1560	2	79	< 0.01	< 20	8	< 2	< 10	16	< 10	3	2		1.91	2.51	6.21
0.33	0.034	0.036	1.47	25	4	132	< 0.01	< 20	2	< 2	< 10	21	< 10	4	2				
0.50	0.048	0.026	1.65	15	5	119	< 0.01	< 20	1	< 2	< 10	22	< 10	5	3				
0.47	0.065	0.086	3.97	716	2	109	< 0.01	< 20	< 1	< 2	< 10	21	17	4	3	794	3.17	1.70	7.53
1.52	0.047	0.083	0.85	49	4	415	< 0.01	< 20	1	< 2	< 10	14	< 10	8	2				
2.42	0.073	0.164	0.04	8	12	210	< 0.01	< 20	< 1	< 2	< 10	137	< 10	8	4				
1.84	0.094	0.151	0.10	4	9	212	< 0.01	< 20	< 1	< 2	< 10	76	< 10	11	3				
0.95	0.035	0.057	0.21	17	4	197	< 0.01	< 20	2	< 2	< 10	36	< 10	8	2				
1.78	0.041	0.045	0.44	46	4	521	< 0.01	< 20	< 1	< 2	< 10	23	< 10	7	4				
0.53	0.027	0.024	0.02	6	2	78	< 0.01	< 20	< 1	< 2	< 10	20	< 10	5	2				
1.12	0.060	0.080	0.03	3	5	512	< 0.01	< 20	3	< 2	< 10	57	< 10	6	2				
2.34	0.051	0.085	0.06	6	7	783	< 0.01	< 20	< 1	< 2	< 10	69	< 10	10	3				
0.63	0.045	0.083	0.60	5	7	259	< 0.01	< 20	< 1	< 2	< 10	78	< 10	9	3				
0.47	0.033	0.097	1.66	3	6	189	0.05	< 20	< 1	< 2	< 10	52	< 10	11	4				
0.27	0.060	0.092	0.08	4	5	203	0.17	< 20	4	< 2	< 10	53	< 10	56	4				
0.21	0.033	0.070	0.97	8	6	22	< 0.01	< 20	< 1	< 2	< 10	8	< 10	10	10				
1.19	0.033	1.32	0.11	4	5	419	0.04	< 20	< 1	22	< 10	61	< 10	27	2				
0.28	0.045	0.063	0.52	9	2	48	< 0.01	< 20	< 1	< 2	< 10	11	< 10	6	6				
0.90	0.051	0.063	0.32	< 2	4	8	0.23	< 20	2	< 2	< 10	58	< 10	3	5				
1.52	0.076	0.085	0.06	4	17	280	0.02	< 20	5	< 2	< 10	186	< 10	9	4				
0.69	0.156	0.110	0.40	2	13	148	0.18	< 20	3	< 2	< 10	135	< 10	10	6				
0.05	0.030	0.088	0.60	8	3	18	< 0.01	< 20	2	< 2	< 10	26	< 10	6	4				
0.05	0.035	0.127	0.33	129	4	27	< 0.01	< 20	< 1	< 2	< 10	28	< 10	10	2				
0.28	0.032	0.083	0.24	7	6	249	< 0.01	< 20	< 1	< 2	< 10	27	< 10	8	2				
0.03	0.025	0.011	17.9	16	< 1	52	< 0.01	< 20	< 1	< 2	< 10	10	< 10	3	23				
0.16	0.021	0.021	1.90	19	< 1	237	< 0.01	< 20	2	< 2	< 10	12	< 10	4	3			1.09	
0.04	0.026	0.042	0.26	5	1	9	< 0.01	< 20	< 1	< 2	< 10	10	< 10	4	3				
< 0.01	0.021	0.009	2.31	155	< 1	7	< 0.01	< 20	< 1	13	< 10	14	< 10	< 1	4		0.600		
0.35	0.045	0.092	0.02	7	11	151	0.15	< 20	4	< 2	< 10	127	< 10	11	8				
0.09	0.034	0.094	0.39	7	4	58	< 0.01	< 20	< 1	< 2	< 10	31	< 10	7	3				
0.06	0.035	0.086	1.07	13	6	129	< 0.01	< 20	1	< 2	< 10	33	< 10	8	3				
0.03	0.027	0.060	0.90	21	5	264	< 0.01	< 20	3	< 2	< 10	11	< 10	9			0.861		
0.05	0.027	0.173	11.3	29	4	29	< 0.01	< 20	< 1	< 2	< 10	32	< 10	8	10				
0.29	0.027	0.146	0.47	10	6	86	< 0.01	< 20	< 1	< 2	< 10	32	< 10	11	3				
0.27	0.034	0.127	0.35	10	7	77	< 0.01	< 20	5	< 2	< 10	38	< 10	10	4				
0.12	0.025	0.091	0.42	7	4	25	< 0.01	< 20	< 1	< 2	< 10	21	< 10	6	3				
0.05	0.029	0.033	1.76	12	6	195	< 0.01	< 20	< 1	< 2	< 10	23	< 10	17	4				
	1.53 0.09 0.63 0.47 0.21 0.55 0.23 0.33 0.30 0.47 1.52 2.42 1.84 0.95 1.78 0.53 1.12 2.34 0.47 0.27 0.21 1.19 0.28 0.90 1.52 0.69 0.05	1.53 0.044 0.09 0.029 0.63 0.048 0.47 0.036 0.21 0.043 0.55 0.033 0.33 0.034 0.47 0.065 1.52 0.047 2.42 0.073 1.84 0.094 0.95 0.035 1.78 0.041 0.53 0.027 1.12 0.060 2.34 0.051 0.47 0.033 0.27 0.060 0.21 0.033 0.28 0.045 0.90 0.051 1.52 0.076 0.05 0.033 0.28 0.045 0.05 0.030 0.05 0.030 0.05 0.030 0.05 0.030 0.025 0.04 0.04 0.026 0.01 0.021 0.03 0.025 </td <td>1.53 0.044 0.058 0.09 0.029 0.016 0.63 0.048 0.135 0.47 0.036 0.057 0.21 0.043 0.012 0.55 0.033 0.032 0.33 0.034 0.036 0.50 0.048 0.026 0.47 0.065 0.086 1.52 0.047 0.083 2.42 0.073 0.164 1.84 0.094 0.151 0.95 0.035 0.057 1.78 0.041 0.045 0.53 0.027 0.024 1.12 0.060 0.080 0.47 0.033 0.097 0.27 0.060 0.082 0.47 0.033 0.097 0.27 0.060 0.092 0.21 0.033 0.070 1.19 0.033 1.32 0.28 0.045 0.083 0.69 0</td> <td>1.53 0.044 0.058 0.10 0.09 0.029 0.016 2.21 0.63 0.048 0.135 2.45 0.47 0.036 0.057 1.15 0.21 0.043 0.012 1.98 0.55 0.033 0.038 1.17 0.23 0.039 0.022 3.29 0.33 0.034 0.036 1.47 0.50 0.048 0.026 1.65 0.47 0.065 0.086 3.97 1.52 0.047 0.083 0.85 2.42 0.073 0.164 0.04 1.84 0.094 0.151 0.10 0.95 0.035 0.057 0.21 1.78 0.041 0.045 0.44 0.53 0.027 0.024 0.02 1.12 0.060 0.080 0.03 2.34 0.051 0.085 0.06 0.47 0.033 0.097</td> <td>1.53 0.044 0.058 0.10 8 0.09 0.029 0.016 2.21 6 0.63 0.048 0.135 2.45 7 0.47 0.036 0.057 1.15 5 0.21 0.043 0.012 1.98 4670 0.55 0.033 0.038 1.17 35 0.23 0.039 0.022 3.29 1560 0.33 0.034 0.036 1.47 25 0.50 0.048 0.026 1.65 15 0.47 0.065 0.086 3.97 716 1.52 0.047 0.083 0.85 49 2.42 0.073 0.164 0.04 8 1.84 0.094 0.151 0.10 4 0.95 0.035 0.057 0.21 17 1.78 0.041 0.045 0.44 46 0.53 0.027 0.024 0.02 <td< td=""><td>1.53 0.044 0.058 0.10 8 3 0.09 0.029 0.016 2.21 6 2 0.63 0.048 0.135 2.45 7 10 0.47 0.036 0.057 1.15 5 5 0.21 0.043 0.012 1.98 4670 <1</td> 0.55 0.033 0.038 1.17 35 4 0.23 0.039 0.022 3.29 1560 2 0.33 0.034 0.036 1.47 25 4 0.50 0.048 0.026 1.65 15 5 0.47 0.065 0.086 3.97 776 2 1.52 0.047 0.083 0.85 49 4 2.42 0.073 0.164 0.04 8 12 1.84 0.094 0.151 0.10 4 9 0.95 0.035 0.057 0.21 17</td<></td> <td>1.53 0.044 0.058 0.10 8 3 635 0.09 0.029 0.016 2.21 6 2 28 0.63 0.048 0.135 2.45 7 10 150 0.47 0.036 0.057 1.15 5 5 179 0.21 0.043 0.012 1.98 4670 <1</td> 41 0.55 0.033 0.038 1.17 35 4 555 0.23 0.039 0.022 3.29 1560 2 79 0.33 0.034 0.036 1.47 25 4 132 0.50 0.048 0.026 1.65 15 5 119 0.47 0.065 0.086 3.97 716 2 109 1.52 0.047 0.083 0.85 49 4 415 2.42 0.073 0.164 0.04 8 12 210	1.53 0.044 0.058 0.09 0.029 0.016 0.63 0.048 0.135 0.47 0.036 0.057 0.21 0.043 0.012 0.55 0.033 0.032 0.33 0.034 0.036 0.50 0.048 0.026 0.47 0.065 0.086 1.52 0.047 0.083 2.42 0.073 0.164 1.84 0.094 0.151 0.95 0.035 0.057 1.78 0.041 0.045 0.53 0.027 0.024 1.12 0.060 0.080 0.47 0.033 0.097 0.27 0.060 0.082 0.47 0.033 0.097 0.27 0.060 0.092 0.21 0.033 0.070 1.19 0.033 1.32 0.28 0.045 0.083 0.69 0	1.53 0.044 0.058 0.10 0.09 0.029 0.016 2.21 0.63 0.048 0.135 2.45 0.47 0.036 0.057 1.15 0.21 0.043 0.012 1.98 0.55 0.033 0.038 1.17 0.23 0.039 0.022 3.29 0.33 0.034 0.036 1.47 0.50 0.048 0.026 1.65 0.47 0.065 0.086 3.97 1.52 0.047 0.083 0.85 2.42 0.073 0.164 0.04 1.84 0.094 0.151 0.10 0.95 0.035 0.057 0.21 1.78 0.041 0.045 0.44 0.53 0.027 0.024 0.02 1.12 0.060 0.080 0.03 2.34 0.051 0.085 0.06 0.47 0.033 0.097	1.53 0.044 0.058 0.10 8 0.09 0.029 0.016 2.21 6 0.63 0.048 0.135 2.45 7 0.47 0.036 0.057 1.15 5 0.21 0.043 0.012 1.98 4670 0.55 0.033 0.038 1.17 35 0.23 0.039 0.022 3.29 1560 0.33 0.034 0.036 1.47 25 0.50 0.048 0.026 1.65 15 0.47 0.065 0.086 3.97 716 1.52 0.047 0.083 0.85 49 2.42 0.073 0.164 0.04 8 1.84 0.094 0.151 0.10 4 0.95 0.035 0.057 0.21 17 1.78 0.041 0.045 0.44 46 0.53 0.027 0.024 0.02 <td< td=""><td>1.53 0.044 0.058 0.10 8 3 0.09 0.029 0.016 2.21 6 2 0.63 0.048 0.135 2.45 7 10 0.47 0.036 0.057 1.15 5 5 0.21 0.043 0.012 1.98 4670 <1</td> 0.55 0.033 0.038 1.17 35 4 0.23 0.039 0.022 3.29 1560 2 0.33 0.034 0.036 1.47 25 4 0.50 0.048 0.026 1.65 15 5 0.47 0.065 0.086 3.97 776 2 1.52 0.047 0.083 0.85 49 4 2.42 0.073 0.164 0.04 8 12 1.84 0.094 0.151 0.10 4 9 0.95 0.035 0.057 0.21 17</td<>	1.53 0.044 0.058 0.10 8 3 0.09 0.029 0.016 2.21 6 2 0.63 0.048 0.135 2.45 7 10 0.47 0.036 0.057 1.15 5 5 0.21 0.043 0.012 1.98 4670 <1	1.53 0.044 0.058 0.10 8 3 635 0.09 0.029 0.016 2.21 6 2 28 0.63 0.048 0.135 2.45 7 10 150 0.47 0.036 0.057 1.15 5 5 179 0.21 0.043 0.012 1.98 4670 <1	1.53 0.044 0.058 0.10 8 3 635 < 0.01	1.53 0.044 0.058 0.10 8 3 635 < 0.01	1.53	1.53 0.044 0.058 0.10 8 3 635 < 0.01	1.53 0.044 0.058 0.10 8 3 635 < 0.01	1.53 0.044 0.058 0.10 8 3 635 < 0.01	1.53 0.044 0.058 0.10 8 3 635 < 0.01	1.53 0.044 0.058 0.10 8 3 635 <.0.01	1.53 0.044 0.058 0.10 8 3 6835 < 0.01	1.53 0.044 0.058 0.10 8 3 6.55 <.0.01	1.53 0.044 0.058 0.10 0 8 3 63.5 0.01 < 20 < 1 < 2 < 10 18 < 10 16 4	1.53 0.044 0.058 0.10 0 8 3 638 <0.01 <0.00 <1 < 0.00 <0.029 0.029 0.016 2.11 6 2 28 <0.01 <0.00 <1 < 0.020 <0.00 <0.029 0.029 0.016 2.11 6 2 28 <0.01 <0.00 <1 <0.00 <0.029 0.029 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Analyta Cymbol	Ι	٨	Cd	Cu	Mn	Mo	Ni	Dh	l _{7n}	IAI	١٨٥	В	ID ₀	D ₀	Bi	Co	Co	Cr.	Eo	Co	Пα	к	10
 	Au	9				_		_	Zn	Al	_		Ba	Be		Ca		Cr	Fe %		Hg		La
Unit Symbol	ppb	_	ppm	ppm	ppm	ppm	ppm		ppm	%		ppm	ppm	ppm	ppm	%	ppm	ppm			ppm		ppm
Lower Limit	5		0.5	AR-ICP	5	AR-ICP	AD IOD		2 AR-ICP	0.01	2	10	10	0.5	2 AR-ICP	0.01	AR-ICP	AD IOD		10 AR-ICP	AD IOD		10 AR-ICP
Method Code PEK19-17 Orig	FA-AA	AR-ICP < 0.2	AR-ICP < 0.5	32	788	AR-ICP < 1	AR-ICP	4R-1CP	59	AR-ICP 2.34	AR-ICP < 2	< 10	391	< 0.5	AR-ICP < 2	AR-ICP 1.24	10	12	AR-ICP 2.69	< 10	AR-ICP < 1	AR-ICP 0.53	
PEK19-17 Ong		< 0.2	< 0.5	33	797	< 1	5		61	2.34	< 2	< 10		< 0.5	< 2	1.25	10	12	2.09	< 10	< 1	0.53	12
PEK19-17 Dup	< 5	< 0.2	< 0.5	33	191	<u> </u>		3	01	2.30	< 2	< 10	392	< 0.5	\ \ \	1.23	10	12	2.71	V 10	<u> </u>	0.55	13
PEK19-20 Dup	< 5																						
GK19-07 Orig	< 5								-			-											
	< 5 < 5																						
GK19-07 Dup	< 5	0.0	. 0 5	-	1500	10	-			0.11	70	. 10	17	.05	. 0	1 1 1	0		2.00	. 10	. 4	0.40	10
GK19-09 Orig		0.3	< 0.5	5	1560	13	- 1	8	68	2.11	73	< 10	17	< 0.5	< 2	1.11	9	8	3.88	< 10	< 1	0.48	12
GK19-09 Dup		0.3	0.6	6	1580	14	5			2.09	73	< 10		< 0.5	< 2	1.06	9	8	3.95	< 10	< 1	0.48	11
GK19-28 Orig	< 5	< 0.2	< 0.5	22	337	< 1	30		56	1.05	4	< 10		< 0.5	3	2.40	2	31	2.26	< 10	< 1	0.10	< 10
GK19-28 Split PREP DUP	< 5	< 0.2	< 0.5	22	311	< 1	29	14	58	1.01	3	< 10	47	< 0.5	< 2	2.09	2	30	2.16	< 10	< 1	0.09	< 10
GK19-31 Orig	6	< 0.2	1.0	3	442	3	2	171	182	0.24	< 2	< 10	35	< 0.5	< 2	0.03	2	24	1.25	< 10	< 1	0.06	< 10
GK19-31 Dup	6	< 0.2	1.1	3	430	3	1	167	180	0.22	< 2	< 10	33	< 0.5	< 2	0.03	2	26	1.21	< 10	< 1	0.06	< 10
DNK19-06 Orig	147	\ 0.2	1.1		+50		'	107	100	0.22	\ <u>`</u>	\ 10	- 55	\ 0.5	``	0.00		-20	1.21	\ 10		0.00	\ 10
DNK19-06 Dup	147							<u> </u>	-			-	<u> </u>										
DNK19-07 Orig	143							-	-	 		-	 										
DNK19-07 Dup	-							-	-	 		-	 										
DNK19-10 Orig	 	> 100	373	764	1650	< 1	2	> 5000	> 10000	1.33	1820	< 10	< 10	< 0.5	2	1.35	4	12	3.98	< 10	3	0.64	< 10
DNK19-10 Dup		> 100	383	777	1670	1	3	> 5000	> 10000	1.37	1900	< 10	< 10	< 0.5	< 2	1.36	4	14	4.06	< 10	2	0.66	< 10
GTKM19-23 Orig		0.3	6.2	9	5220	2	< 1	11	421	2.81	16	11	101	1.5	< 2	> 10.0	5	4	2.35	10		0.18	20
GTKM19-23 Dup		0.3	6.1	9		1	< 1	12	413	2.72	13	11	98	1.5	< 2	> 10.0	5	4	2.26	< 10	< 1	0.17	20
GTKM19-33 Orig	21	0.5	0.1	3	3110			12	410	2.12	10	- ''	30	1.5	\ _	> 10.0	3	-	2.20	V 10	\ '	0.17	20
GTKM19-33 Dup	23																						
GTKM19-38 Orig		14.7	15.1	26	797	1	3	296	1170	1.00	15	13	89	0.7	< 2	1.14	10	8	1.51	< 10	< 1	0.54	15
GTKM19-38 Dup		14.3	14.9	25	778	1	2	285	1140	1.00	15	12		0.7	< 2	1.10	9	8	1.47	< 10	< 1	0.54	15
GTKM19-43 Orig	< 5	1 1.0	1 1.0		770	·	_	200	11.10	1.00	10		02	0.7	`-	1.10	Ů	Ŭ	1.17	V 10	` '	0.01	- 10
GTKM19-43 Dup	< 5																						
SK19-04 Orig	< 5	2.9	12.6	140	5600	17	1	65	810	0.46	35	< 10	17	< 0.5	2	> 10.0	16	3	2.31	< 10	2	0.32	11
SK19-04 Split	< 5	2.9	12.0	137	5470	18	1	62	760	0.57	33	< 10		< 0.5	4	> 10.0	16	3	2.37	< 10	2	0.40	11
PREP DUP																							
SK19-10 Orig		> 100																					
SK19-10 Dup		82.2																					
SK19-01A Orig		3.4	23.4	25	3770	10	4	162	2300	1.42	73	30	164	0.6	< 2	6.75	13	3	2.29	< 10	3	0.77	14
SK19-01A Dup		3.5	23.9	25	3820	5	4	166	2270	1.47	75	31	166	0.6	< 2	6.80	13	2	2.35	< 10	2	0.80	14
Method Blank		< 0.2	< 0.5	< 1	< 5	< 1	< 1	< 2	< 2	< 0.01	< 2	< 10	< 10	< 0.5	< 2	< 0.01	< 1	< 1	< 0.01	< 10	< 1	< 0.01	< 10
Method Blank		< 0.2	< 0.5	< 1	< 5	< 1	< 1	< 2	< 2	< 0.01	< 2	< 10	< 10	< 0.5	< 2	< 0.01	< 1	< 1	< 0.01	< 10	< 1	< 0.01	< 10
Method Blank		< 0.2	< 0.5	< 1	< 5	< 1	< 1	< 2	< 2	< 0.01	< 2	< 10	10	< 0.5	< 2	< 0.01	< 1	< 1	< 0.01	< 10	< 1	< 0.01	< 10
Method Blank																							
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Analyte Symbol	Mg	Na	Р	s	Sb	Sc	Sr	Ti	Th	Те	TI	U	V	W	Υ	Zr	Ag	Pb	Zn	Au
Unit Symbol	%	%	%	%	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	%	g/tonne
Lower Limit	0.01	0.001	0.001	0.01	2	1	1		20	1	2	10	1	10	1	1	3	0.003	0.001	0.03
Method Code	AR-ICP	AR-ICP			AR-ICP	AR-ICP	AR-ICP	AR-ICP		AR-ICP	AR-ICP	-	AR-ICP	AR-ICP	AR-ICP	AR-ICP	ICP- OES	ICP- OES	ICP- OES	FA- GRA
PEK19-01 Split PREP DUP	0.80	0.105	0.077	< 0.01	4	5	822	0.22	< 20	6	< 2	< 10	54	< 10	8	7				
PEK19-06 Orig	0.53	0.097	0.027	< 0.01	< 2	2	74	0.03	< 20	3	< 2	< 10	25	< 10	2	3				
PEK19-06 Dup	0.53	0.096	0.027	< 0.01	< 2	2	71	0.03	< 20	2	< 2	< 10	25	< 10	2	2				
PEK19-17 Orig	1.25	0.103	0.087	< 0.01	< 2	7	189	0.23	< 20	2	< 2	< 10	88	< 10	8	6				
PEK19-17 Dup	1.29	0.107	0.088	< 0.01	< 2	7	191	0.24	< 20	2	< 2	< 10	87	< 10	8	5				
PEK19-20 Orig																				
PEK19-20 Dup																				
GK19-07 Orig																				
GK19-07 Dup																				
GK19-09 Orig	0.59	0.053	0.083	1.54	2	5	123	0.21	< 20	< 1	< 2	< 10	50	< 10	14	14				
GK19-09 Dup	0.58	0.055	0.082	1.56	4	5	117	0.20	< 20	6	< 2	< 10	50	< 10	14	15				
GK19-28 Orig	0.66	0.043	0.048	0.01	2	2	528	< 0.01	< 20	< 1	< 2	< 10	25	< 10	3	1				
GK19-28 Split PREP DUP	0.63	0.059	0.045	0.01	5	2	466	< 0.01	< 20	< 1	< 2	< 10	23	< 10	3	1				
GK19-31 Orig	0.02	0.067	0.009	0.01	< 2	< 1	6	< 0.01	< 20	< 1	< 2	< 10	5	< 10	< 1	2				
GK19-31 Dup	0.02	0.063	0.009	0.01	< 2	< 1	5	< 0.01	< 20	< 1	< 2	< 10	5	< 10	< 1	2				
DNK19-06 Orig																				
DNK19-06 Dup																				
DNK19-07 Orig																		1.91	2.49	
DNK19-07 Dup																		1.92	2.53	
DNK19-10 Orig	0.46	0.064	0.085	4.02	699	2	115	< 0.01	< 20	< 1	< 2	< 10	21	18	4	3				
DNK19-10 Dup	0.48	0.067	0.086	3.92	733	2	104	< 0.01	< 20	< 1	< 2	< 10	22	15	4	3				
GTKM19-23 Orig	0.27	0.062	0.093	0.08	5	5	204	0.17	< 20	3	< 2	< 10	54	< 10	56	4				
GTKM19-23 Dup	0.27	0.059	0.091	0.07	2	5	203	0.17	< 20	4	< 2	< 10	53	< 10	55	4				
GTKM19-33 Orig																				
GTKM19-33 Dup																				
GTKM19-38 Orig	0.09	0.033	0.096	0.40	7	4	59	< 0.01	< 20	2	< 2	< 10	31	< 10	7	3				
GTKM19-38 Dup	0.09	0.035	0.092	0.39	7	4	56	< 0.01	< 20	< 1	< 2	< 10	31	< 10	7	4				
GTKM19-43 Orig																				
GTKM19-43 Dup																				
SK19-04 Orig	0.05	0.029	0.033	1.76	12	6	195	< 0.01	< 20	< 1	< 2	< 10	23	< 10	17	4				
SK19-04 Split PREP DUP	0.05	0.029	0.035	1.70	11	7	192	< 0.01	< 20	< 1	2	< 10	27	< 10	16	3				
SK19-10 Orig																				
SK19-10 Dup																				
SK19-01A Orig	0.16	0.049	0.124	0.27	5	7	144	0.07	< 20	3	< 2	< 10	64	< 10	14	2				
SK19-01A Dup	0.16	0.050	0.126	0.26	5	7	145	0.08	< 20	4	< 2	< 10	66	< 10	14	2				
Method Blank	< 0.01	0.012	< 0.001	< 0.01	< 2	< 1	< 1	< 0.01	< 20	< 1	< 2	< 10	< 1	< 10	< 1	< 1				
Method Blank	< 0.01	0.012	< 0.001	< 0.01	< 2	< 1	< 1	< 0.01	< 20	< 1	< 2	< 10	< 1	< 10	< 1	< 1				
Method Blank	< 0.01	0.014	< 0.001	< 0.01	< 2	< 1	< 1	< 0.01	< 20	< 1	< 2	< 10	< 1	< 10	< 1	< 1				
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